Exploring Sandstone Body in Weak Electrical Resistivity Contrast with AMT Data*

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Abstract

AMT method with the economical and convenient superiority plays a key role in exploring the sandstone-type uranium deposits in China, which mainly meets to four problems such as the thickness of overlying strata, delineating the shape of the significant sand bodies, whether there are buried structures and knowing the basement relief. Exploring the sand body shape is the key one among such problems because sand body provides the room of uranium deposits and is the prerequisite for exploring uranium mineralization. Through an example of outlining the sandstone layer within the mudstone layers, the ability can be improved to recognize the electrical resistivity anomaly among the weak electrical property contrast by adjusting the inversion model’s scale. A route to deal with the problem was given by inverting different scale models followed by checking whether the anomalies of each inversion are reliable. Finally, the geo-electrical model was to be determined by comparing results of different scale model.

Keywords

AMT, Inversion Model Scale, Recognizing the Electrical Resistivity Anomaly

1. Introduction

Up to 1980’s, natural source EM method had not been applied to shallow geophysical exploration, while the MT method acquiring below several hundred Hz natural EM signals was mainly developed. AMT (the natural-source audiomagnetotelluric method) has become a key geophysical exploration method in several decades.

*AMT, the natural-source audio-magnetotelluric method.
With speedy development of micro-computer and signal processing and emergence of the impedance tensor decomposition, the commercial AMT instruments become mature up to 1990s and are applied for different purposes. With the advantage of economy and convenience as well as the ability to explore shallow objects (less than 2000 meters), many users are attracted to apply AMT to different fields such as ore exploration, engineer surveying and looking for water resources, etc. The more applied, the more techniques were promoted. In the past decades, many researchers were focusing on forward and inversion techniques of EM problems and there were new algorithms, which resulted in the significance of the feasible 2D inversion algorithms in field practices. However, there is little focus on model making. There are only several articles discussing the model making, such as Wannamaker et al. (1985) [1] and Weaver (1994) [2] giving general suggestions.

In the late 20th century, China began to devote exploring and mining leachable stand-stone type uranium Deposits, which enhanced AMT application in this field. AMT method in exploring uranium deposit exploration can give information of four aspects, such as the thickness of overlying strata, delineating the shape of the significant sand bodies, whether there are buried structures and knowing the basement relief. Uranium deposits exist in Erlian basin, the petroleum producer in northern China. Nuheting uranium deposit was discovered many years ago, but several other uranium deposits were found in recent years. To update exploring techniques and extend the area of uranium deposits, CNNC initiated Enlarging Uranium Resource and Evaluation technology Research for Building Mining Base project in 2010, which included several basins such as Erlian basin. Many new geophysical and geological data were acquired in this project to improve understanding geological units setting and determining new uranium deposits. An AMT profile (YC profile) was recorded in 2014, locating at the eastern Erlianhot and the northern Ulanqab depression (Figure 1 and Figure 2). During processing data of this AMT profile, using general data processing method could not identify the significant stand stone body because there are much thicker mudstone layers underground.

The particular aims of this paper are to present an AMT application of recognizing sandstone body in the weak resistivity contrast at Erlian basin of China and to give an AMT data inverting route to deal with the problem in such situation.

2. Geologic Setting and Resistivity Well-Logging Curve

2.1. Geologic Setting [3]

Erlian basin was meso-cenozoic inland basin developed in Tianshan-Mongolia-Xingan Variscan fold belts, which are the south-east margin of the Central Asia Mongolia geosynclinals fold belt.

Basin base is composed of the Proterozoic or Paleozoic metamorphic rock series and intermediate acid intrusive rocks of Variscan and Yanshan period. The cover of this basin is mainly comprised of Cretaceous formation, which includes
Figure 1. Geotectonic elements diagram of Erlian basin (after No. 208 Geological Party, CNNC. and Liu et al. (2013)).

Figure 2. Site map (1, Sand, clay etc.; 2, sand gravel; 3 sandy mudstone; 4, sand stone, mudstone; 5, sandy conglomerate, mudstone; 6, mudstone; 7, mudstone, sand stone; 8, mud stone, siltstone; 9, sandy conglomerate, mud stone, siltstone; 10, granite; 11, granodiorte; 12, borehole; 13, uranium occurrence; 14, AMT surveying profile) (after No. 208 Geological Party, CNNC. and Liu et al. (2013)).
Aershan formation (K₁a), Tenggeer formation (K₁t), Saihan formation (K₁s) and Erlain formation (K₂e) from bottom to top. The major tectonic line is in E-W direction.

Aershan formation (K₁a) was composed of coarse detrital sedimentary formed in stage of strongly pulling apart. The rocks of the lower Tegeer formation are pebbled sandstone, or siltstones with mudstone, while the upper’s are mudstones. The rocks of Saihan formation are sandstones intercalated by mudstones, lignite or pebbly mudstones with sandstones and the sandstones of this formation provided the room of uranium ore-bodies. The rocks of the lower Erlian formation are sandy conglomerates, sandstone or mudstone and so on, while the upper’s are mudstones.

### 2.2. Resistivity Well-Logging Curve

There is the borehole E14-7 at a distance 9200 m on the profile, and its well logging resistivity curve and lithological column are plotted in Figure 3. For simple and convenient purposes, here only part of the profile is discussed from 8000 m to 9800 m. In Figure 3, the resistivity values of the sandstones are little more than that of the mudstones from the depth 165 m to 230 m, although resistivity values of the well logging curve are very lower especially at the depths from 30 m to 165 m and below 275 m only approximately 8 Ω·m. At the depths where the resistivity curve varies steeply such as about 180 m, 230 m and 275 m there generally exit pebbled sandstone, or siltstones with mudstone.

### 3. AMT Data Acquiring and Processing

A 27-km long YC profile in NNW direction (Figure 2) was recorded in 2014.

![Figure 3. Lithological column and resistivity well-logging curve of borehole E14-7 (after field well logging data and Liu et al. (2013)).](image)
with V8 multi-function receiver the AMTC-30 induction coils of Phoenix Geophysics Limited.

The AMT station space is 100 meters and the electrode spacing is 50 meters with tensor layout using non polarized electrodes. Little human interference is conducive to AMT data recording. In order to further improve data quality, each station’s data were recorded for over 25 minutes.

The first step to process the AMT data is to convert each station’s time serial data to the frequency domain before calculating the cross power spectrum. Then, the estimation of each station’s impedance is followed by seeking a corresponding resistivity model which also satisfies a specified regularization function through inversion. The time serial data stacked by long time recording provide plenty of geo-electric field information. The impedance of each station is estimated through the robust method by TSP program’s kernel [4]. Sample data curves of apparent resistivity of the profile are showed in Figure 4. These curves are descending with frequency decrease and are very smooth for most stations. Fortunately, the “dead frequency band [5]” shows very weakly here. In Figure 4, due to the small resistivity value differences between TE and TM and the similar shapes of the TE and TM, it’s inferred that the surveying area is mainly one-dimensional and the inverting AMT with 2D technique can meet the demand. The resistivity values estimated from AMT data are less than 20 ohm-meters for most stations and their various range is within several ohm-meters among the whole frequencies. The lithology described above agrees with the characteristics of these resistivity curves.

4. 2D Inversion

Developed by Randy Mackie, Inversion software (WinGLink®) finds regularized
solutions (Tikhonov Regularization) to the two-dimensional inverse problem for magnetotelluric data using the method of nonlinear conjugate gradients (William, et al., 2001) [6]. The forward model simulations are computed using finite difference equations generated by network analogs to Maxwell’s equations (Swift and JR., 1971) [7]. The program inverts for a 2D rectangular mesh of resistivity blocks, extending laterally and downwards beyond the central detailed zone, and incorporating topography (WinGLink® User’s Guide) [8]. The shallow meshes are of the same height, and from certain depth the mesh’s heightens with some spreading proportional factor.

There is a difference between the inverted geo-electrical model (Figure 5(A)) and the well-logging curve near the E14-7 borehole at the depth between 165 m and 230 m although the inversion misfit is less than 5% and the observed resistivity curves are similar with the response of the inverted model (Figure 7). Because of the coherence relationship between electrical anomaly and the model used in the inversion, a coarse mesh may integrate the anomaly parts and the host in one cell, which result in mixing their differences especially within very small contrast of this example. The paper solved this problem by compiling the characteristics of AMT data and WinGLink software. First, getting a generally underground resistivity distribution with the inverted coarse model produced by

![Figure 5](image.png)

**Figure 5.** Resistivity models of 2D inversion with different scale models. The red cure is resistivity well-logging curve. The rectangle delineates the high resistivity part, and the red circle presents the resistivity value from high to low with increase of the depth.
WinGLink. Then, projecting this resistivity distribution to a fine mesh produced from a coded program and inverting it. The thinking of this program is similar to the route of producing a fine mesh of the WinGLink software. However, there are significant differences between them especially at the algorithm. The height or width of a cell can be set respectively, and there are no more adding elements beyond the maximum detecting depth calculated from the whole data. Trial and error approach is used to balance total element number and model scale, while ensuring that cells are kept sufficiently fine for accurate numerical calculation. By inverting YC profile AMT data with a new model produced by this method, a new geo-electrical model was obtained in the Figure 5(B) with 5% of the inversion misfit.

There is a resistivity layer near the depths from 165 m to 275 m in Figure 5(B), which generally agrees with the well logging curve resistivity anomaly part at the same depth range where the curve’s resistivity values are only several ohmmeters higher than the top and the below. In order to check whether this anomaly layer exists, two ways of constraining the area of the anomaly inversions were used. One, model C in Figure 6, is set with lower resistivity values than that of model B (Figure 5); the other is the higher one as model D in Figure 6. The response of the inverted constrained models and no-constrained model and observed curves are plotted near the borehole E14-7 as Figure 7.

Figure 6. Constraint inversion results. The white rectangles with white dots are constrained with different resistivity values. The constrained resistivity value in model C is 3 Ω·m, while 50 Ω·m in model D.
There is a clear bias between the constrained model and unconstrained model from the observed curves. There is a resistivity anomaly layer exiting at the depths from 165 m to 275 m near the borehole indeed.
5. Conclusions

Studying key sandstone body’s shape is a key task for exploring the leachable sandstone type uranium deposits, for the sandstone body is the room where the ore may exist. Many sandstone bodies are surrounded by the conductive mudstone layers, which add the difficulty in exploring such sandstone bodies. Especially, the low resistivity contrast adds the difficulty in imaging such anomaly.

The analysis above shows that the sandstone layer or body (a little higher resistivity layer), which agrees with the well logging data and is within the conductive layers, was imaged through adjusting the model’s scale by the identical inversion algorithm. It’s inferred that the different scale the models indeed influenced the result of the inversion and adjusting model’s scale to invert data are a key step for AMT data inversion especially in such little contrast environment. So, it’s recommended to calculate more times with two steps to meet such situation. The first step is to invert different scale models, then to check the anomalies of each inversion by the above way or the known information. This inversion route can defend missing some anomalies hidden in the data, especially for the weak resistivity contrast condition.

References


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