

Novel Power Law of Turbulent Spectrum

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Abstract

This paper is concerned with novel power law of turbulent energy spectrum and the relevant experiment in tidal current. The power law in the inertial sub-range has been proposed in such a way that the power of the one-dimensional turbulent energy spectrum varies from 0 to -2 approximately, but it is accompanying the small oscillation with increasing the wave number. The well-known Kolmogorov $-5/3$ power law is merely one facet, to appear within the present proposed novel power law. The turbulent energy spectra (S_u , S_v , S_w) in x -, y - and z -directions, respectively oscillate with the wave number. It is found that the turbulent in the tidal currents is three dimensional, and the intermittence of momentum transport is a predominant and characteristic feature in tidal current.

Keywords

Energy Spectrum, Kinetic Theory, Tidal Current, Inertial Sub-Range, Turbulence Measurement, Power Law

1. Introduction

Late Dr. Shunichi Tsugé (1932.4.1-2003.6.21) has elaborated a new theory of turbulence on the basis of Newton mechanics and mathematical probability or statistical mechanics so as to elucidate the motion of a system consisting of N particles with mathematics and computer [1]-[4]. So far, his theory has been successfully applied to boundary layer flow [5], turbulent combustion [6]-[9], mixing layer flow [10], grid-produced turbulence [11]-[14], Bénard convection [15] [16], turbulent energy spectrum [17]. Most crucial feature of Tsugé's theory is that it includes no empirical constant, contrary to almost all the other existing turbulent theories such as Prandtl mixing length theory and $k - \varepsilon$ Model. In this sense, Tsugé's theory is only one true turbulent theory among the many, for soundness and usefulness of his theory have been well confirmed by the relevant experiments [7]-[19]. Major Tsugé's papers on turbulence research have been collected by the senior author [18] [19], which are ac-

cessible at National Diet Library in Tokyo.

Intermittency in turbulence is an essential feature in laboratory [20] [21], so measurements of turbulence have been also conducted in geophysical flow to examine it [22]-[26]. However, even so tremendous efforts to clarify the turbulence phenomena have been continued by many researchers in the world since the end of 19 century by Reynolds [27], still relatively little is known about this. Only success known in this research field except for Tsugés achievements is made by Kolmogorov [28], who obtain a power law of turbulent energy spectrum in the inertial sub-range in terms of dimensional analysis, this is known as $-5/3$ law widely.

It is, therefore, natural to examine what law can be derived by Tsugé' theory, if it is rigorous theory of turbulence currently existing in this globe. Indeed, this prompts us to conduct this study.

The structure of this paper is as follows:

1) Introduction, 2) Liouville equation and BBGKY hierarchy, 3) Navier-Stokes equation is only valid for laminar flow, but not for turbulent flow, 4) Experimental (4.1 Velocity probes, 4.2 Underwater tower, 4.3 Experimental procedure and technique, 4.4 Location and site conditions, 4.5 Results of turbulence measurements, 4.6 Turbulence statistics), 5) Proposition of novel power law of turbulent energy spectrum and its comparison with the experiment, and 6) Conclusion.

The main purpose of the present paper is to derive the novel power law of turbulent energy spectrum, based on the theory and experiment.

2. Liouville Equation and BBGKY Hierarchy

It is well known that the Liouville equation describes the motion of N particles, and it is $6N$ Dimensions. However, problem is that number of particles involved is too many to trace their trajectories. For example, order of 10^{20} molecules is contained in the air of 1 cm^3 or in a water droplet of radius of 1mm. Because information to specify position (x, y, z) and velocity (u, v, w) for each the molecule is required, one must know the position and velocity for N molecules.

There is no question that if we could obtain the solution of Liouville equation completely, it is said that motion of particles is solved, so that turbulence may not be problem anymore. However, this task is so difficult that no genius and no computer have not ever succeeded. Non the less, there is still some room that human beings would solve the Liouville equation describing motion of $N (\rightarrow \infty)$ particles some day in future.

Bogoliubov [29] has proposed so-called "Synchronization Principle" as a practical mathematical idea to overcome the forgoing despairing situation. According to this principle, "since any particle is considered to be equal to each other, except for objective particles, one is allowed to smear out the rest particles by integrating with respect to the position as well as velocity". For example, integration to smear the unnecessary particles ranging from $(s + 1)^{\text{th}}$ to N^{th} may be done in order to obtain the distribution function for s -molecules, and we get it as follows,

$$f_s(Z_1, \dots, Z_s) = \int f_N dZ_{s+1} \dots dZ_N$$

with

$$Z_s \equiv (\mathbf{x}_s, \mathbf{v}_s)$$

In this way, the smeared particles now form the background for the objective particles. Surprisingly, this method has succeeded to reduce number of objective particles from N to 1. This whole procedure is called BBGKY hierarchy theory. BBGKY hierarchy is summarized in **Table 1**, where N is total number of particles.

Tsugé [1] has rigorously proven based on the kinetic theory that two-particle level of Liouville equation is reduced to Kármán-Howarth equation [30], which was derived by them heuristically or phenomenologically. In another words, this means that Kármán-Howarth equation has acquired kinetic theoretical background.

The kinetic theoretical base for Reynolds averaged Navier-Stokes equation, $\langle (NS)_j \rangle = 0$ has been given by Tsugé & Sagara [3]. It is realized that Liouville equation for $N = 1$ particle, $D(f) = J(f_2)$ in the kinetic theory corresponds to Reynolds averaged Navier-Stokes equation in fluid mechanics.

The kinetic theoretical base for Navier-Stokes equation, $(NS)_j = 0$ has been provided by Chapman [31] and Enskog [32] independently in 1917. It has been proven mathematically that Liouville equation for $N = 1$ particle with binary molecular chaos, $D(f) = J(ff)$ with $f_2 = ff$, in the kinetic theory corresponds to Navier-Stokes equation in fluid mechanics. That is, Navier-Stokes equation never includes two-point, or double correlation, so that this equation is not even relevant to turbulence in general. This statement is also supported by the isotropic assump-

The diameter of the sphere is of 42 mm and in a current of 0 - 1.00 m/s it is capable of resolving velocity component fluctuations from 0 to better than 4 HZ. The shape of the probe is shown in this figure.

For a turbulence intensity of 0% - 20%, the effect of inertia on the probe is found, by experiments in a calibration rig, to be insignificant compared to the drag. Assuming the effect of inertia force on the probe is negligible, the voltage readouts from the component strain bridges are proportional to the components of drag force on the sphere. The threshold velocity of the mean velocity component is 0.02 m/s, and for the transverse and vertical components is 0.004 and 0.006 m/s, respectively. The voltages are amplified and low pass filtered to prevent aliasing errors. The velocity probe was calibrated using a reference flow meter that had been calibrated in the laboratory by volumetric measurement. The departure from the theoretical calibration was found to be less than 5%.

4.2. Underwater Tower

The underwater tower [23] which was developed to support the probes is sketched in **Figure 2**.

It consisted of a column formed of two 4 m lengths of 102 mm diameter steel pipe, mounted on a base of three radial arms, each of which was made of two steel angles $127 \times 76 \times 10$ mm, with spikes for stabilizing the legs in position on the sea bed. The outer end of each leg was tied to the top of the column with a cable tensioned by a hand winch. The lower part of the tower had a universal joint for the vertical alignment of the column on different sea bed configurations. Ropes wrapped around the column were used to minimize the vibration due to vortex shedding in flowing water. The damped natural frequency of the tower was about 10 Hz, which was much higher than that of the energy containing eddies. Thus, the effect due to the vibration of the column is considered to be negligibly small.

4.3. Experimental Procedure and Technique

During this study, the field experiments were conducted from "Triton", an 8 m cabin launch, "Hydraulics", a 4.3 m aluminium runabout, and a raft, of the Mechanical Engineering Department, Monash University. Triton housed the minicomputer, power supplies and the other electronic equipment for the data recording system. Hydraulics was used to carry out the topographic survey around the site for the tidal cycle during the data collection, and for transportation between site to site. The raft was made of four 0.2 m³ petrol drums with a small wooden deck and equipped with one-ton hand winch. This raft was used to lift and lower down the three steel legs, which were used to stand and stabilize the underwater column at the site.

4.4. Location and Site Conditions

A total of 15 cruises were conducted in North Arm, a tidal channel in Westernport Bay, Victoria, Australia, as shown in **Figure 3**. These cruises were undertaken over 3 years, viz., from May 1978 and to November 1981,

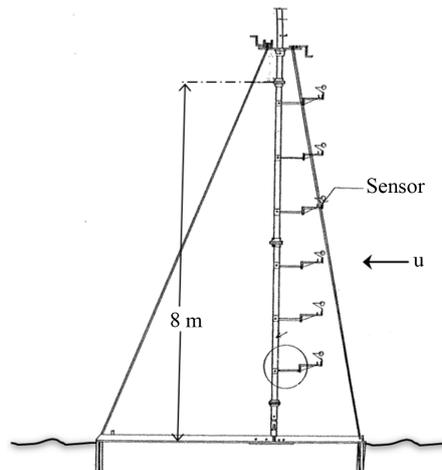


Figure 2. Schematic diagram of the instrument tower.

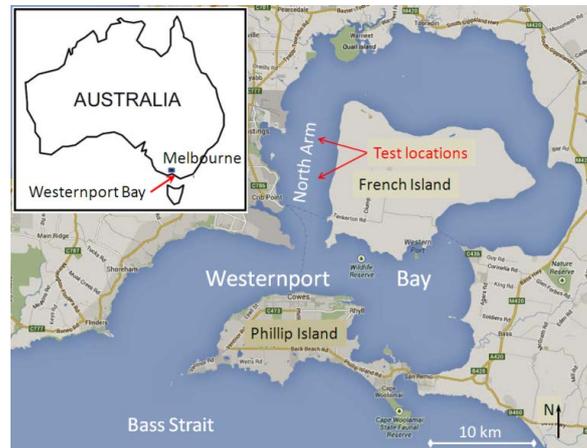


Figure 3. Location of the experimental study. Base map from Google Maps.

following loss of an earlier tower with all equipment. Full utilization of this period was prevented by severe weather, instrument failure and damage, and staff availability.

Four cruises were made in the tidal channel: three in a section of channel with small sand dunes and one in a very rough bed consisting of big and steep sand dunes. Measurements were made of the turbulent velocity fluctuations, bottom topography, wave height, temperature and salinity at different depths. In particular, velocity fluctuation data were collected throughout the tidal cycle of 12.4 hours. Altogether ten 910-second long simultaneous records of the three velocity components (u , v , w) at 6 depths (1, 2, 3, 4, 5 and 6 m above the bed) and thirteen 910-second long records of the three velocity components at respective, 3 depths (1, 2 and 3 m above the bed). Additionally, several 1820-second long records were used to examine the fluctuations having longer period. Mean flow velocities determined from the records at 1, 2, 3, 4, 5, and 6 m above the sea bed.

The effects of wind and surface waves on the flow are negligible, for the channel is protected from the effects of the predominant strong winds and days of data collection were chosen such that the sea was calm and the wind speed less than 5 m/s. Wave height was measured at the instrument tower using a staff gauge, and never exceeded a few centimeters, with wave periods being generally less than 1.0 second. It is quite unlikely that such small waves would cause detectable velocities at the depth of measurement.

Temperature and salinity profiles were made over the depth. In general, the variations of temperature and salinity were found to be random with ranges of 0.2°C and less than 0.1 ppt, respectively. There was no evidence of a thin surface layer having different temperature or salinity. The lack of such gradients precludes the existence of internal waves.

The sea bed at both locations was sandy, with small dunes at one site and large dunes at the other slightly different results were obtained at the two sites, but only the results from the latter site are presented here. In some measurements over a large dune bed there is evidence that the velocity meter fixed 1m above the bed was in the lee of the upstream dune for part of the measurement time.

It is possible that some error may be introduced into the mean flow due to the zero drift of the instruments, mainly due to thermal effects, but it is estimated that the combined error should be less than 5%. This would not affect the measurement of the turbulence.

4.5. Results of Turbulence Measurements

Figure 4 shows a typical time record of the velocity fluctuation components (u , v , w), together with the Reynolds stress uw at 2 m above the sea bed.

Fluctuations of u , v , and w are distributed around zero with intermittent events, while those of uw indicate much larger instantaneous values changing from a positive value to negative one, and vice versa, where u , v , and w are turbulent velocities in the main, transverse and vertical flow directions, respectively. It is worth noting that the present result for the Reynolds stress in this figure adds another piece of evidence that the intermittent phenomena observed in wind tunnel and laboratory boundary layers can be scaled up to dimensions of geophysical interest [37].

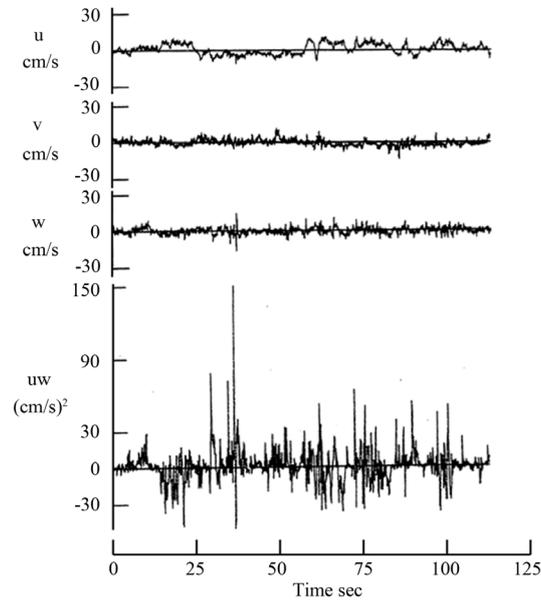


Figure 4. Typical time records of the velocity fluctuation components (u , v , w), together with the Reynolds stress uw at 2 m above the sea bed.

4.6. Turbulence Statistics

The probability distributions, not shown, of the turbulent velocity fluctuations, together with the turbulent intensity and the turbulent scales in the component directions were determined at the six depths of the measurement in the tidal channel. It was found that the probability distributions of each of the velocity components (u , v , w) were approximately Gaussian, while those of the cross product $u \cdot w$ were far from Gaussian having large skewness and kurtosis [23].

It is, therefore, suggested that the intermittency of momentum transport is a predominant and characteristic feature of most geophysical boundary layers, as to be also demonstrated in the present turbulent velocities and energy spectra.

The scatter between records is quite considerable, so no strong systematic variations of turbulent intensity σ_u/U , σ_v/U , σ_w/U , and $\left\langle (uw)^2 \right\rangle^{1/2} / U^2$ are apparent with respect to the velocity range of 0.2 - 0.7 m/s, and height range of 1 - 6 m above the sea bed, though a weak trend towards higher turbulence intensity at locations closer to the sea bed is shown. It must be noted that somewhat higher turbulence intensities are found in case of the rough bottom configuration comparing with the smooth one. The 95% confidence interval for the normalized turbulence intensities for each component are:

$$\sigma_u/U = 8.63 \mp 0.87\%$$

$$\sigma_v/U = 7.47 \mp 0.96\%$$

and,

$$\sigma_w/U = 6.51 \mp 0.80\%$$

These results indicate that the turbulent intensity in the main flow direction is largest, while that in the vertical direction is the smallest, but only by a very small margin.

4.7. Spectral Analysis

In obtaining a statistically accurate spectral analysis of a turbulence record, an ensemble of time series would be necessary to estimate the spectra by averaging all together. Under controlled conditions such as in the laboratory, any long time series under a steady flow condition can be established relatively easily, and thus by dividing it

into shorter ones an ensemble of time series can be obtained. However, in case of tidal flow in the tidal channel as the present study, it is very difficult to obtain even a shorter time series under steady condition. The change in the mean velocity of two successive 910second long records, during three hours of peak flow, is normally within 3 cm/s. The sphere current probes are sometimes fouled by seaweed, for a short time series, even though it is not frequent, but for a long time series the chance of being interrupted by seaweed is much higher.

In this analysis, the spectra are presented as wave number spectra $S(k)$ on log-log plots in order to identify the wave number k dependence clearly. Note this plot is convenient to compare the present results with the existing ones. The wave number spectra $S(k)$ may be transformed from the frequency spectra $S(f)$ by adopting Taylor's hypothesis, in which the relation of $L = UT$ stands, where L is the length scale, U the local mean flow velocity, and T the period. The wave number k is equal to $1/L$, so this is different from the radian wave number $2\pi/L$. The wave number spectrum is thus given as $S(k) = S(f)U$.

Figure 5 and **Figure 6**, respectively, show typical turbulence spectrum for each of the components (S_u , S_v , S_w) at 1 m and 2 m from the sea bed. It is evident that in the low wave number range S_u is the largest, while S_w is the

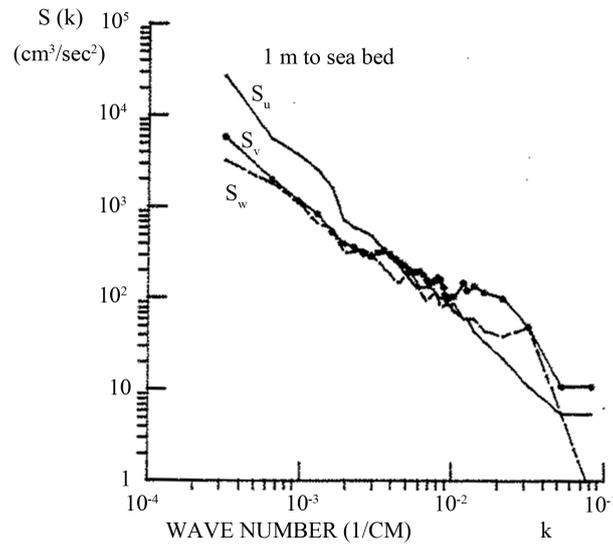


Figure 5. Three components of turbulent energy spectra (S_u , S_v , S_w) at 1 m above the big and steep dunes on the sea bed.

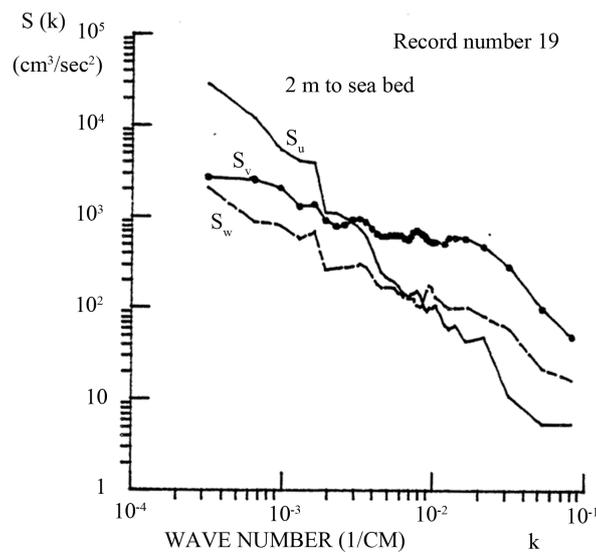


Figure 6. Three components of turbulent energy spectra (S_u , S_v , S_w) at 2 m above the big and steep dunes on the sea bed.

smallest. However, at higher wave numbers, the component spectra do not show this trend, but, despite the data scatter of each the measurement, it suggests anisotropy as depicted in these figures.

The magnitude of each of the spectral components decreases slightly with increasing the elevation from the sea bed, but increases with increasing the mean flow velocity U . These spectra never follow $-5/3$ power law due to Kolmogorov [28], but decrease intermittently and/or periodically with wave number, though in average they follow a power law between $-5/3$ and -2 , as predicted by Tsugé [17] and others. Thus, any sound power law of turbulent energy spectrum must express these basic characteristics properly, as to be proposed in the next section.

5. Proposition of Novel Power Law of Turbulent Energy Spectrum and Its Comparison with the Experiment

As early as in 1941, Kolmogorov [28] derived $-5/3$ power law of turbulent energy spectrum, which looks to have been supported by experiments at least “partially” in later years. It is, however, rather surprising that this law is based on purely dimensional analyses by imposing an intuitive hypothesis of “local homogeneity” for small-scale turbulence. It is more surprising that this famous law has little to do with the dynamic equations of fluid.

Proposed is a novel power law of turbulent energy spectrum in the inertial sub-range based on Tsugé turbulent equation [1] [6] governing turbulent flow: A correction term Z is added to the well-known one-dimensional Kolmogorov power law $P_{11} = Ck_1^{-5/3}$, where P_{11} is one-dimensional turbulent energy spectrum, C the constant depending on the initial conditions, and k_1 the x -direction wave number as follows

$$P_{11} = Ck_1^{-5/3+Z}$$

with

$$Z = \alpha_0 + \sum_{n=1}^{\infty} [\alpha_n \cos(nk_1) + \beta_n \sin(nk_1)] \quad \text{for } -1/3 \lesssim \alpha_0 \lesssim 5/3$$

where n is the mode number of fluid vibrations, and coefficients α_n and β_n are small compared with α_0 . This new power law has been discovered heuristically, based on the theoretical results by Tsugé [17], together with the turbulent measurements in the tidal current by Osonphasop [23] and. Osonphasop & Hinwood [22] [24].

The comparison of the present power law with the experiment suggests us the wide flexibility of this law, and indeed it is realized that this law can interpret any data in principle. It should be noted that the Kolmogorov power law of $P_{11} = Ck_1^{-5/3}$ has only a limited value, for this law has been derived by the dimensional analyses, so that it is not relevant to any dynamical equation governing the fluid motion. This is because it represents a merely straight line on the log-log scaled graphic paper.

Because turbulence is highly complicated motion of fluid particles in general, it must be time dependent and three dimensional. This is why the Kolmogorov power law has been inquired critically in the light of the sound theoretical and experimental facts.

Figure A2 illustrates Tsugé’s theoretical result on the power law (see Appendix). This figure illustrates how the power law depends on the initial conditions and oscillates with increasing the wave number k_1 . Whereas **Figure 6** delineates Osonphasop’s turbulent energy power spectrum for u , v , and w , respectively, obtained by turbulent measurements at 2 m above the sea bed in the tidal current [23]. This figure indicates that the power spectrum in x -, y -, and z -directions, respectively, neither follow the Kolmogorov power law nor draws straight line, but oscillates depending on each the wavenumber as being captured by the proposed universal power law of the turbulent energy spectrum. The infinite trigonometric series in Z -term will be used as approximate to any oscillation appearing in the energy spectrum.

The present proposed power law of turbulent energy spectrum is of significance for the practical point of view, for it is expressed by a simple mathematical form to be used by engineers. It is believed that discovery of the present law is an essential contribution to the progress of science if one recalls its wide applicability, covering many engineering professional fields such as fluid mechanics, hydraulics, oceanography, geophysical fluid dynamics, astrophysics, and many others.

6. Conclusions

New insights and new knowledge obtained through this study are summarized as follows:

1. It has been confirmed that Navier-Stokes equation has no potential to describe turbulence except for isotropic and homogeneous one.
2. Kolmogorov's $-5/3$ power law appears in the small region of k_1 in the inertial sub-range, but -2 power law prevails in larger region of k_1 . It is, therefore, inferred that neither $-5/3$ power law of the turbulent energy spectrum nor the -2 power law is universal in the inertial sub-range.
3. It is found that though numeral power laws in the inertial sub-range exist, but no law is adequate, so that novel power law of turbulent energy spectrum has been proposed as

$$P_{11} = Ck_1^{-5/3+Z}$$

with

$$Z = \alpha_0 + \sum_{n=1}^{\infty} [\alpha_n \cos(nk_1) + \beta_n \sin(nk_1)] \quad \text{for } -1/3 \lesssim \alpha_0 \lesssim 5/3$$

where n is the mode number of fluid vibrations, and coefficients α_n and β_n are small compared with α_0 .

It is clear that this parameter Z is not constant, but depends on the wave number k_1 , for the spectra oscillate with it. This fact has been confirmed both by Tsugé' theory and the present experiment in the tidal current.

4. The time records and spectra for the turbulent velocities have clearly demonstrated that the turbulence in the tidal current is highly three-dimensional.
5. It is reassured that the intermittence of momentum transport is a predominant and characteristic feature of geophysical boundary layer.
6. Tsugé equation or Kármán-Howarth equation retains the two-point or double correlation, but Navier-Stokes equation does not: It is, therefore, absolutely necessary for us to use the former equation to solve any problem on turbulence in order not to waste our time.

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Appendix: Brief Introduction How to Derive the Present Power Law of Turbulent Energy Spectrum

The complete account of this Appendix is discussed in Tsugé's paper [17], so that only a brief introduction on the paper will be made here.

The actual form of $q_1(s)$ written in axially isotropic form $q_1(s_1, \sigma)$ with $\sigma^2 = s_2^2 + s_3^2$ is,

$$\begin{aligned} q_1(s_1, \sigma) &= \partial^2 / \partial s_1^2 \left\{ \alpha \left[(s_1 - s_0)^2 + \sigma^2 \right]^{1/2} - \alpha \left[(s_1 + s_0)^2 + \sigma^2 \right]^{1/2} \right\} \\ &= \alpha \left\{ -2 \left[(s_1 - s_0)^2 + \sigma^2 \right]^{-3/2} + 3\sigma^2 \left[(s_1 - s_0)^2 + \sigma^2 \right]^{-5/2} \right. \\ &\quad \left. + 2 \left[(s_1 + s_0)^2 + \sigma^2 \right]^{-3/2} - 3\sigma^2 \left[(s_1 + s_0)^2 + \sigma^2 \right]^{-5/2} \right\} \end{aligned} \quad (1)$$

Note that there is no contribution from q_s to q_1 . For $|s_1| \gg s_0$, this expression approaches to

$$q_1 \sim 6\alpha s_0 s_1 \left(-2/s^5 + 5\sigma^2/s^7 \right), \quad (2)$$

which is no more than a quadruple field.

Figure A1 depicts streamlines of the fictitious flow generated by a pair of dipoles, which are given by the following potentials,

$$\begin{aligned} \phi^{(1)} &= \partial \phi^{(0)} / \partial s_1, \\ \phi^{(0)} &= \alpha \left\{ \left[(s_1 - s_0)^2 + \sigma^2 \right]^{-1/2} - \left[(s_1 + s_0)^2 + \sigma^2 \right]^{-1/2} \right\}. \end{aligned} \quad (3)$$

This axially-symmetric flow can be alternatively represented by using stream function ψ as

$$\begin{aligned} q_1 &= \partial \phi^{(1)} / \partial s_1 = \sigma^{-1} \partial(\sigma\psi) / \partial \sigma \\ q_{\perp} &= \partial \phi^{(1)} / \partial \sigma = -\sigma^{-1} \partial(\sigma\psi) / \partial s_1 \end{aligned} \quad (4)$$

from which we have the following relation,

$$\psi = -\partial \phi^{(0)} / \partial \sigma. \quad (5)$$

The flow pattern $\psi = \text{constant}$, shows quadruple-like structure at far field, $|s| \gg s_0$, toward which the longitudinal vortices are stretched stream-wise and getting thicker at the same time. On the returning path to the dipole core they are chopped off and trim the aspect ratio, getting into the dissipation region.

Let us obtain the power spectrum for inertial sub-range. The actual form of the pseudo-singularity enables us to calculate power spectrum for one-dimensional wave number k_1 in the inertial sub-range. This spectrum function $P_{11}(k_1)$ is written by definition,

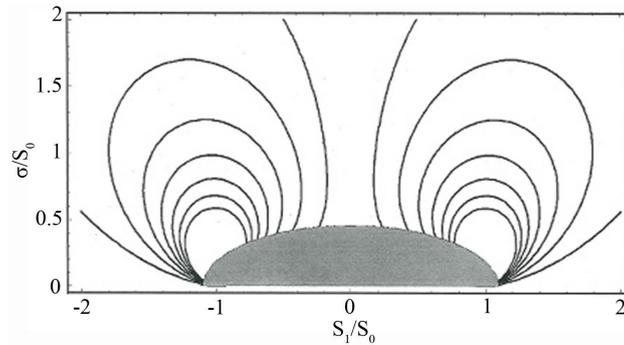


Figure A1. Pseudo-singularity formed by a pair of dipoles in the inertial sub-range of the eddy space. The shaded part represents the dissipation region. After Tsugé [17].

$$\langle u_1'^2 \rangle = \int_{-\infty}^{\infty} P_{11}(k_1) dk_1 \quad (6)$$

The Kolmogorov's consequence of the similitude analysis that P_{11} , having the dimension of $v^2 l \sim \varepsilon^{2/3} l^{5/3}$, where v is the velocity, l the eddy size, and ε the dissipation, leads in the language of wave number space as

$$P_{11} \sim k_1^{-5/3} \quad (7)$$

with $k_1 = l^{-1}$, while the present similitude analysis is that P_{11} , having the dimension of $v^2 \sim a l^2$, where a is the acceleration of flow, provides

$$P_{11} \sim k_1^{-2} \quad (8)$$

The present view into the existing formula for the power spectrum is rather straightforward: The concrete form of P_{11} can be derived by comparing expressions for the observable quantity of fluctuation correlation

$$\langle u_1'^2 \rangle = R \cdot P \cdot l^3 \int_{-\infty}^{\infty} dk_1 \int_{-\infty}^{\infty} dk_2 dk_3 f_1 f_1^* , \quad (9)$$

which results in

$$P_{11} = l^3 \int_{-\infty}^{\infty} dk_2 dk_3 f_1 f_1^* \quad (10)$$

where used is made of the definition of power spectrum function (6). The above integral (10) is transformed into the one in s -space as follows,

$$P_{11}(k_1) = (2\pi l)^{-3} \int_{-\infty}^{\infty} ds_1 \int_{-\infty}^{\infty} d\hat{s}_1 \exp[-ik(s_1 - \hat{s}_1)] \int_0^{\infty} \sigma d\sigma q_1(s_1, \sigma) q_1(\hat{s}_1, \sigma) \quad (11)$$

Since the integration spans over the whole s -space, we have yet to have solution q_1 inside the dissipation range $[O(l_k^3)]$, where viscous flow prevails essentially. However, the volume of the dissipation range is by far smaller than the inertial sub-range $[O(l_0^3)]$, the following potential flow solution might be sufficient for the present purpose:

$$\begin{aligned} q_1(s_1, \sigma) &= \partial^2 / \partial s_1^2 \left\{ \alpha / [(s_1 - s_0)^2 + \sigma^2]^{1/2} - \alpha / [(s_1 + s_0)^2 + \sigma^2]^{1/2} \right\} \\ &= \alpha \left\{ -2 / [(s_1 - s_0)^2 + \sigma^2]^{3/2} + 3\sigma^2 / [(s_1 - s_0)^2 + \sigma^2]^{5/2} \right. \\ &\quad \left. + 2 / [(s_1 + s_0)^2 + \sigma^2]^{3/2} - 3\sigma^2 / [(s_1 + s_0)^2 + \sigma^2]^{5/2} \right\} \end{aligned}$$

to be integrated over its own region. Then, the integral can be converted into a double one as follows,

$$P_{11}(k_1) = 2 / (2\pi l)^3 \cdot \int_0^{\infty} d(\sigma^2) Q(\sigma^2, k_1)^2 \quad (12)$$

with

$$Q(\sigma^2, k_1) = \int_{s_1^+}^{\infty} \sin k_1 s_1 \partial^2 \phi^{(0)} / \partial s_1^2 ds_1 \quad (13)$$

where the boundary contour $s_1 = s_1^+(\sigma^2)$ is given by

$$s_1^+(\sigma^2) = s_0(1 + \delta) \left\{ 1 - \sigma^2 / [s_0^2(2\delta + \delta^2)] \right\} \quad (14)$$

The one-dimensional power spectrum (10) is depicted in **Figure A2**. The parameter δ corresponds to the slenderness ratio of the worm as detected by the direct numerical simulation [10], which is estimated as the order of $O(10N)^{-1}$, where N is number in the order from 1 to 3. For a certain range of this parameter, the power spectrum shows $k_1^{-5/3}$ dependence, then with increase in the value of δ , it changes to k_1^{-2} for $\delta \gg 1$, where the pair of dipoles are regarded as a quadruple asymptotically.

It may be worth noting that the power spectrum P_{11} is not straight, but sinuous, depending on the wave number k_1 . This clearly suggests us that the power spectrum possesses intrinsic oscillatory character.

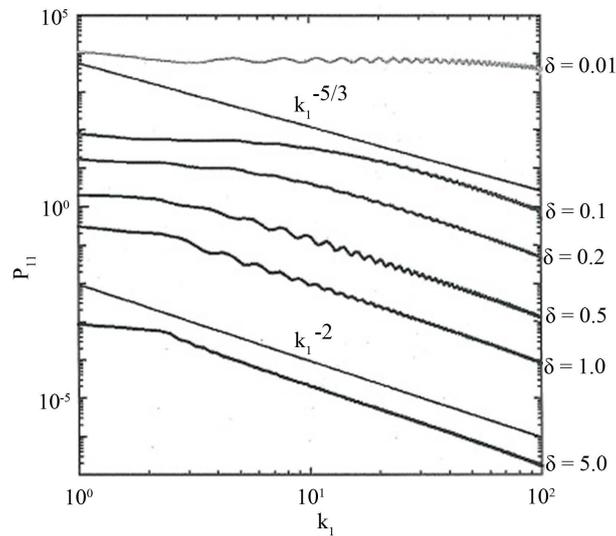


Figure A2. One-dimensional power spectrum P_{11} against one-dimensional wave number k_1 with the slenderness ratio of worm δ as the parameter. The calculated spectrum is averaged over five neighboring points to make the comparison with $-5/3$ and or -2 power law easier. After Tsugé [17].

Figure A2 indicates the following points.

1. For $\delta \leq 0.01$, P_{11} is almost independent of k_1 , but oscillates.
2. For $0.01 < \delta < 0.1$, the greater δ is, the more P_{11} rapidly decreases. Note $-5/3$ power law appears in this range of δ .
3. At $\delta \approx 0.1$, P_{11} decreases with increasing k_1 , following -2 power, if >40 .
4. For $\delta > 0.1$, there exists a region of k_1 , following -2 power law. It may be worth noting that with increasing δ , the range of k_1 following -2 power law becomes wider than otherwise.
5. For $\delta \gg 0.1$, P_{11} decreases with k_1 , following a negative power law having slightly greater than -2 of the slope in almost all region of k_1 .

Note that the one-dimensional power spectrum P_{11} is oscillatory, but not steady if k_1 is smaller than a certain value.