Review: the charnockite problem, a twenty first century perspective

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ABSTRACT
Beginning of the twentieth century was marked by coinage of a new rock name, Charnockite, first described as a hypersthene-bearing granite from Southern India. Since then charnockites have been described from most of the continents and mostly restricted to high-grade belts. Later half of the last century saw a lively debate over an igneous versus metamorphic origin. However, two factors acted as deterrents for the resolution of the debate. First, charnockites and associated rocks occur in a variety of different structural setting and display diverse field relations, attesting to possible different mode of origin. Second and possibly more important is the lack of consensus on the nomenclature of charnockites and associated rocks and this is commonly linked with the metamorphic versus magmatic perspective. Scanning the literature of this period makes one believe that both metamorphic and magmatic hypotheses are valid, but applicable to different field setting only. Before critically evaluating individual cases, it is imperative that a uniform approach in nomenclature should be agreed upon. It is proposed that name charnockite be adopted for any quartzofeldspathic rock with orthopyroxene, irrespective of its mode of occurrence, structural setting and mode of origin. The associated more mafic varieties, be better described as mafic granulite, rather than basic charnockite. For the patchy charnockites of east Gondwana (including parts of Peninsular India, Sri Lanka and Antarctica), metamorphic transformation from amphibolite facies gneiss, by two different mechanisms: CO₂ ingress from deep level, and drop in fluid pressure, has been proposed. However, all such patchy occurrence is not amenable to explanation by metamorphic transformation. In some instances, migmatisation of older charnockitic rocks is evident. Also progressive charnockitisation relating patchy charnockite to banded variety could be argued against on two counts: grain-size relation and time-relation. Larger bodies or bands have been explained as magmatic, but in many instances, from geochemical consideration alone. The compositional variation, commonly encountered in many high-grade belts, if not described in terms of field relation, may lead to wrong notions of magmatic differentiation of mantle-derived melts. Crustal melting of dry granulite facies source rocks has been proposed from geochemical and isotopic data of charnockitic intrusions. This model proposes high-temperature melting of previously dehydrated and dry granulite source rocks. However, tectonic perturbation subsequent to granulite facies metamorphism that might have been responsible for such high temperatures, is not well constrained in this model. Finally, with advent of high-pressure dehydration-melting experiments in the nineties, dehydration-melting of mafic to intermediate composition, syn-kinematic with granulite facies metamorphism has been proposed.

Keywords: Incipient Growth; Progressive Charnockitisation; Plutonic Charnockite; Partial Melting; Plutonic Metamorphism

1. INTRODUCTION
Holland [1] first described charnockite from south India, as hypersthene-bearing granite; and Howie [2] introduced the concept of plutonic metamorphism: charnockite magma emplaced at lower crustal depth, resulting in slow recrystallisation under great heat and uniform pressure. Recent researches, particularly, dehydration melting under granulite conditions in the deep crust, both experimental and empirical results is compatible with the concept of plutonic metamorphism.

Since then charnockite has been described from most
of the continents [3-9]. However, a variety of different mode of occurrence and structural setting, particularly the patchy occurrence first reported by Pichamuthu from Karnataka, south India [10] led to a lively debate over an igneous versus metamorphic origin. The first attempt for resolution of the charnockite problem was made by Ravich [11].

Another contentious issue relates to nomenclature, and to this day, this issue remains unresolved and no consensus among practicing earth scientists is a serious deterrent.

2. NOMENCLATURE

As more and more occurrences are reported, controversies on nomenclature cropped up and the IUGS classification, based on feldspar ratio, could not be uniformly implemented, even for the purported plutonic charnockites. On the one hand, many practicing earth scientists would use IUGS nomenclature, even for purported metamorphic charnockites [12-13]. It is noteworthy that not all the patchy occurrences are charnockite sensu stricto [14,16,17]. On the other hand, many of the reported plutons are enderbite or charno-enderbite, but described as charnockite by some workers [17-18].

Again, it has been noted from many granulate terranes that large-scale bodies commonly include charnockite and enderbite, along with intermediate varieties, but are not distinguishable in the field [15,19,20]. Lack of consensus on charnockite nomenclature continues and some recent publications use various terms like charnockitic gneiss of tonalite-trondjemite affinity, enderbite, enderbitic charnockite [21] and charnockites, charnockitic gneiss of tonalite-trondjemite affinity, enderbite, enderbitic charnockite [21] and charnockites, charnockitic rocks, chemically quartz-monzodiorite, quartz monzonite, granodiorite and granite [22]. It is important to note that orthopyroxene also occurs in high-temperature pelitic granulites, which should not be confused with charnockitic rocks [23].

SUMMARY:
The only plausible solution could be a general name for any quartzofeldspathic rock with orthopyroxene as charnockite (except of course high aluminous pelites), irrespective of the mode of occurrence, structural setting and mineralogical-chemical variations within each occurrence; the associated more mafic varieties may be described as mafic granulite, rather than basic charnockite, as first proposed by us [17]. The chemical classification then may follow Streickeisen’s scheme for common plutonic rocks and special names like enderbite etc may be omitted.

3. MODE OF ORIGIN

Naha et al. [24] noted that charnockitic rocks in south India occur in a variety of different structural settings, attesting to different styles and time-relations. Since 1960, when Pichamuthu first described patchy charnockites from Kabbaldurga in south India, the focus shifted to metamorphic transformation. Moreover, Newton and Hansen [25] questioned the possibility of slow cooling (and hence magmatic charnockite) and recrystallization of relatively dry granitic to intermediate magma under deep seated conditions, as proposed by Holland and Lambert [26]. Lack of experimental evidence on the primary crystallization of orthopyroxene from such H2O under saturated SiO2 rich liquid was their main argument and this created a strong bias in favor of metamorphic transformation. However, Kramers and Ridley [27] considered the evolution of the fluid phase during crystallization in the presence of orthopyroxene, and showed that fluid saturation curve is reached at the field of high CO2/H2O ratios and hence fluid inclusions are predicted. They further argued that “the patchy distribution of amphibolite & granulite facies TTG rocks in some high-grade terrains could be accounted for in this way”.

Melting experiments since the nineties, moreover, have highlighted the possibility of primary crystallization of orthopyroxene by dehydration melting reactions in the deep crust.

3.1. Metamorphic Transformation

From many localities in south India and Sri Lanka, “patchy” charnockites have been described as “arrested growth”, “in situ” charnockites or charnockitisation of amphibolite facies gneisses [28-39].

Two suggested mechanisms of this transformation: CO2-influx and drop in fluid pressure are reviewed in the following paragraphs.

Influx of CO2 rich fluid from deep mantle source along structural weak zones has been proposed by several workers [25,28,29,32,33,40], and Newton [15] mentioned three criteria for recognition of charnockitisation by CO2 influx, namely, 1) diffuseness of patchy alteration, unlike discrete veins; 2) occurrences closely associated with warping of foliation or dilation cracks; 3) open system alteration- often loss of mafic constituents and gain of Na and Si; Y and sometimes Rb are characteristically depleted. Some of these criteria are not ubiquitous, as argued by Bhattacharya and co-workers [14,16,17]. From Kerala and from Chilka area of the Eastern Ghats belt, these workers have argued that, 1) diffuse boundaries of the charnockite patches could have been produced by migmatisation of older charnockitic bodies by a granitic melt; 2) at Elavattum and Kottavattum quarries in Kerala, the apparent disposition along conjugate fractures [41], are actually disrupted segments of fold limbs (Figure 7 in Reference [14]). In Chilka Lake area the charnockite patches occur as elongate bodies parallel to sub horizontal F3 fold axis and along shear
planes with sub horizontal direction of maximum stretching; hence these weak zones are shallow structures and cannot act as channelways for fluid ingress from deeper levels. From the classical area of south India, Kabbaldurga, Bhattacharya and co-workers argued that the charnockite patches are usually not emplaced along the system of fractures, that are common in this region; and 3) four varieties of Peninsular gneisses: granite, trondhjemite, granodiorite and tonalite and three varieties of charnockite: granite, trondhjemite and tonalite are recognized in the quarry and charnockite patches occur within all varieties of peninsular gneisses; hence chemical similarity between close-pairs, cited as evidence of in-situ transformation by several workers, could be fortuitous [17]. The reported abundance of CO₂-rich fluid inclusions in patchy charnockites has been cited as evidence for the process of charnockitization by fluid-streaming [42] But Sen and Bhattacharya [16] argued that CO₂-enriched fluid inclusions may be due to preferential loss of H₂O by crystal plastic deformation and/or open system processes, as suggested by Hollister [43] and Buick and Holland [44]. For the patchy charnockites in the Eastern Ghats belt, CO₂-rich fluid streaming was also assumed by several workers [45-47]. But the possibility of large-scale influx of CO₂-rich fluids in the Eastern Ghats was ruled out by several workers [48,49]. Also deep mantle source of CO₂-rich fluid is not evident, while Bhowmik et al. [49] presented isotopic evidence of local sedimentary source (calc-granulites) in a granulite suite from the Eastern Ghats belt.

Raith and Srikantappa [41] proposed an alternative mechanism of this transformation. According to this hypothesis, arrested charnockitization is internally controlled; during near-isothermal uplift, the release of carbonic fluids from decrepitating inclusions in the host gneiss into developing fracture zones, resulting in a change in fluid regime and development of an initial fluid-pressure gradient, triggering the dehydration reaction. What is common, however, between the two hypotheses, is development of “arrested charnockite” in structural weak zones. For the Kabbaldurga occurrence, Bhattacharya and Sen [17] pointed out that “charnockite veins at Kabbal are usually not emplaced along the system of fractures that are common in this region.”

Time relations between charnockites and enclosing gneisses, as also between patchy occurrence and massive bodies, are important constraints, for validating or otherwise of the hypothesis of in-situ transformation. Naha et al. [24] pointed out that charnockites of Dharwar craton have formed in at least two distinct phases separated in time and possibly by different mode of origin. And Bhattacharya and Sen [17] pointed out that “patchy charnockites seen in Kerala and in the Eastern Ghats are mostly non-pegmatitic”; “the coarser-grained patches could very well be modified versions of the smaller patches”; and “... are basically earlier than the enclosing gneisses”. It is imperative, therefore, to consider individual cases of “patchy charnockite”, in terms of field relations and if possible, in terms of isotopic age relations. For the Chilka Lake case in the Eastern Ghats belt, Bhattacharya et al. [50] reported older zircons in the patchy charnockite to those of the host leptynite/granite gneiss.

Another point of contention is the proposed link between patchy charnockite and massive charnockite, particularly in South India. Srikantappa et al [34] proposed progressive charnockitisation, from some locales in the Kerala Khondalite belt. But Sen and Bhattacharya [15] argued that size relation between smaller patches and adjacent larger bands (supposedly final product) does not support this hypothesis. Sen and Bhattacharya [15] further argued on the evidence of field relation between them, that larger bands are actually older. On the other hand, the proposed genetic link between the incipient/arrested charnockite of the transition zone in South India to regional scale granulites (massive charnockite), is strongly influenced by the CO₂-influx hypothesis [19,29,31-33,51]. According to this model ascent of the carbonic fluid front to higher crustal levels, results in pervasive fluid flow and wholesale granulitization of the deeper crustal domains. However, Raith and Srikantappa [41] argued on the evidence of field relations, petrological, geochemical and isotopic data, that development of arrested charnockites is a late-stage phenomenon; and regional-scale granulites could have been generated by dehydration melting processes.

**SUMMARY:**

Proposed hypothesis of charnockitisation, either by CO₂ influx or drop in fluid pressure, could indeed be applicable for individual cases; but each would require additional data pertaining to structural setting and field structural data attesting to time relation. Additionally, isotopic data would resolve the issue in favor or against the hypothesis of progressive charnockitisation. It is emphasized here that patchy occurrence itself should not be taken as prototypes of incipient charnockite.

### 3.2. Magmatic Origin

Since Howie [2] proposed the hypothesis of plutonic metamorphism, large-scale charnockitic rocks have been described from many granulite terranes [20-22,52-57]. Subba Rao and Divakara Rao [53] described charnockitic rocks of intrusive origin from Eastern Ghats Belt, and identified two groups, namely basic granulite and charnockite. From geochemical angle, these authors proposed that protoliths of these charnockitic rocks are the fractionated products of a melt, which was derived from metasomatised mantle, and that these were affected
by a depletion event probably coeval with granulite facies metamorphism. Although, two groups were said to be identified “based upon field relations and chemistry”, the actual field relation between basic granulite and charnockite is not described in this publication. Moreover, as noted earlier by several workers, local structural setting and sample locations are important criteria, and without these information, the applicability of the mantle melting model proposed by these authors can be questioned. In this context, it is important to note that from detailed field mapping and structural analysis in the Chilka Lake area of the Eastern Ghats belt, India, Bhattacharya et al [57] argued that “certainly an igneous protolith which has suffered granulate facies metamorphism (as evidenced by inter-layered basic granulites) is a distinct possibility”. It is unfortunate that some workers concluded that in the Eastern Ghats belt, age relations may be deduced from field relations, but neither do they present any data, nor refer to published information; hence their conclusion that “intruding magmas are either mantle-derived (basic granulites, enderbites and charnockites with crustal contribution)…..” remains questionable [58]. Bhattacharya and co-workers described two types of field relations between charnockitic rocks and metapelitic rocks. First type is the interbanding of the two lithologies; the time relation is uncertain, though both may have undergone granulate facies metamorphism together [57,59]. The other type of field relation between the two lithologies is all the more complex; large-scale bodies of charnockitic rocks usually occur as separate exposures, and no contact between the two could be observed; no pelitic enclaves were observed in charnockites. Only on the basis of the sequence of deformation structure, a tentative correlation has been proposed: mafic granulite, occurring as folded enclaves in charnockite, could be correlated to intrafolial folds in pelitic granulites [57,60]. Dobmeier and Raith [13] also observed that “since the enderbitic and metasedimentary rocks have identical structural histories, the emplacement (of enderbitic/tonalitic magma) happened prior to the discernible deformation…” in the Chilka Lake area.

Magmatic origin of charnockite is also proposed by several workers in the nineties. Kilpatrick and Ellis [7] described Charnockite Magma Type, or C-type, from different areas, with distinctive geochemical signatures. This C-type magma was considered to be derived by melting of a dry granulate source. It should be noted that this C-type magma is not strictly charnockite sensu stricto, but varies between charno-enderbite and charnockite (see $K_2O/Na_2O$ ratios and $SiO_2$ values in Table 1 of Reference [7]). Also the melting here is considered to have been post-granulate facies metamorphism and a crustal-melting event. Melting of dry granulate-facies source rocks, for Antarctic charnockites, was also proposed by some workers from geochemical and iso-

topic data [20,55,61]. On the other hand, Sheraton et al. [54] argued that more mafic varieties may be largely mantle-derived. It is important to note that these reports on Antarctic charnockites show a range of composition from quartz monzodiorite through granodiorite to adamellite. Hence, discrimination between charnockite and enderbite magma, in massif-type or intrusive charnockite, was considered inappropriate by these authors. This model proposes high-temperature melting of previously dehydrated and dry granulate source rocks. But tectonic perturbation subsequent to granulate facies metamorphism that might have been responsible for such high-temperatures is not well constrained in this model.

A partial melting interpretation for vein type charnockite was advocated by Hansen and Stuk [62], and these authors reported orthopyroxene-bearing leucosomes, of tonalitic to granodioritic composition, within mafic bodies of granulate facies rocks from California.

Finally, melting experiments, particularly dehydration-melting experiments of mafic to intermediate rocks in the nineties have added a new dimension to the problem of charnockite genesis [63-66]. These experiments demonstrate a) significant melting at 8 to 10 kbar and temperatures in excess of $850^\circ C$; these values are commonly recorded from many granulate terrains; b) the residual assemblage of two-pyroxene-plagioclase-quartz ± garnet, clearly resemble mafic granulate, that are frequently found associated with massif-type charnockite; c) melt compositions in hornblende-dehydration melting range from tonalite-granodiorite-trondjhemite, while hornblende-biotite combined melting produced granitic melts.

From the classic area, Kabbaldurga, in South India, Bhattacharya and Sen [17] presented a new interpretation of vein type charnockite. These authors proposed hornblende and biotite dehydration melting in two types of mafic granulites observed in the area, producing two types of charnockitic vein, of tonalitic and granitic compositions respectively. Besides the field features, such as orthopyroxene-bearing leucosomes within mafic granulate enclaves in the peninsular gneiss; these authors presented comparative mineral compositions in the charnockite veins and mafic granulate enclaves and bulk compositions of the charnockite veins, and these are compatible with the results of experimental melting, referred to above.

For the massif-type charnockite in the Eastern Ghats belt, India, Kar et al. [56] proposed a hornblende-dehydration melting in mafic rocks, now occurring as cognate xenoliths, under granulate facies conditions. Additionally these authors reported two types of mafic granulites, namely prograde hornblende-bearing mafic granulate, interpreted as restitic granulate and two-orthopyroxene mafic granulate, interpreted as peritectic segregations.
And unlike Subba Rao and Divakara Rao’s [53] mantle-melting model for the Eastern Ghats charnockite, these authors described a crustal melting phenomenon, coeval with granulite facies metamorphism. From pressure-temperature estimates and P-T path constraints, these authors further argued that melting could have occurred in thickened continental crust undergoing decompression. Bhattacharya et al. [67] established the link between partial melting and granulite facies metamorphism with isotopic data. Kar et al. [56] further pointed out that trace element partitioning in dehydration melting is likely to be complex, because incongruent melting reactions result in two sets of solid mineral phases, residual and peritectic [68]. Hence quantitative modeling is inappropriate when the process involves reactions producing a variety of solid peritectic phases. Trace element partitioning then could be considered as a two stage process; to some extent correlated with different degrees of partial melting. At low degree of melting the main process is melt-restitie separation, whereas at higher degrees of melting peritectic-melt separation becomes more important [69-71].

SUMMARY:
Although magmatic origin of charnockites, particularly for the large scale bodies, are evident in many cases, the question relating to either mantle-melting or crustal melting and in case of crustal melting, the actual melting process and conditions remain debatable in many cases. Dehydration-melting in mafic to intermediate rocks and conditions remain debatable in many cases. Dehydration-melting in mafic to intermediate rocks under granulite facies conditions could be the most potential hypothesis for the massif-type charnockite, provided prograde hornblende/biotite bearing mafic granulite enclaves are observed. Thus the concept of plutonic metamorphism may return with new vigor.

REFERENCES


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