Relationship between the Atmospheric CO\textsubscript{2} and Climate Indices by Wavelet-Based Multifractal Analysis

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Abstract

Atmospheric concentrations of greenhouse gases are rising, leading to a positive radiative forcing of climate and an expected warming of surface temperatures. In general, fractal properties may be observed in the time series of the dynamics of complex systems. To study the relation between the atmospheric CO\textsubscript{2} concentration and the climate indices, we investigated the change of fractal behavior of the CO\textsubscript{2}, the carbon isotope ratio ($\delta^{13}$C) of atmospheric CO\textsubscript{2}, the El Niño-Southern Oscillation (ENSO), the Pacific Decadal Oscillation (PDO), and the North Atlantic Oscillation (NAO) indices using the multifractal analysis. When the atmospheric CO\textsubscript{2} growth rate was large, the multifractality of CO\textsubscript{2}, $\delta^{13}$C in CO\textsubscript{2}, ENSO, and NAO was large and the changes were large from the change of fractality. The changes of CO\textsubscript{2} and ENSO were closely related and the influence of the CO\textsubscript{2} on the ENSO was strong from the change in fractality and wavelet coherence. When the El Niño occurred, the CO\textsubscript{2} growth rate was large. The CO\textsubscript{2} related to PDO, NAO, and global temperature from the change in fractality and wavelet coherence. Especially, the changes of CO\textsubscript{2} and global temperature were closely related. When the global warming hiatus occurred, the multifractality of the global temperature was weaker than that of CO\textsubscript{2} and the change of the global temperature was stable. These findings will contribute to the research of the relation between the atmospheric CO\textsubscript{2} and climate change.

Keywords

Atmospheric CO\textsubscript{2}, $\delta^{13}$C, ENSO, PDO, NAO, Wavelet, Multifractal

1. Introduction

Atmospheric concentrations of greenhouse gases are rising, leading to a positive
radiative forcing of climate and an expected warming of surface temperatures.

Many of the El Niño events appear to be associated with a net transfer of CO₂ from the biosphere to the atmosphere. ENSO events produce droughts, particularly in tropical regions, which can cause drought and forest fires and increase CO₂ (Keeling et al., 1995; Langenfelds et al., 2002).

The North Atlantic Oscillation (NAO) and the Pacific-North America (PNA) indices are correlated with the observed atmospheric CO₂ growth rate (Maruyama et al., 2004).

Various objects in nature show the so-called self-similarity or fractal property. Monofractal shows a roughly similar pattern at different scales and is characterized through a fractal dimension. Multifractal is a non-uniform, more complex fractal and is separated into many sub-sets characterized through different fractal dimensions. Fractal property can be observed in the time series representing dynamics of complex systems as well. A change of fractality occurs with a phase transition and changes of state. The multifractal properties of daily rain were studied in two contrasting climates: an East Asian monsoon climate with drastic rain variability and a mild climate with moderate rain variability (Svensson et al., 1996). In both the climates, the frontal rain shows monofractality and the convective-type rain shows multifractality.

Hence, climate change can be interpreted from the view of fractals. A change of fractality may be observed when the climate changes. We use the wavelet transform to analyze the multifractal behavior of the climate index. Wavelet methods are useful for the analysis of complex non-stationary time series. The wavelet transform allows good multifractal analysis to be performed (Muzy et al., 1991). We used the wavelet transform to analyze the multifractal behavior of the climate index. We concluded that a climatic regime shift corresponds to a change from multifractality to monofractality of the Pacific Decadal Oscillation (PDO) index (Maruyama et al., 2015).

To study the relation between the atmospheric CO₂ concentration and the climate indices, we investigated the change of multifractal behavior of the CO₂, the carbon isotope ratio (δ¹³C) of atmospheric CO₂, the Southern Oscillation Index (SOI), and the NAO index using the multifractal analysis.

2. Data and Method of Analysis

The SOI, Niño3.4, PDO, NAO indices, and global mean surface air temperature anomalies provided by NOAA’s Climate Prediction Center, USA (CPC) were used. The SOI is a standardized index based on the observed sea level pressure differences between Tahiti and Darwin, Australia. The SOI is one measure of the large-scale fluctuations in air pressure occurring between the western and eastern tropical Pacific (i.e., the state of the Southern Oscillation) during El Niño and La Niña episodes. Prolonged periods of negative (positive) SOI values coincide with abnormally warm (cold) ocean waters across the eastern tropical Pacific typical of El Niño (La Niña) episodes.
Atmospheric CO₂ concentrations (ppm) derived from in situ air measurements at Mauna Loa, Observatory, Hawaii was used. Monthly atmospheric ¹³C concentrations (per mil) in CO₂ derived from flask air samples at Mauna Loa Observatory was used. Annual mean growth rate of CO₂ for Mauna Loa obtained from NOAA was used.

For the examination, we used the Daubechies wavelet, which is widely used in solving a broad range of problems, e.g., self-similarity properties of a signal and signal discontinuities. We made use of a discrete signal that was fitted the Daubechies mother wavelet with the capacity of correct inverse transformation. Thus, we can precisely calculate the following best \( \tau(q) \), which can be regarded as a characteristic function of the fractal behavior. We can define the \( \tau(q) \) from the power-law behavior of the partition function, as shown in equation (2). We then computed the scaling of the partition function \( Z_q(a) \), which is defined as the sum of the \( q \)-th powers of the modulus of the wavelet transform coefficients at scale \( a \), where \( q \) is the \( q \)-th moment. In our computation, the wavelet-transform coefficients did not grow zero. Thus, for a correct calculation, the summation was considered for the whole set. Muzy et al. (1991) defined \( Z_q(a) \) as the sum of the \( q \)-th powers of the local maxima of the modulus to avoid dividing by zero. We got the following partition function \( Z_q(a) \):

\[
Z_q(a) = \sum W_q[f](a,b)^q
\]  

where \( W_q[f](a,b) \), \( a \) and \( b \) are the wavelet coefficient of function \( f \), a scale parameter, and a space parameter, respectively. \( W_q[f](a,b) \) is defined as below.

\[
W_q[f](a,b) = \frac{1}{\sqrt{\text{det} \Phi}} \int_{-\infty}^{\infty} f(t) \phi^q \left( \frac{t-b}{a} \right) dt
\]  

where \( f(t) \) is data and \( \phi \) is wavelet function. For small scales, we expect

\[
Z_q(a) \sim a^{\tau(q)}
\]

First, we examined the changes in \( Z_q(a) \) in the time series at a different scale \( a \) for each moment \( q \). We plotted the logarithm of \( Z_q(a) \) against that of time scale \( a \). Here \( \tau(q) \) is the slope of the fitted straight line for each \( q \). Next, we plotted \( \tau(q) \) versus \( q \). The time window was advanced by one year, which was repeated. The time window was fixed to 6 years, when a moderate change in fractality was observed. Monofractal and multifractal signals were defined as follows. For \( \tau(q) \), a monofractal signal corresponds to a straight line, while a multifractal signal is nonlinear (Frisch & Parisi, 1985). We calculated the \( R^2 \) value, which is the coefficient of determination, for the fitted straight line. If \( R^2 \geq 0.98 \), the time series is monofractal and if \( 0.98 > R^2 \), it is multifractal.

We calculated the \( \tau(q) \) of different moments \( q \) for individual records in the Niño3.4 index. In Figure 1, the \( \tau(q) \) between 1980 and 1994 is shown. The data were analyzed in six-year sets, e.g., \( \tau(q) \) of n80 was calculated between 1980 and 1985, and that of n81 was calculated between 1981 and 1986. To investigate the change of fractality, the time window was then shifted forward one year and \( \tau(q) \)
Figure 1. $\tau(q)$ for individual Niño3.4 index between 1980 and 1994.

was calculated from n80 up to n89. A monofractal signal would correspond to a straight line for $\tau(q)$, while a multifractal signal would be nonlinear. Nearly all the multifractality observed is due to the negative value of $q$, i.e., small fluctuations are more inhomogeneous than large fluctuations. In Figure 1, the data sets were monofractal in the cases of n80, n81, n82, n85, n86, n87, and n89 and were multifractal in the cases of n83, n84, and n88.

We plotted the value of the $\tau(-6)$ in each index. The negative large values of the $\tau(-6)$ show large multifractality. For the $\tau(q)$, $q = -6$ is the appropriate number to show the change of $\tau$.

3. Results

3.1. The Relation between the Atmospheric CO$_2$ and $\delta^{13}$C in CO$_2$

The time series of atmospheric CO$_2$ and $\delta^{13}$C in CO$_2$ are shown in Figure 2. Changes in $\delta^{13}$C of CO$_2$ can be used in global carbon-cycle models to elucidate the relative roles of oceanic and terrestrial uptake of fossil-fuel CO$_2$ (Francey et al., 1985). The CO$_2$ increased and $\delta^{13}$C decreased with time, hence there was an inverse relationship between CO$_2$ and $\delta^{13}$C, which indicated that these changes were mainly generated by activities of terrestrial biosphere.

We investigated the relation between the CO$_2$ and $\delta^{13}$C. The $\tau(-6)$ of the CO$_2$, and $\delta^{13}$C are shown in Figure 3 (top). The red square shows monofractality and the green circle shows multifractality for the 6 years centered on the year plotted. For instance, the green circle for 1980 in the CO$_2$ shows multifractality between 1977 and 1982. The data was excluded from Figure 3 (top) for cases where we could not distinguish between monofractality and multifractality. The changes in fractality of the CO$_2$ and $\delta^{13}$C were inverse for 1995-2010, which corresponded to the change of the CO$_2$ and $\delta^{13}$C in CO$_2$. When the CO$_2$ growth rate was large as shown in 3.2, the $\tau(-6)$ of $\delta^{13}$C was small and the multifractality was large. The changes of CO$_2$ and $\delta^{13}$C in CO$_2$ were related to each other.
Figure 2. The time series of CO$_2$ and $\delta^{13}$C in CO$_2$.

Figure 3. The $\tau$(−6) of CO$_2$ and $\delta^{13}$C in CO$_2$ (top). Wavelet coherence (middle) and phase (bottom) between the CO$_2$ and $\delta^{13}$C in CO$_2$. In the wavelet phase, the positive value shown by the blue and pink shading means that the CO$_2$ leads the $\delta^{13}$C and the negative value shown by the green, yellow and red shading means that the $\delta^{13}$C leads the CO$_2$. 
We studied the relationship between the CO$_2$ and $\delta^{13}$C by means of wavelet coherence. We show the wavelet coherence and phase using the Morlet wavelet between the CO$_2$ and $\delta^{13}$C in Figure 3 (middle) and (bottom), respectively. The coherence between the CO$_2$ and $\delta^{13}$C in 1 - 2 year scale was very strong for 1980-2010, when the changes of the fractality both were very similar and the phase of CO$_2$ was delayed by $\pi$ from the $\delta^{13}$C.

3.2. The Relation between the CO$_2$ Growth Rate and the Multifractality of CO$_2$

We investigated the relation between the annual mean CO$_2$ growth rate and the multifractality of CO$_2$. The annual mean CO$_2$ growth rate for Mauna Loa and the $\tau(-6)$ of the CO$_2$ are shown in Figure 4. The changes of the CO$_2$ growth rate and the $\tau(-6)$ of CO$_2$ were very similar. When the CO$_2$ growth rate was large, the $\tau(-6)$ of CO$_2$ was small and multifractality was large. When the El Niño (65 - 66, 72 - 73, 82 - 83, 86 - 87, 94 - 95, 97 - 98, 02 - 03, 09 - 10) occurred, the CO$_2$ growth rate was large. When the La Niña (73 - 74, 75 - 76, 99 - 00, 07 - 08) occurred, the CO$_2$ growth rate was small.

3.3. The Relation between the CO$_2$ and ENSO

We investigated the relation between the CO$_2$ and SOI. The $\tau(-6)$ of the CO$_2$, and SOI are shown in Figure 5 (top). The changes in fractality of the CO$_2$ and SOI were very similar for 1965-2010 especially for the 1970s and 1990s. When the CO$_2$ growth rate was large, the ENSO event occurred and the $\tau(-6)$ of SOI was small and multifractal was strong. The changes of CO$_2$ and ENSO were related to each other. We studied the relationship between the CO$_2$ and SOI by means of wavelet coherence. We show the wavelet coherence and phase using the Morlet wavelet between the CO$_2$ and SOI in Figure 5 (middle) and (bottom), respectively. The coherence between the CO$_2$ and SOI in two year scale was strong for 1970-2000, when the changes in the fractality both were very similar and the lead of the CO$_2$ was observed. The proportion of eastern versus central Pacific-type El Niño events increases with CO$_2$ (Stevenson et al., 2012).

Figure 4. The annual mean CO$_2$ growth rate for Mauna Loa (red) and the $\tau(-6)$ of the CO$_2$ (green).
Figure 5. The $\tau(-6)$ of CO$_2$ and SOI (top). Wavelet coherence (middle) and phase (bottom) between the CO$_2$ and SOI. In the wavelet phase, the positive value shown by the blue and pink shading means that the CO$_2$ leads the SOI and the negative value shown by the green, yellow and red shading means that the SOI leads the CO$_2$.

We investigated the relation between the CO$_2$ and Niño3.4. The $\tau(-6)$ of the CO$_2$, and Niño3.4 are shown in Figure 6 (top). The changes in fractality of the CO$_2$ and Niño3.4 were very similar for 1970-1995. When the CO$_2$ growth rate was large, the ENSO event occurred and the $\tau(-6)$ of Niño3.4 was small and multifractal was strong. The changes of CO$_2$ and ENSO were related to each other. We show the wavelet coherence and phase using the Morlet wavelet between the CO$_2$ and Niño3.4 in Figure 6 (middle) and (bottom), respectively. The coherence between the CO$_2$ and Niño3.4 in one year scale was strong for 1960-2010, when the lead of the CO$_2$ was observed.

3.4. The Relation between the CO$_2$ and PDO

We investigated the relation between the CO$_2$ and PDO. The $\tau(-6)$ of the CO$_2$,
and PDO are shown in Figure 7 (top). The changes in fractality of the CO$_2$ and PDO were very similar for 1965-2010, especially for the 1970s and 1990s. The changes of CO$_2$ and PDO were related to each other.

We studied the relationship between the CO$_2$ and PDO by means of wavelet coherence. We show the wavelet coherence and phase using the Morlet wavelet between the CO$_2$ and PDO in Figure 7 (middle) and (bottom), respectively. The coherence between the CO$_2$ and PDO in 1 - 2 year scale was strong for the 1990s, when both the changes in fractality were very similar and the lead of the PDO was observed.

### 3.5. The Relation between the CO$_2$ and NAO

We investigated the relation between the CO$_2$ and NAO. The $\tau(-6)$ of the CO$_2$, and NAO are shown in Figure 8 (top). The changes in fractality of the CO$_2$ and NAO were very similar for 1960-2000. When the CO$_2$ growth rate was large, the $\tau(-6)$ of NAO was small and multifractal was strong. The changes of CO$_2$ and NAO were related to each other.
We studied the relationship between the CO$_2$ and NAO by means of wavelet coherence. We show the wavelet coherence and phase using the Morlet wavelet between the CO$_2$ and NAO in Figure 8 (middle) and (bottom), respectively. The coherence between the CO$_2$ and NAO in one year scale was strong, and the lead of the CO$_2$ and NAO was observed.

### 3.6. The Relation between the Atmospheric CO$_2$ and Global Temperature

We investigated the relation between the CO$_2$ and global temperature. The global mean surface temperature anomalies is shown in Figure 9. The rate of global temperature increase slowed during 1950-1975 and 1998-2012. The change between 1998 and 2012 was often termed the “global warming hiatus” (Medhag et al., 2017). The τ(−6) of the CO$_2$, and global temperature are shown in Figure 10 (top). The changes in fractality of the CO$_2$ and global temperature were very similar for 1970-2000, when the rate of global temperature increased largely. The changes of CO$_2$ and global temperature were related to each other.
Figure 8. The $\tau(−6)$ of CO$_2$ and NAO (top). Wavelet coherence (middle) and phase (bottom) between the CO$_2$ and NAO. In the wavelet phase, the positive value shown by the blue and pink shading means that the CO$_2$ leads the NAO and the negative value shown by the green, yellow and red shading means that the NAO leads the CO$_2$.

Figure 9. Global mean surface temperature anomalies.

We studied the relationship between the CO$_2$ and global temperature by means of wavelet coherence. We show the wavelet coherence and phase using the Morlet wavelet between the CO$_2$ and global temperature in Figure 10.
Figure 10. The \( \tau(\cdot 6) \) of CO\(_2\) and global temperature (top). Wavelet coherence (middle) and phase (bottom) between the CO\(_2\) and global temperature. In the wavelet phase, the positive value shown by the blue and pink shading means that the CO\(_2\) leads the global temperature and the negative value shown by the green, yellow and red shading means that the global temperature leads the CO\(_2\). (middle) and (bottom), respectively. The coherence between the CO\(_2\) and global temperature in two year scale was very strong, when the rate of global temperature increased largely. The lead of the global temperature was observed.

4. Discussion

4.1. The Influence of the \( \delta^{13}\)C on the CO\(_2\) Concentration

The CO\(_2\) increased and \( \delta^{13}\)C decreased with time, so there was an inverse relationship between CO\(_2\) and \( \delta^{13}\)C, which indicated that these changes were mainly generated by activities of terrestrial biosphere. The \( \delta^{13}\)C decreased due to combustion of fossil fuel with isotopically light CO\(_2\) (Nakazawa et al., 1993).

When the CO\(_2\) growth rate was large, the \( \tau(\cdot 6) \) of the CO\(_2\) was small and multifractality was large and the change was large. The changes in fractality of the
CO₂ and δ¹³C were very similar for 1985-2010. When the CO₂ growth rate was large, the τ(−6) of δ¹³C was small and the multifractality was large. The changes of CO₂ and δ¹³C in CO₂ were closely related.

4.2. The Influence of the ENSO, PDO, NAO, and Global Temperature on the CO₂ Concentration

The changes in fractality of the CO₂ and SOI were very similar for 1965-2010. The coherence between the CO₂ and SOI in 2-4 year scale was strong for 1970-2000 and the lead of the CO₂ was observed. The CO₂ differences between the central and western Pacific Ocean correlate well with the SOI. There is more (less) midtropospheric CO₂ in the central Pacific and less (more) midtropospheric CO₂ in the western Pacific during El Niño (La Niña) events (Jiang et al., 2013). When the CO₂ growth rate was large, the ENSO event occurred and the τ(−6) of SOI was small and multifractal was strong. The changes of CO₂ and ENSO were closely related. The influence of the CO₂ on the ENSO (SOI and Niño3.4) was strong from the change of fractality and wavelet coherence. When the El Niño (La Niña) occurred, the CO₂ growth rate was large (small). The relationship between the CO₂ and the SOI values is shown (Matsueda et al., 2015).

The changes in fractality of the CO₂ and PDO were very similar for 1965-2010 especially for the 1970s and 1990s and the changes of CO₂ and PDO were related to each other. The changes in fractality of the CO₂ and NAO were very similar for 1965-2000. The coherence between the CO₂ and NAO in one year scale was strong and the leads of the CO₂ and NAO were observed. When the CO₂ growth rate was large, the τ(−6) of NAO was small and multifractality was strong. The changes of CO₂ and NAO were related to each other. When the rate of global temperature increased largely, the changes in fractality of the CO₂ and global temperature were very similar and the coherence between the CO₂ and global temperature was very strong. For 1998-2012 (global warming hiatus), the multifractality of the global temperature was weaker than that of CO₂. The change of the global temperature was more stable than that of CO₂. The changes of CO₂ and global temperature were closely related. Hence, the CO₂ related to PDO, NAO, and global temperature from the change in fractality and wavelet coherence.

5. Conclusion

To study the relation between the atmospheric CO₂ concentration and the climate indices, we investigated the change of fractal behavior of the CO₂, the carbon isotope ratio (δ¹³C) of atmospheric CO₂, SOI, Niño3.4, PDO, and NAO indices using the multifractal analysis. We showed the change of factuality by plotting the τ-function and used the wavelet coherence. The main findings are summarized below.

1) When the atmospheric CO₂ growth rate was large, the multifractality of CO₂, δ¹³C in CO₂, ENSO, and NAO was large and the changes were large from the change of fractality.

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2) The changes of CO₂ and ENSO were closely related and the influence of the CO₂ on the ENSO was strong from the change in fractality and wavelet coherence. When the El Niño occurred, the CO₂ growth rate was large.

3) The CO₂ related to PDO, NAO, and global temperature from the change in fractality and wavelet coherence. Especially, the changes of CO₂ and global temperature were closely related. When the global warming hiatus occurred, the multifractality of the global temperature was weaker than that of CO₂ and the change of the global temperature was stable.

These findings will contribute to the research of the relation between the atmospheric CO₂ and climate change.

Conflicts of Interest

The author declares no conflicts of interest regarding the publication of this paper.

References


