The Influence of the Planets, Sun and Moon on the Evolution of the Earth’s Axis

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Abstract

To study climate evolution, we consider the rotational motion of the Earth. The law of angular momentum change is analyzed, based on which the differential equations of rotational motion are derived. Problems with initial conditions are discussed and the algorithm of the numerical solution is presented. The equations are numerically integrated for the action of each planet, the Sun and the Moon on the Earth separately over 10 kyr. Results are analyzed and compared to solutions of other authors and to observation data.

Keywords: Angular Momentum, Differential Equations, Numerical Solution

1. Introduction

In the history of the Earth the periodical changes of sedimentary layers on continents and in oceans, their chemical composition and magnetic properties are observed. The oscillations of oceans levels and repeating traces of activity of glaciers are traced also. The climatic changes are determined by different factors, including: the solar heat quantity acting to the Earth, which depends on angle of fall of solar beams on the Earth’s surface, the distance from the Sun and duration of illumination. In the 1920s the Yugoslavian scientist M. Milancovitch [1] created the astronomical theory of ice ages in which these three factors are expressed by a tilt between the Earth’s orbital and equatorial instant planes (obliquity), an eccentricity of an orbit and position of a perihelion. Subsequently calculations of Milankovich was repeated by other researchers: Brouwer and Van Woerkom [2], Sh. Sharaf and N. Budnikova [3], A. Berger and V. Loutre [4], J. Laskar et al. [5], and also others.

As a result of interaction of Solar system bodies there are changes of orbits of the planets. When rotating, the Earth becomes stretched in its equatorial plane. Therefore each of external bodies creates a moment of force, which causes precession and nutation of the Earth’s spin axis. The joint motion of the Earth’s spin axis (or equatorial plane) and orbital plane produces a tilt between the two moving planes, which controls the terrestrial insolation.

In the above mentioned works the problem of orbital movement was solved approximately by analytical methods, within the framework of the so-called theory of secular perturbations. In the second problem of the rotational motion, the second-order differential equations for Earth rotation have been commonly reduced to first-order Poisson equations and were also solved approximately and analytically.

Each of the mentioned above followings groups of the researchers used more perfect analytical theories of secular perturbations. However, the theories of rotary motion were based on the Poisson equations with taking into account only the action of the Moon and the Sun. With the today’s computing facilities and advanced processing tools, the simplifications applied in the analytical method are no longer indispensable. We have integrated the equations of the Sun, the planets and the Moon orbital motion for 100 Myr with the help of a numerical method [6]. The results of a numerical integration of the rotational motion equations are presented in this work.

According to P. Laplace [7], Isaac Newton apparently was the first who solve the problem of Earth’s rotation. Since then the equations of rotary movement have been repeatedly derived by different authors. Thus they were based on different theorems of the mechanics, used different systems of coordinates, different designations and different methods of solution. All authors began to simplify equations during the deriving them with the purpose of using analytical methods for solution. In the great majority these equations were reduced to Poisson equations. Thus, there are no standard non-simplified
differential equations of the Earth rotary movement in the literature, which could be integrated without any change.

Besides, numerical integration causes a number of problems. For example it is impossible to solve the task of initial conditions without understanding of all the features and details of the differential equations deriving. Frequently there are contradictions between treatment of complex Earth’s spin in astronomy and requirements, which follow from the mechanics laws. Therefore we had to analyze different ways of the deriving of the equations and the simplest way submitted in the present work was chosen. Taking into account the large derivation only the basic mathematical transformations of the differential equations are given here and the scheme of derivation of others is described.

The appearance of precise observation systems about 50 years ago made it possible to investigate the short-term dynamics of the Earth’s spin. This dynamics was explained by theories of precession and nutation. Researchers were forced to enter additional effects besides the basic gravitational influence because there were the divergence between these theories and observation [8]. For instance, a geopotential correction was applied to estimate the gravitational interaction of the Earth’s mass on long-term intervals. Therefore they are taken into account at long-term Earth evolution.

To understand the distinction between the theory and observation there were models of a nonrigid Earth or of a structured Earth in which every structure, for example, the core, followed its own motion [11]. Ice redistribution in polar areas leads to change of the moments of inertia on long-term intervals. Therefore they are taken into account at long-term Earth evolution.

Practically all additional effects are not determined as precisely, as gravitational forces. Influence number of them has hypothetical character. Some of them are offered by experts from different areas of physics and experts in the theoretical and celestial mechanics should apply them on belief without a strict analysis. Relativistic additives are concerned to those. These forces depend not only on distance between interacting bodies, but on their velocities as well [12,13]. Systems with such connections are known as non-holonomic systems in theoretical mechanics. Contradictions between equations of motion and laws of conservation are familiar to experts in the General Theory of Relativity. Therefore, energy methods of mechanics demand correction for nonholonomic systems.

All these small additional actions are based on divergences between computations of the basic gravitational action on Earth’s spin and observation. However, this calculation was executed approximately as it was mentioned. There are a number of simplifications in the derivation of the Earth’s rotational motion differential equations, which can give difference in computations and observation data as it will be shown below. Therefore, one represents the great interest in obtaining the most accurate solutions given no doubts that the discrepancy between calculations and observation are really explained by other factors. This position is more actual in the long-term rotational motion research. Additional small short-term observational actions may yield unreal results in millions years intervals. Thereby, only Newtonian gravitational interaction on rotational motion of axisymmetrical Earth is described further.

2. The Law of Angular Momentum Change and Its Consequences

According to the theorem of momentum relative to moving center of mass [14], the rotational motion of a mechanic system is described as the law of angular momentum change:

\[
\frac{dK_0}{dt} = \sum m_1 (F_i)
\]

where \(K_0\) is the angular momentum of a mechanic system relative to the center \(O\) in the non-rotating coordinates \(x_O y_O z_O\) (Figure 1), and \(m_0 (F_i)\) is the moment of the applied force \(F_i\) relative to this center in the same coordinates.

We provided a number of experiments (Figure 1) to understand better the mechanism of gravitation force action on a rotating body. Let a whirligig (Figure 1(a)) rotates on a table. Having its own axis counterclockwise inclination from a vertical at the angle \(\theta\), it is treated by gravitation force \(P\) creating the moment of force \(m_O\), which is counterclockwise also. In case of the suspended wheel (Figure 1(b)) the torque \(m_O\) will be clockwise. Experiments were executed at different \(\theta\) angles and angular speeds of these bodies. Direction of precession of the whirligig and a suspended wheel axes were opposite in all experiments.

Now we consider the law of angular momentum (1) on an example of rotational movement of three bodies submitted on Figure 1 in more details. Let a whirligig (Figure 1(a)) be at angle \(\theta\) to the axis \(z_1\) and spin about its axis \(z\) at the angular velocity \(\phi\), and the gravitational force \(P\) be applied to its gravity center. The torque
Figure 1. Precession of rotational bodies: (a) Whirligig on bearing surface \(x_Oy_1\); (b) Wheel suspended at the point \(O\); (c) Free Earth. 1 and 2—Earth’s equatorial and orbital planes; 3—orbital plane of the body \(B\).

(moment of force) \(m_\phi = P_\theta \cdot \sin\theta\) makes the whirligig axis rotate about the point \(O\) at the angular velocities \(\dot{\theta}\) and \(\dot{\psi}\). The vector of the whirligig absolute angular velocity is

\[
\omega = \dot{\phi} + \dot{\theta} + \dot{\psi}
\]  

(2)

Let a axisymmetrical whirligig with the moment of inertia \(J_\phi\) relative to \(z\) has the velocity \(\dot{\phi}\) which is much faster than the velocities of its axis \(\dot{\theta}\) and \(\dot{\psi}\). If we neglect the deviation of \(K_\phi\) from \(z\) (Figure 1(a)), the whirligig approximate angular momentum \(K_\phi \approx K_\phi \approx J_\phi \cdot \dot{\phi}\) is directed along \(z\), and law (1) becomes written approximately as

\[
K_\phi \approx m_\phi
\]  

(3)

This approach is employed in the elementary gyroscope theory. The torque \(m_\phi\) being normal to \(K_\phi\) (Figure 1(a)), the angular momentum remains invariable in magnitude \(K_\phi = J_\phi \cdot \dot{\phi}\) but changes its direction \(\Delta K\) in the plane \(x_Oy_1\) according to \(m_\phi\). Then, for the time \(\Delta t\), the angular momentum receives the increment \(\Delta K = K_\phi \cdot m_\phi \cdot \Delta t\) which causes its rotation with the angle \(\Delta\psi\) about \(z_1\). If the increment \(\Delta K\) is expressed via the angle increment \(\Delta\psi\), \(\Delta K = K_\phi \cdot \Delta\psi = K_\phi \cdot \sin\theta \cdot \Delta\psi\) the rotation velocity becomes

\[
\dot{\psi} = \lim_{\Delta t \to 0} \Delta\psi / \Delta t = m_\phi \cdot K_\phi \cdot \sin\theta
\]  

(4)

Inasmuch as the whirligig spin axis coincides with the angular momentum \(K_\phi\) in this formulation, it will rotate in the same way as \(K_\phi\), i.e., will precess counterclockwise at the velocity \(\dot{\psi}\). After the moment \(m_\phi\) is substituted into (4), the precession velocity becomes

\[
\dot{\psi} = P_\theta / (J_\phi \cdot \sin\theta)
\]  

(5)

In the applied approximation, the precession velocity does not depend on the tilt of the whirligig spin axis \(\dot{\theta}\) and moreover, there are no nutation oscillations (changes of angle \(\dot{\theta}\)).

In case of a wheel suspended at the point \(O\) (Figure 1(b)), \(m_\phi\) is directed clockwise, and the wheel axis precesses also clockwise at a velocity governed by (5).

Figure 1(c) shows the action of body \(B\) on the rotating Earth. If the Earth is centrosymmetrical, the action of \(B\) on the Earth’s parts nearest and farthest to \(B\) corresponds to the forces \(F_1\) and \(F_2\), whose resultant passes through the center \(O\). For an oblate Earth, the purchases of these forces become farther from the center \(O\). Then, \(F_1\) increases while \(F_2\) decreases producing the torque \(m_\phi\) directed clockwise as in Figure 1(b). Therefore, the Earth’s spin axis will precess clockwise at a velocity given by (4).

Above we considered the rotation at an angular velocity \(\dot{\phi} > 0\), i.e. counterclockwise. If the rotation changes its direction from counterclockwise to clockwise, the precession changes correspondingly, i.e. the velocity changes sign. It was proved by experiments.

It is possible to establish the precession cycles of the Earth’s spin axis by analyzing the action body \(B\) on Earth (Figure 1(c)). The body \(B\) produces the maximum torques \(m_{\phi max}1\) of the same direction at points \(B_1\) and \(B_1\), which are zero when the body is in the Earth’s equatorial plane \(B_1\) and \(B_1\), i.e., the torque changes twice from 0 to \(m_{\phi max}1\) during a single orbiting cycle of \(B\). Therefore, the precession and nutation cycles of the Earth’s spin axis equal the half orbiting periods of the Sun, the Moon, and the planets relative to the moving
Earth’s axis Oz, or to the moving equatorial plane. The Earth and the planets may approach in their orbital motion. If the approachment occurs at the points $B_1$ and $B_3$, the maximum torque increases to $m_{1\text{max}2z}$. Thus, the Earth’s axis oscillates with periods corresponding to cycles at which the Earth approaches with planets are especially the closest to it at the points $B_1$ and $B_3$. Note, that with the purpose of simplification Figure 1(c) the three planes are shown with intersection at line $B_1B_2$: 1 and 2—Earth’s equatorial and orbital planes, 3—orbital plane of the body $B$. Therefore these oscillation periods are modulated because the orbits of the Earth and the planets do not belong to the same plane and are rather elliptical than circular.

The torque produced by body $B$, depends also on parameters of its orbit, in particular from a tilt between the equatorial plane 1 (Figure 1(c)) and the orbital plane 3. Therefore the Earth’s spin axis will oscillate with the period of orbit precession of the body $B$ relative to the Earth’s moving equatorial plane. For example, the precession cycle of the Moon’s orbit relative to the Earth’s orbital plane is equal 18.6 yrs. The Earth’s spin axis will oscillate with this cycle at the Moon action. There will not be an essential difference between this cycle and the cycle of the Moon’s orbital precession the Earth’s equatorial plane is practically motionless.

3. Differential Equations of Rotational Motion

3.1. Earth’s Angular Momentum

There arises a problem in the derivation of the rotation equations: as the body rotates the moment of inertia or the angular velocity components are changed depending on the coordinate system. Therefore, we first use the coordinates in which the moments of inertia remain invariable and then we proceed to the coordinates where the angular velocities are independent of the body rotation.

Motion of the Solar system bodies is considered in non-rotating barycentric ecliptic coordinates $x_{10}y_{10}z_{10}$ (Figure 2) connected to the Earth’s orbital plane fixed at $T_0$ epoch. Axis $x_{10}$ is pointed toward the vernal equinox. Let the non-rotating system of geocentric ecliptic coordinates $x_{1}y_{1}z_{1}$ with the center of mass at $O$ moves translationally relative to the system $x_{10}y_{10}z_{10}$. Axis $z$ of the rotating geocentric ecliptic coordinates $xyz$ is directed along the Earth’s own angular velocity $\phi$ and axis $x$ at the initial moment $t = 0$ lies in the Greenwich Meridian plane. The Earth’s absolute angular velocity $\omega$ in system $x_{1}y_{1}z_{1}$ with the projections $\omega_x$, $\omega_y$, $\omega_z$ to rotating axes of system $xyz$ will be $\omega = i\omega_x + j\omega_y + k\omega_z$.

In the law (1), angular momentum $K_o$ is produced by all the masses of rotating Earth in system $x_{1}y_{1}z_{1}$, and $\sum m_{1}(F_{i})$ is moment of forces produced by all the Solar system bodies. First define derivation of the angular momentum and then the sum of torques.

As in system $x_{1}y_{1}z_{1}$ the Earth rotates with angular velocity $\omega$, thus any of its element $dM$ with a radius vector $r = ix + jy + kz$ (Figure 2) moves with velocity $v = \omega \times r$ and has angular momentum $m_{1}(dMv) = r \times v dM$ relative to the center $O$.

After integration taking into account the entire Earth’s mass $M$, $K_o$ is

$$K_o = \int m_{1}(dMv) = \int r \times v dM$$

(6)

Having differentiated $K_o$ on time and then having substituted vectors $r$, $\omega$ and $v$, after transformation we obtain derivatives of the projections of the angular momentum to axes of rotating system $xyz$

$$K_{o1} = J_x \epsilon_x - J_y \epsilon_y - J_z \epsilon_z - J_x (\epsilon_y + \omega_z \epsilon_z) - (J_y - J_z) \omega_x \epsilon_z - J_z (\epsilon_y - \omega_z \epsilon_z)$$

(7)

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\[ \dot{K}_{Ox} = J_x e_x - J_{yz} (e_x - \omega_x \omega_z) - J_{yz} (e_x + \omega_x \omega_z) \]  
\[ -(J_x - J_z) \omega_x \omega_z - J_{ys} (\omega_z^2 - \omega_x^2) \]  
\[ \dot{K}_{Oy} = J_y e_y - J_{xz} (e_y - \omega_y \omega_z) - J_{xz} (e_y + \omega_y \omega_z) \]  
\[ -(J_y - J_z) \omega_y \omega_z - J_{yx} (\omega_z^2 - \omega_y^2) \]  
\[ \dot{K}_{Oz} = J_z e_z - (J_x - J_y) \omega_x \omega_y, \]  
\[ \dot{K}_{Oy} = J_y e_y - (J_z - J_x) \omega_y \omega_z, \]  
\[ \dot{K}_{Oz} = J_z e_z. \]  

where \( e_x, e_y, e_z \) are projections of the Earth angular acceleration in system \( x'y'z' \) to axes of rotating system \( xyz \);  
\[ J_x = \int (y^2 + z^2) dM, \quad J_y = \int (z^2 + x^2) dM, \]  
\[ J_z = \int (x^2 + y^2) dM - \text{the Earth’s axial moments of inertia}, \]  
and \( J_{yx}, J_{yz}, J_{xz} \) are its centrifugal moments of inertia.

As the system \( xyz \) is connected with the rotating Earth, the moment of inertia does not depend on the Earth’s rotation and remain constants. Today’s knowledge of the Earth’s density distribution is insufficient for definition of the moments of inertia. As it will be shown below, it is possible to define only the relation between two moments of inertia \( J_x \) and \( J_y \), from observable angular velocity \( \psi \) (precession), but not their absolute values. Last years the three-axial model of Earth is described where the third moment of inertia is estimated by potential of the gravitational distribution on Earth’s surface. However, this method can not give precise estimation of the moments of inertia. Note, that in case of three-axial Earth it is required to add terms with centrifugal moments of inertia because of nonsymmetrical distribution of potential as shown from (7)-(9). Traditionally, terms with these moments are unaccounted and terms containing third moment \( J_z \) are in equations though in last forms \( J_z \) is equal \( J_z \). Nevertheless, avoiding of complications with unaccounted terms it is considered axisymmetrical Earth \( J_x = J_y, \) and \( J_{yz} = J_{xz} = J_{xz} = 0 \). Then derivatives of the angular momentum (7)-(9) are:

For axes with zero centrifugal moments of inertia the moments \( J_x, J_y, J_z \) are named principal moments of inertia.

### 3.2. The Law of Angular Momentum in Euler Angles

Position of rotating system \( xyz \) (Figure 2) relative to non-rotating system \( x'y'z' \) is determined by Euler angles: \( \varphi, \psi, \theta \), here precession angle \( \psi \) defines position of nodes line \( OK \), where moving equatorial plane intersects fixed orbital plane (ecliptic), nutation angle \( \theta \) defines a tilt between equatorial plane and ecliptic and \( \varphi \) is angle of the Earth’s own rotation around axis \( z \). Figure 2 shows directions of this angles accepted in theoretical mechanics. In astronomy angles, which are directed according to observably movements, are accepted: \( \psi = 2\pi - \psi, \varphi = \varphi \) and \( \theta = -\theta \). Directions of angular velocities \( \psi, \theta, \phi \) are shown on Figure 2, i.e. turn of angles \( \psi, \theta \) and \( \varphi \) is visible counterclockwise from the side of these vectors. It is seen from Figure 2 that Euler’s velocities stay constant at the Earth’s turn. Therefore let all the variables and equations will be expressed in Euler variables.

As it is possible to present the vector of absolute angular velocity as (2), thus in projections to axes \( x, y, z \) it is

\[ \omega_x = \phi_x + \psi_x + \dot{\theta}_x, \]  
\[ \omega_y = \psi_x + \dot{\psi}_x + \dot{\theta}_x, \]  
\[ \omega_z = \dot{\phi}_x + \dot{\psi}_x + \dot{\theta}_x. \]  

where \( \phi_x, \psi_x, \dot{\theta}_x \) etc. —projections of the Euler angular velocities to axes \( xyz \). Having calculated expressions (11) with the help of Figure 2 it is received Euler well-known equations:

\[ \omega_x = \psi \sin \theta \sin \phi + \theta \cdot \cos \phi, \]  
\[ \omega_y = \psi \sin \theta \cos \phi - \dot{\theta} \cdot \sin \phi, \]  
\[ \omega_z = \phi + \psi \cdot \cos \theta. \]  

After differentiation (12) the angular accelerations in Euler variables become:

\[ \epsilon_x = \psi \sin \theta \cos \phi + \dot{\psi} \cdot \cos \theta \cos \phi \]  
\[ + \theta \cdot \sin \phi \cos \theta \cos \phi, \]  
\[ \epsilon_y = \psi \sin \theta \sin \phi + \dot{\psi} \cdot \cos \theta \sin \phi \]  
\[ - \theta \cdot \cos \phi \sin \theta \cos \phi, \]  
\[ \epsilon_z = \dot{\phi} + \psi \cdot \cos \theta - \psi \cdot \sin \theta. \]  

Having substituted angular velocities (12) and angular accelerations (13)-(15) in (10), the projections of derivatives of the angular momentum depending on Euler angles are:

\[ \dot{K}_{Ox} = J_x \psi \sin \theta \sin \phi + J_x \dot{\theta} \cos \phi \]  
\[ + (2J_y - J_z) \psi \theta \cos \phi \sin \phi \]  
\[ + J_y \dot{\psi} \sin \theta \cos \phi - J_y \dot{\theta} \sin \phi \]  
\[ + (J_z - J_y) \psi^2 \sin \theta \cos \phi \cos \phi; \]
\[ \hat{K}_{ij} = J_i \psi \sin \theta \cos \phi - J_j \psi \sin \phi \\
+ (2J_i - J_j) \psi \cos \theta \cos \phi \\
- J_i \psi \phi \sin \phi \sin \theta - J_i \psi \phi \cos \phi \\
-(J_j - J_i) \psi^2 \sin \theta \sin \phi \sin \phi; \]

(17)

\[ \hat{K}_{ij} = J_i \phi \sin \theta \cos \theta + J_j \phi \sin \theta \sin \phi, \]

(18)

Let us transform the law of the angular moments (1) to Euler angles. The moments of forces can be defined directly through forces for the actions of each body on an Earth mass element (see Figure 2). However, hereinafter at their series representation they are reduced to other terms. Here we define the moments through the force function \( U \) by a traditional method to safe continuity with previous works. As it is known from the theorem of theoretical mechanics, the moments of forces in a direction of axes are equal derivatives of force function on an angle of turn relatively to this axis. Therefore in system \( OK_2z \) with the axes located on vectors of angular velocities \( \psi \), \( \theta \) and \( \phi \), the moments of forces are (Figure 2):

\[ m_{\phi \psi} = m_{\psi \phi} = \partial U / \partial \psi, \quad m_{\phi \theta} = m_{\theta \phi} = \partial U / \partial \theta, \]

\[ m_{\psi \phi} = m_{\phi \psi} = \partial U / \partial \phi. \]

Taking into account these moments we project the right and left part of (1) on the axes of system \( OK_2z \):

\[ \hat{K}_{ij} = \partial U / \partial \psi, \quad \hat{K}_{ij} = \partial U / \partial \theta, \quad \hat{K}_{ij} = \partial U / \partial \phi \]

(19)

Let us express derivatives of the angular momentum in Equation (19) through derivatives of force function depending on the Earth’s spin in Euler’s variables:

\[ \psi = \left( \frac{\partial U}{\partial \psi} - (2J_i - J_j) \psi \theta \sin \phi \cos \phi + J_j \phi \theta \phi \sin \phi \phi \right) \]

\[ \frac{1}{J_j} \left( \frac{\partial U}{\partial \phi} + (J_j - J_i) \psi^2 \sin \phi \cos \phi - J_i \phi \psi \sin \phi \phi \right) \frac{1}{J_i} \]

(22)

where the derivative of the angular momentum \( \hat{K}_{ij} \) is defined by the third formula in the equations (19) for the law of angular momentum.

### 3.3. The Force Function in Cartesian Coordinates

Now we define force function for the action of body \( B_i \) with mass \( M_i \) (Figure 2) on the Earth by the traditional way presented, for example, in [15]. Its coordinates in rotating system \( xyz \) are \( x_i, y_i, z_i \). The force function of the action of this body on the Earth’s element of mass \( dM \) with coordinates \( x, y, z \) is

\[ dU = \frac{GM_i dM}{r_{ii}}, \]

where \( G \) is the gravitation constant,

\[ r_{ii} = \sqrt{(x_i - x)^2 + (y_i - y)^2 + (z_i - z)^2} \]

is the distance between the element of mass \( dM \) and the body \( B_i \). Having summarized on all Earth’s mass \( M \) and on all bodies \( n \) the force function is:

\[ U = \sum_{i=1}^{n} GM_i \int \frac{dM}{r_{ii}}, \]

(24)

where \( dM = \rho dxdydz; \quad \rho = \rho(x, y, z) \) is the Earth’s density.

It is obviously impossible to integrate Equation (24) because today’s distribution of the Earth’s density is studied only qualitatively. Therefore direct definition of the force function for the action of external bodies on the Earth as mechanical system is impossible. All further solutions are in simplification of Equation (24) and in derivation of the force function depending on the Earth’s moments of inertia \( J_x, J_y, J_z \). Correlation between them are defined on observable rate of the Earth’s precession.

Taking into account that the distance \( r_i \) between the Earth and the body \( M_i \) is considerably exceeds distance \( r \) to the element \( dM \) the Equation (24) is simplified. Let the line segments \( OD = \xi \) and \( AD = h \). Having applied the theorem of Pythagoras to cathetus \( h \) of both rectangular triangles \( \Delta ODA \) and \( \Delta ADB \), we obtain:
\[ r_{ij} = \sqrt{(r_i - \xi_i)^2 + (r_j - \xi_j)^2} = r_i \sqrt{1 + b_i} \]  

(25)

where \( b_i = -2r_i \xi_i + r_i^2 \).

As \(|b_i|<1\), then subintegral function \( 1/r_{ij} \) in (24) is represented as Taylor series at \( b \) powers. Taking into account terms that are no more the fourth order in relation to \( \sqrt{|b_i|} \), the force function is in form

\[
U = \sum_{i=1}^{n} GM_i \left[ \frac{\xi_i}{r_i} + \frac{3\xi_i^3 - r_i^2}{2r_i} + \frac{5\xi_i^5 - 3\xi_i^3 r_i^2}{2r_i^3} \right] + \frac{35\xi_i^4 - 30\xi_i^2 r_i^2 + 3r_i^4}{8r_i^5} \]  

\[
\frac{dM}{8r_i^5} \]  

(26)

Line segment \( \xi_i \) (Figure 2) is expressed through coordinates \( x, y, z \) of the element \( dM \) by means of direction cosines \( \alpha_i, \beta_i, \gamma_i \):

\[
\xi_i = \hat{l}_i x + \hat{m}_i y + \hat{n}_i z \]  

(27)

where

\[
\hat{l}_i = \cos \alpha_i = x_i/r_i ,
\hat{m}_i = \cos \beta_i = y_i/r_i ,
\hat{n}_i = \cos \gamma_i = z_i/r_i . \]  

(28)

By definition, the axes of a body, which are on intersection of planes of symmetry, are the principal axes of inertia. Axes of rotating system \( \text{Oxyz} \) are directed along the Earth’s principal axes of inertia. Therefore integrals of type \( \int x^k f(x, y, z) dM \), where \( k \)—an odd integer, and \( f(x, y, z) \)—even function of coordinates, will consist of two equal on size and opposite on sign parts at the intervals \( x<0 \) and \( x>0 \), i.e. these integrals are equal zero in the sum. Because, according to (27), variable \( \xi_i \) is proportional to coordinates \( x, y, z \), the integrals in (26), depending on \( \xi_i \) and \( \xi_i^2 \), are equal zero. Note, that these integrals are necessary to consider in case of the non-symmetrical Earth and absence of symmetry on one of the planes \( \text{zOx} \) or \( \text{zOy} \), i.e. the force function as well as the derivatives of angular momentum (7)-(9) will include centrifugal moments of inertia.

Integral of the first term in (26) is the Earth’s mass. Numerator of the third term with the account (27) is

\[
3\xi_i^2 - r_i^2 = 3 \left( x_i^2 - \hat{m}_i^2 x_i^2 - \hat{n}_i^2 x_i^2 + \hat{m}_i^2 y_i^2 + \hat{n}_i^2 z_i^2 \right) - x_i^2 - y_i^2 - z_i^2 + 2 \left( \hat{l}_i \hat{m}_i y_i + \hat{l}_i \hat{n}_i x_i + \hat{n}_i \hat{m}_i y_i \right) \]  

(29)

As \( \hat{l}_i^2 + \hat{m}_i^2 + \hat{n}_i^2 = 1 \), thus in (29) it is taken into, that \( \hat{l}_i^2 = 1 - \hat{m}_i^2 - \hat{n}_i^2 \). The terms integration having co-ordinates in the second power gives the moments of inertia and the last ones gives centrifugal moments as \( J_{xy} = \int xy dM \) which are zero for the Earth with principal axes along \( x, y, z \).

After substitution (29) in (26) and integration we obtain the force function without the last fifth term

\[
U = \sum_{i=1}^{n} GM_i \left[ \frac{M}{r_i} + \frac{J_x + J_z - 2J_y + 3m^2 \left( J_x - J_z \right) - 3n^2 \left( J_x - J_z \right)}{2r_i^3} \right] \]  

(30)

In this case the action on the Earth’s rotational motion is defined by the third and the fifth terms in (26) as they depend on Euler angles. Having divided the fifth term on the third, we find that the order of their relation is equal \( R_E^5/r_0^2 \), where \( R_E \)—radius of the Earth. This ratio has the greatest value for the Moon, and is equal \( 3 \cdot 10^{-4} \). Further the bodies’ action on the Earth’s spin is considered with such relative error. Therefore additional effects should be considered, if the relative value of their action has the order \( 3 \cdot 10^{-4} \) and more.

After the direction cosines \( \hat{m} \) and \( \hat{n} \) are substituted to (28), force function (30) becomes

\[
U = \sum_{i=1}^{n} GM_i \left[ \frac{M}{r_i} + \frac{J_y + J_z - 2J_x + 3m^2 \left( J_y - J_z \right) - 3n^2 \left( J_y - J_z \right)}{2r_i^3} \right] \]  

(31)

where \( y_i \) and \( z_i \)—coordinates of the body \( M_i \) in rotating system \( \text{xyz} \).

### 3.4. The Moments of Forces in Euler Angles

We express \( y_i \) and \( z_i \) through coordinates \( x_{1(1)j(1)} \) of system \( x_{1(1)1(1)} \) (Figure 2). Let us draw the line \( OL_1 \) perpendicularly to \( OK \) in the plane \( x_1Oy_1 \). Coordinates of the body \( Bi \) will give projection \( OL_2 = x_i \sin \psi + y_i \cos \psi \) on this line and then coordinate \( z \) of the body \( Bi \) is

\[
z_i = z_{i(1)} \sin \theta - OL_i \sin \theta \]  

(32)

Having defining projections of the coordinates \( x_i \) and \( y_i \) to the axes \( OK \) and \( OL \) analogously, coordinate \( y_i \) is

\[
y_i = -x_i \left( \cos \phi \sin \varphi + \sin \psi \cos \varphi \cos \phi \right) + y_i \left( \cos \psi \cos \varphi \cos \theta + \sin \psi \sin \phi \right) + z_i \sin \theta \cos \phi \]  

(33)

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Earlier we neglected the centrifugal moments of inertia. It is admissible, if the Earth is axisymmetrical, or axes \( x \) and \( y \) lie in symmetry planes. As there is no evidence on presence of such planes we should consider the axisymmetrical Earth with the equal moments of inertia relative to axes \( x \) and \( y \), i.e. \( J_x = J_y \). Therefore we do not present bulky expressions for the non-axisymmetrical Earth. If necessary they may be received on below-mentioned sequence with use of expression (33). For the axisymmetrical Earth, after substitution (32) in (31), taking into account \( J_z = J_x \), force function becomes:

\[
U = \sum_{i=1}^{n} \frac{3GM_i}{2r_i^3} \left[ -\left( J_z - J_x \right) \cdot \left( x_i \sin \theta \sin \psi - y_i \sin \theta \cos \psi + z_i \cos \theta \right)^2 \right]
\]

(34)

After differentiation (34) on Euler angles \( \varphi, \psi, \theta \) and reductions we obtain the moments of forces

\[
m_{\varphi} = \frac{\partial U}{\partial \varphi} = 0
\]

(35)

\[
m_{\psi} = \frac{\partial U}{\partial \psi} = \sum_{i=1}^{n} \frac{3GM_i}{r_i^5} \left( J_z - J_x \right) \cdot \left[ \sin^2 \theta \sin \psi \cos \psi \left( x_i^2 - y_i^2 \right) - x_i y_i \sin^2 \theta \cos (2\psi) \right. \\
\left. + z_i \sin \theta \cos \theta \cdot \left( x_i \cos \psi + y_i \sin \psi \right) \right]
\]

(36)

\[
m_{\theta} = \frac{\partial U}{\partial \theta} = \sum_{i=1}^{n} \frac{3GM_i}{2r_i^3} \left( J_z - J_x \right) \left[ \sin (2\theta) \cdot \left( x_i^2 \sin^2 \psi + y_i^2 \cos^2 \psi - z_i^2 - x_i y_i \sin (2\theta) \right) \right. \\
\left. + 2z_i \cos (2\theta) (x_i \sin \psi - y_i \cos \psi) \right]
\]

(37)

### 3.5. Differential Equations

The law (1) in projection to the axis \( z \) of rotating system \( xyz \) according to (19), has a form \( K_{\varphi z} = \partial U/\partial \varphi \). We consider the rigid Earth during derivation of the equations: \( J_x = \text{const} \). The moment of inertia of the axisymmetrical Earth \( J_{\text{ez}} \) may be changed for the action of redistribution of a water cover, thawing of glaciers, moving of continents, etc. In order to estimate the influence of these factors, it is entered \( J_x \neq \text{const} \). Therefore, according to (35), \( \partial U/\partial \varphi = 0 \), then \( K_{\varphi z} = 0 \) and taking into account (10) at \( J_x = J_{\text{zo}} \) it becomes \( J_{\text{ez}} = 0 \). After integration it is \( J_x \omega_x = \text{const} = J_{\text{zo}}\omega_{\text{zo}} \), that may be written in form

\[
\omega_z = \omega_{\text{zo}} J_{\text{zo}} / J_z
\]

(38)

where \( J_{\text{zo}} \) and \( \omega_{\text{zo}} \) are the Earth moment of inertia and its projection of angular velocity in an initial epoch. Because \( \omega_z = \phi + \psi \cdot \cos \theta \) (according to Equation (12)), thus, taking into account (38), we obtain angular velocity of the Earth’s own spin

\[
\phi = \omega_{\text{zo}} J_{\text{zo}} / J_z - \psi \cdot \cos \theta
\]

(39)

From expression (39) follows that angular velocity of the Earth’s own spin \( \phi \), which is not connected with motion of vector of angular speed \( \omega \), may be changed via redistribution of the Earth’s moment of inertia and alteration of precession rate. Expression (39) for angular velocity of the Earth’s own spin may be used for an estimation of the Earth’s parts motion influence on its angular speed.

Having designated a projection of the Earth’s angular velocity \( \omega_j = \omega_{\text{zo}} = \omega_{\text{ek}} = \text{const} \), expression (39) for the rigid Earth \( J_{\text{ez}} = J_{\text{eo}} = \text{const} \) we rewrite in a kind

\[
\phi = \omega_{\text{ek}} - \psi \cdot \cos \theta
\]

(40)

After substitution of \( \phi \), \( K_{\text{ek}} = 0 \) and the derivatives of the force function (36) and (37) in Equations (22) and (23), differential equations of the Earth’s rotational movement are

\[
\dot{\psi} = -2\psi \omega_{\text{zo}} \cos \theta / \sin \theta + \left( J_{\text{zo}} \omega_{\text{ek}} / J_z \right) \sin \theta \\
- \sum_{i=1}^{n} \frac{3GM_i E_i J_z}{r_i^7 J_z} \left\{ 0.5 \sin (2\theta) \left( x_i^2 - y_i^2 \right) \\
- x_i y_i \cos (2\theta) + z_i \cos \theta \left( x_i \cos \psi + y_i \sin \psi \right) \right\}
\]

(41)

\[
\ddot{\theta} = 0.5 \psi^2 \cos (2\theta) / J_z \left( J_{\text{zo}} \omega_{\text{e}} / J_z \right) \sin \theta \\
- \sum_{i=1}^{n} \frac{3GM_i E_i J_z}{2r_i^7 J_z} \left\{ \sin (2\theta) \left[ x_i^2 \sin^2 \psi \\
+ y_i^2 \cos^2 \psi - z_i^2 - x_i y_i \sin (2\theta) \right] \\
+ 2z_i \left( x_i \sin \psi - y_i \cos \psi \right) \cos (2\theta) \right\}
\]

(42)

where \( E_j = (J_z - J_x) / J_z \) is the Earth’s dynamical ellipticity; \( \omega_{\text{ek}} \) projection of the Earth’s absolute angular velocity to its axis \( z \), and ratio of moments is \( J_z / J_x \). For the Earth’s own spin \( \omega_{\text{e}} = \text{const} \) and \( \phi \) changes according to (40), thus value \( \omega_{\text{e}} \) may be obtained as a result of averaging of the measured values \( \phi \) and \( \psi \cdot \cos \theta \) on time, i.e. \( \omega_{\text{e}} = \phi + \psi \cdot \cos \theta \).

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4. Initial Conditions and the Earth’s Dynamical Ellipticity

We study the action of separate bodies on the Earth, therefore at integration Equations (41) and (42) it is necessary to set initial values for alteration of rates \( \psi \) and \( \theta \), which are determined by influence of each body. At any set of initial values \( \psi \) and \( \theta \) there begins a transient response, which may become whether a steady-state regime in a long time or not to be steady state at all. We are interested in the Earth’s behavior at steady-state regime. Let the second derivatives \( \ddot{\psi} \) and \( \dot{\theta} \) in Equations (41) and (42) are small in relation to the other terms at steady-state regime, i.e. \( \ddot{\psi} \approx 0 \). As \( \dot{\omega}_E \) exceeds the derivatives \( \ddot{\psi} \) and \( \dot{\theta} \) on 6 orders therefore at neglecting terms with \( \ddot{\psi} \times \theta \) and \( \ddot{\psi}^2 \) in Equations (41) and (42) they become

\[
\dot{\theta} \approx \sum_{i=1}^{n} \frac{3GM_i E_d \sin \theta}{2 \omega_{E r_i}^2} \left[ \sin (2\psi) \left( x_{i1}^2 - y_{i1}^2 \right) - 2x_{i1}y_{i1} \cos (2\psi) \right] + 2z_{i1} \frac{\cos \theta}{\sin \theta} \left( x_{i1} \cos \psi + y_{i1} \sin \psi \right)
\]

\[
\ddot{\psi} \approx -\sum_{i=1}^{n} \frac{3GM_i E_d}{2 \omega_{E r_i}^2 \sin \theta} \left[ \sin (2\theta) \right] \left[ x_{i1}^2 \sin^2 \psi + y_{i1}^2 \cos^2 \psi - z_{i1}^2 - x_{i1}y_{i1} \sin (2\psi) \right] + 2z_{i1} \left( x_{i1} \sin \psi - y_{i1} \cos \psi \right) \cos (2\theta) \]

Equations (43) and (44) are identical to the Poisson equations; their approximate solutions are applied in astronomical theories of climate change. We will define initial values of derivatives from these equations, setting coordinates \( x_{i1}, y_{i1} \) and \( z_{i1} \) of bodies acting on the Earth during an initial epoch \( t = 0 \). Initial values of the angles are: \( \psi_0 = 0 \) and \( \theta_0 = \epsilon_0 \), where \( \epsilon_0 = 0.4093197563 \) is the inclination due to an initial epoch, for example 30.12.1949.

There is the unknown parameter \( E_d = (J_z - J_s) / J_z \), which defines relation between the Earth’s moments of inertia. Today’s knowledge of the Earth’s density distribution is insufficient for definition of the moments \( J_z \) and \( J_s \) with a demanded accuracy. Therefore the Earth’s dynamical ellipticity \( E_d \) is defined by comparison of the calculated precession rate with the observations. According to [16,17] the Earth’s precession has clockwise direction relative to the fixed ecliptic and during a tropical year is

\[
\rho_t = 50''.37084 + 0''.00493T
\]  
(45)

where \( T \) is counted in tropical centuries from a fundamental epoch 1900.0 with \( JD = 2415020.3134 \). In an initial epoch it is \( \rho_0 = 50''.3733046''/yr \). Observable average precession rate is described by Equation (45). Now we define its calculated value according to (44). If an initial epoch coincides with an epoch of reference frame, then \( \theta = \epsilon_0 \) and \( \psi = 0 \). Then Equation (44) is written

\[
\ddot{\psi}_0 = -E_d \sum_{i=1}^{n} \frac{3GM_i}{\omega_{E r_i}^2} \left[ \cos \epsilon_0 \left( y_{i1}^2 - z_{i1}^2 \right) - z_{i1} \nu_{i1} \cos \left( 2\epsilon_0 \right) / \sin \epsilon_0 \right]
\]

(46)

where \( \nu_{i1} = x_{i1} / r_i, \) \( \nu_{i1} = y_{i1} / r_i, \)

\[ \nu_{i1} = z_{i1} / r_i \] — dimensionless coordinates of acting bodies.

If a reference epoch does not coincide with an epoch of reference frame, \( \theta \neq \epsilon_0 \), and \( \psi \neq 0 \).

In the course of cyclic relative movement of bodies round the Earth all coordinates change similarly, for example \( -x_{i1\text{max}} \leq x_{i1} \leq x_{i1\text{max}} \). Here \( x_{i1\text{max}} = y_{i1\text{max}} = z_{i1\text{max}} = 1 \), and \( x_{i1\text{max}} = x_{i1} / r_i, \) where \( x_{i1} / r_i \) — the tilt between the body’s orbital plane and equatorial plane. Since Pluto has the most value of the tilt \( i = 0.7 \) [6], for all planets \( x_{i1\text{max}} < 1 \).

The second term in a square bracket (46) will change in limits \( \pm \sin i \cos \left( 2\epsilon_0 \right) / \sin \epsilon_0 \) and averaging on time will make it zero. The maximum value of the factor of the first term \( \left( y_{i1}^2 - z_{i1}^2 \right) / \sin \epsilon_0 \) is 1, and minimum is zero, therefore and averaging on time will give 0.5. Then all the part in square brackets at averaging is \( 0.5 \cos \epsilon_0 \).

Value of geocentric distances to the planets changes in limits \( \left| R_{Gi} - R_{Gi\text{sf}} \right| \leq r_i \leq R_{Gi} + R_{Gi\text{sf}} \), where \( R_{Gi}, R_{Gi\text{sf}} \) and \( R_{Hi}, R_{Hi\text{sf}} \) — perihelions and afelions of the planets and the Earth, accordingly. Then average geocentric distance to the planets is

\[
r_{i\text{es}} = \left( R_{Gi} + R_{Gi\text{sf}} + \left| R_{Gi} - R_{Gi\text{sf}} \right| \right) / 2
\]  
(47)

which is \( r_{i\text{es}} = \left( R_{Gi} + R_{Gi\text{sf}} - \left| R_{Gi} - R_{Gi\text{sf}} \right| \right) / 2 \) for internal planets, \( R_{Gi} + R_{Gi\text{sf}} \) for external, where \( a_i, a_g \), and \( e_s, e_G \) — semimajor axes of heliocentric orbits and orbits eccentricity of the planets and the Earth accordingly. Accordingly, for the Sun and the Moon average geocentric distances are

\[
r_{s\text{es}} = a_e, \quad r_{m\text{es}} = a_M
\]  
(48)

where \( a_M \) — average semiaxis of the Moon’s orbit. Then average value of speed of precession, according to (46), one gets

\[
\ddot{\psi}_{i\text{es}} = -E_d \sum_{i=1}^{n} \frac{1.5GM_i \cos \epsilon_0}{\omega_{E r_i^{i\text{es}}}}
\]

(49)

The negative sign testifies clockwise direction of precession. Due to equality condition between calculated
average precession rate and observed one

\[ \psi_{0w} = -p_{10}/(Y_p \cdot rd), \]

where

\[ Y_p = 365.25636042 \cdot 24 \cdot 3600 \text{ — duration of a sidereal year in sec, and } \]

\[ rd = 180 \cdot 3600/\pi \text{ we find the dynamical ellipticity} \]

\[ E_{da} = p_{10}/\left( Y_p \cdot rd \cdot \sum \frac{1.5GM_i \cos \theta_i}{a_i r_{im}^3} \right) \] (50)

where index \( a \) and \( E_{da} \) denote approximate method of calculation of this value.

Having derivated Equations (49) and (50) we averaged enough approximately expression (46) on time, for example, the average values \( r^3_i \) and \( r_i \) can essentially differ. If laws of planets motion \( x_i(t), y_i(t), z_i(t) \) could be found, there were known average value \( \psi_0 \) on time span \( T \) by integrating expression (46) on time at interval from 0 to \( T \). However, expression (46) is approximate since it is obtained from the equations of motion at neglecting \( p \) and \( \theta \). Therefore Equation (50) represents the Earth’s dynamical ellipticity approximately. Its more exact value will be obtained after numerical solution of the Earth’s rotational movement Equation (41) and (42) and comparisons of averaged from short-period oscillations of dynamics \( p \) during known observation with observable dynamics of average precession for this period.

The value of dynamical calculated according to (50) under the data used by us, is \( E_{db} = 3.324257 \times 10^{-3} \). Traditionally dynamical ellipticity is calculated with the account only the Moon and the Sun. In this case, under our data \( E_{dsm} = 3.324268 \times 10^{-3} \) i.e. differs at unit of the sixth positional notation. As it is seen from (50), the dynamical ellipticity depends on averaged distance \( r_{im}^3 \), thus the small changes of its value sufficiently change \( E_{da} \).

For instance value of this distance for the Moon \( r_{im} = a_M = 3.838 \times 10^{-3} \text{ m} \), which we obtained at reference conditions at epoch 30.12.49 from ephemerides DE406, gives dynamical ellipticity \( E_{dsm} = 3.324770 \times 10^{-3} \). To comparison here are ellipticities used by various authors: Bretagnon and Simon [11] \( E_{db} = 3.273767 \times 10^{-3} \); F. Roosbeek and V. Dehant [18] \( E_d = 3.273767 \times 10^{-3} \); Laskar et al. [5] \( E_d = 3.28005 \times 10^{-3} \) with using \( a_M = 38447.980 \text{ m} \) and \( p_{10} = 50.290966 \text{ rad/yr} \).

As we see, values \( E_d \) in expression (50) change depending on the accepted data in the second significant figure. The value dynamical ellipticity, taking into account the data accepted by us, is designated as \( E_{db} \), as the sun.

Table 1 gives Earth’s axis mean precession rates for the action of the Sun, the Moon and the planets calculated according to (49) and \( E_d = E_{db} \). The most value of speed of precession is caused by the Moon, the action of the Sun is twice less, and Venus and Jupiter cause the greatest action from the planets. As it is seen from Table 1, difference between speed of precession \( \psi_{0w} \) for the action of all the bodies and observable value \( p_{10} \) is 3.14 \( \times 10^{-3} \% \).

Due to accepted dynamical ellipticity \( E_{ds} \) and \( \psi_0 = 0 \), initial values \( \psi_0 \) and \( \theta_0 \) are calculated according to (43) and (44) for the action of the separate bodies and all together, see Table 1. Note, that expressions (43) and (44) represent values of derivatives \( \psi_0 \) and \( \psi_0 \) approximately. Their exact value is given by observation. There are no these values for the action of separate bodies but there are observation data for the simultaneous action of all the bodies on the Earth. However, as Equations (43) and (44) are approximate, total values \( \psi_0 \) and \( \psi_0 \) (taking into account all the bodies) will differ from the observation data.

5. Solution Algorithm

The coordinates \( x_{10}, y_{10}, z_{10} \) of the bodies \( M_i \) acting the Earth in Equations (41) and (42) are referred to the fixed geocentric ecliptic coordinates (Figure 2). Numerical solution of the orbital problem gives the change of their orbit parameters \( i, \phi_{2}, \phi_{3}, R_p, e \) over a time span 100 Myr [6]. Due to this solution there are the bodies’ positions saved with 10 kyr interval. To integrate equations of rotational motion one needs the bodies’ positions during any moment of time.

The bodies’ positions are usually set on the base of the theory of secular perturbations during solving rotational motion equations. Bretagnon et al. [11] used ephemerides DE403 for numerical integration of these equations for 1968-2023 yr interval. These methods are not suitable

<table>
<thead>
<tr>
<th>No.</th>
<th>Bodies</th>
<th>( \psi_{0w} )</th>
<th>( \psi_0 )</th>
<th>( \theta_0 )</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>Sun</td>
<td>-7.73095 ( \times 10^{-3} )</td>
<td>-1.20483 ( \times 10^{-9} )</td>
<td>-7.09773 ( \times 10^{-8} )</td>
</tr>
<tr>
<td>2</td>
<td>Mercury</td>
<td>-1.02693 ( \times 10^{-9} )</td>
<td>1.42035 ( \times 10^{-10} )</td>
<td>2.95303 ( \times 10^{-11} )</td>
</tr>
<tr>
<td>3</td>
<td>Venus</td>
<td>-1.86489 ( \times 10^{-10} )</td>
<td>-7.17466 ( \times 10^{-9} )</td>
<td>-2.99417 ( \times 10^{-9} )</td>
</tr>
<tr>
<td>4</td>
<td>Mars</td>
<td>-6.82598 ( \times 10^{-10} )</td>
<td>-8.87394 ( \times 10^{-12} )</td>
<td>1.26014 ( \times 10^{-10} )</td>
</tr>
<tr>
<td>5</td>
<td>Jupiter</td>
<td>-5.19060 ( \times 10^{-10} )</td>
<td>1.58365 ( \times 10^{-9} )</td>
<td>4.60379 ( \times 10^{-10} )</td>
</tr>
<tr>
<td>6</td>
<td>Saturn</td>
<td>-2.53056 ( \times 10^{-10} )</td>
<td>1.06473 ( \times 10^{-10} )</td>
<td>2.27373 ( \times 10^{-10} )</td>
</tr>
<tr>
<td>7</td>
<td>Uranium</td>
<td>-4.74889 ( \times 10^{-11} )</td>
<td>1.45508 ( \times 10^{-12} )</td>
<td>2.78046 ( \times 10^{-14} )</td>
</tr>
<tr>
<td>8</td>
<td>Neptune</td>
<td>-1.46677 ( \times 10^{-11} )</td>
<td>-6.38274 ( \times 10^{-13} )</td>
<td>8.18066 ( \times 10^{-13} )</td>
</tr>
<tr>
<td>9</td>
<td>Pluto</td>
<td>-6.92755 ( \times 10^{-14} )</td>
<td>3.47064 ( \times 10^{-14} )</td>
<td>1.91721 ( \times 10^{-14} )</td>
</tr>
<tr>
<td>10</td>
<td>Moon</td>
<td>-0.01657</td>
<td>-1.83206 ( \times 10^{-10} )</td>
<td>-9.29602 ( \times 10^{-10} )</td>
</tr>
<tr>
<td>All bodies</td>
<td>-2.44209 ( \times 10^{-2} )</td>
<td>-</td>
<td>-</td>
<td></td>
</tr>
</tbody>
</table>
for long-time integration. Therefore we developed the follow algorithm to compute the motion laws of the bodies $x_i(t), y_i(t), z_i(t)$. Their coordinates are defined on the base of our solutions about motion of the Solar system bodies in fixed barycentric equatorial coordinates. Using orbital parameters $i, \phi_{2i}, \phi_{n}, R_p$, e Cartesian coordinates are recalculated in polar $r_i$ and $\phi_i$ in orbital plane. To estimate the body position during new moment of time $t_1 = t + \Delta t$ one may use expressions for elliptical motion given in [12,13,20,21]: $t = t(r)$—motion law in implicit form and $r(\theta)$ — trajectory equation. Polar radius of the body $r_i$ during new moment of time $t_1$ is predicted under Taylor’s formula of the second order, checked on dependence $t_i(r_i)$ and then discrepancy for some iterations comes almost to zero. After that from the trajectory equation there obtained polar angle $\phi_i(r_i)$ in a return form. At small ellipticities $e \cdot 0$ during new moment of time $t_1$ under Taylor’s formula the angle $\phi_i$ is predicted, checked under the movement law $t_i(r(\phi_i))$, then the discrepancy for some iterations also practically comes to zero.

This algorithm differs from the traditional. The intermediate parameters: mean anomaly $M$ and eccentric anomaly $E$, which feature in traditional description of body motion on an elliptic orbit, are not used here.

The bodies polar coordinates $r_i, \phi_i$ during any moment of time are recalculated in coordinates of translationally moving geocentric system $x_i y_i z_i$ via orbital parameters $i, \phi_{2i}, \phi_{n}, R_p$. Since the orbital parameters change from one epoch to another, they are interpolated on a parabola on each step. Thus, the algorithm, which allows defining the bodies coordinates $x_i, y_i, z_i$ acting the Earth at any moment from the considered interval of time, is developed. Note, that the Moon’s orbit has more short oscillation periods comparably with the other bodies of planets. Therefore the algorithm for the Moon was modified upon with their account.

We applied the Runge-Kutta integration of Equations (41) and (42). These algorithms of calculation of coordinates and numerical integration are programmed on the FORTRAN. The computation was run on personal computers and supercomputers MVS-1000 at Keldysh Institute of Applied Mathematics of the Russian Academy of Sciences (Moscow) and NKS-30T at Siberian Branch of the RAS Computing Center (Novosibirsk). As a result of numerical experiments we accepted the initial integration step $\Delta t = 1 \cdot 10^{-4}$ yr.

6. Numerical Integration Results

6.1. The Sun’s Action

Figure 3, a gives results of integration Equations. (41) and (42) over 0.1 yr for the action of the Sun on the Earth’s rotational motion. The precession angle $\psi$ decreases since zero value making fluctuations with the amplitude $\psi_{01} = 5.63 \times 10^{-8}$ rad = 13.47 mas and the period $T_1 = 0.99348 \; d$, where: $mas$ is $10^{-3}$ angular seconds, $d$ is duration of a day in sec. The nutation angle is presented as difference $\Delta \theta = \theta - \theta_0$, where $\theta_0 = 0.4931976 \; rad$. As it is seen from the graph $\Delta \theta(T)$ value of the tilt between moving equatorial plane and fixed ecliptic increases, making fluctuations with amplitude $\psi_{01} = 2.60 \times 10^{-8}$ rad = 5.27 mas and same period $T_1$. Rate of precession $\psi$ oscillates round some mean value with period $T_1$ and amplitude $\psi_{01} = 1.57 \times 10^{-2}$ rad/cyr. Rate of nutation oscillates with the same period $T_1$ and amplitude $\psi_{01} = 6.17 \times 10^{-3}$ rad/cyr round almost zero value.

As it is seen from Figure 3, a initial values of the rates $\theta_0$ and $\psi_0$ do not exceed their extreme values. Therefore an error in definition $\theta_0$ and $\psi_0$ causes only change of an initial phase of fluctuations. Hence, in this case dynamics of the rates $\theta_0$ and $\psi_0$ is not changed at an error of initial conditions described in p. 14. There are similar properties of solutions for the action of other bodies, thus this property will be kept for the action of all bodies. Consequently, the initial conditions in this case calculated according to (43), (44), may be specified on a phase of observed daily oscillations $\theta$ and $\psi$.

As the accepted initial conditions, according to (43), (44), are artificial, there arises the question. Which values of derivatives $\dot{\theta}_0$ and $\dot{\psi}_0$ might be, if a long-time action of a single body on the Earth preceded the beginning of calculation? Apparently, there would be an occurrence of equilibrium and derivatives would come nearer to the mean values. With this purpose there were defined mean values $\psi_{01} = 1.61891 \times 10^{-2}$ and $\theta_{01} = -1.32398 \times 10^{-3}$ according to the first daily oscillations (see Figure 3(a)). Integrating Equations (41)-(42) with initial values $\psi_{01}$ and $\theta_{01}$ has resulted in reduction of daily oscillations in 20 times (see Table 2). Hereat character of solutions behavior for the long-time actions has not changed.

At averaging derivatives in new solutions there obtained mean values $\psi_{02} = 1.59342 \times 10^{-2}$ and $\theta_{02} = -1.08328 \times 10^{-3}$, submitted in Table 2. Repeated calculation with these initial values has resulted in diminution of amplitudes of oscillations in several times. Thus, there were executed 4 calculations with specification of the mean values of initial rates. Consequently, amplitudes of daily oscillations have decreased in several hundred times, according to Table 2.

It is possible to account the final values of derivatives $\psi_0$ and $\theta_0$ as steady-state for long-time action of a single body on the Earth. For the action on the Earth
several bodies there, apparently, will not be such steady-state condition for each body. Therefore further we consider the action of single bodies at initial conditions, according to (43), (44).

Figure 3(b) shows the Earth’s axis dynamics for 1 yr. The precession angle increases clockwise making oscillations with period $T_2 = 0.5$ yr and amplitude $\psi_{min} = 5.98 \times 10^{-6}$ rad $\cdot$ yr$^{-1}$ in the reference epoch (30.12.1949). It is explained by the most deviation of the Sun from the equatorial plane, which makes the most moment of force due to inertia. Therefore $\psi$ changes more than in 20 times relative to an average. The rate $\dot{\psi}$ changes in more relative bounds in relation to its mean value $\dot{\psi}_{min} \approx 0$. However, at such considerable changes of the instant rates $\psi$ and $\dot{\psi}$ the mean values remain constants during all integration interval. It testifies stability of the integration method. And as mean values $\psi_{max}$ correspond with observations, it testifies reliability of the solution of the entire problem.

Note, that under consideration of the action of the body $B$ on the Earth (Figure 1) there was shown that the torque $m_0 < 0$ and the precession rate $\psi < 0$. As a result of integration (41), (42) we have received that precession rate takes positive values also (Figure 3(b)). This result is caused by dynamics: under the action of torque the Earth’s axis speeds up, according to (4), and its motion proceeds even in absence $m_0$ due to inertia. Therefore $\psi_{max}$ is more than the value given by (4), and $\psi_{min} > 0$.

Here we do not submit graphs of the Earth’s angular velocity dynamics $\phi$, since, according to (40), the hanges of $\phi$ have contrary sign to the changes of the precession rate $\psi$. This conclusion is proved by results of other authors, for example, [9].

Table 2. Amplitudes of daily oscillations for the action of the Sun at specification of the steady-state initial rates.

<table>
<thead>
<tr>
<th>No.</th>
<th>Initial rates</th>
<th>Daily amplitudes</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>$\psi_0$</td>
<td>$\dot{\psi}_0$</td>
</tr>
<tr>
<td>1</td>
<td>$-1.20483 \times 10^{-6}$</td>
<td>$-7.09773 \times 10^{-6}$</td>
</tr>
<tr>
<td>2</td>
<td>$-1.61891 \times 10^{-7}$</td>
<td>$1.32398 \times 10^{-3}$</td>
</tr>
<tr>
<td>3</td>
<td>$-1.59342 \times 10^{-7}$</td>
<td>$1.08328 \times 10^{-3}$</td>
</tr>
<tr>
<td>4</td>
<td>$-1.59710 \times 10^{-7}$</td>
<td>$9.09860 \times 10^{-4}$</td>
</tr>
</tbody>
</table>

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6.2. The Moon’s Action

The moon has the strongest action on the Earth’s axis dynamics. Figure 5(a) gives the daily oscillations of the parameters with the period $T_1 = 0.99348 \text{ d}$ and amplitudes $\psi_{\alpha_1} = 7.29 \cdot 10^{-8} \text{ rad} = 15.04 \text{ mas}$, $\theta_{\alpha_1} = 2.83 \cdot 10^{-6} \text{ rad} = 5.83 \text{ mas}$, $\psi_{\alpha_1} = 1.66 \cdot 10^{-2} \text{ rad/cyr}$ and $\theta_{\alpha_1} = 6.63 \cdot 10^{-3} \text{ rad/cyr}$. Against daily oscillations there appear half-monthly oscillations with period $T_2 = 13.5143 \text{ d}$, which are well visible on 1 yr interval (see Figure 5(b)). The precession angle has a linear trend and also oscillates with the half-monthly period and amplitude $\psi_{\alpha_2} = 1.16 \cdot 10^{-6} \text{ rad} = 240.10 \text{ mas}$. For 10 yr interval the parameters behave analogically, therefore these results are not submitted here.

The further changes of parameters we will consider for 100 yr interval (Figure 5(c)). Here there are new oscillations with period $T_3 = 18.6 \text{ yr}$. Along with linear the change of precession angle it has oscillations with the period $T_3$ and the amplitude $\psi_{\alpha_3} = 0.98 \cdot 10^{-4} \text{ rad} = 20286.33 \text{ mas}$. The nutation angle oscillates with this period at amplitude of fluctuations $\psi_{\alpha_3} = 5.23 \cdot 10^{-5} \text{ rad} = 10797.80 \text{ mas}$.

As it was told in p. 2, due to the analysis of the theorem of momentum the presence of the oscillations caused by orbital precession of the acting body $B$ (Figure 1(c)) has been established. At integration (41), (42) with the Moon action to the Earth, we have obtained oscillations with period $T_3 = 18.6 \text{ yr}$, which are caused by precession of the Moon orbital axis.

For 1000 yr interval the behavior of the precession $\psi$ and the nutation $\theta$ angles are similar to Figure 5(c). Except oscillations with the period $T_3 = 18.6 \text{ yr}$ others are not present. The mean precession rate is $\psi_{\alpha_3} = 1.70 \cdot 10^{-2} \text{ rad/cyr}$. Oscillations with the period $T_3 = 18.6 \text{ yr}$ also proceed on 10 kyr interval. The trend of change of the nutation angle $\theta$ under the cosine law is tracked.

Figure 4. Action of the Sun on the Earth’s rotation: $a$—during 10 yr; $b$—during 100 yr. See captions on Figure 3.

Figure 5. Action of the Moon on the Earth’s rotation: $a$—during 0.1 yr; $b$—during 1 yr; $c$—during 100 yr. See captions on Figure 3.
6.3. Action of Venus and Other Planets

Equations (41) and (42) were integrated separately for the action of all the planets. The Venus makes the most influence on the Earth’s axis. Figure 6(a) shows the Earth’s axis dynamics under the action of the Venus on 0.1 yr interval. The precession and the nutation angles oscillates with daily periods and amplitudes $\psi_{a1} = 2.15\cdot10^{-12}$ rad = $0.443\cdot10^{-3}$ mas; $\theta_{a1} = 0.855\cdot10^{-12}$ rad = $0.176\cdot10^{-3}$ mas. As we see, the daily amplitudes for the Venus action are in 4 orders less than for the Sun action.

The rates $\dot{\psi}$ and $\dot{\theta}$ also oscillate with daily period $T_1 = 0.99348$ d and the amplitudes considerably exceeding a mean values. As amplitudes of daily oscillations shade oscillations with bigger periods, we do not submit $\psi$ and $\theta$ diagrams on the subsequent intervals.

Figure 6(b) represents 10 yr dynamics of the precession and nutation. $\Delta \theta$ change shows that the second period of oscillations with duration about 1.6 yr is tracked. These oscillations are expressed by horizontal lines of the change $\Delta \theta$. Their period $T_2 = 1.6$ yr is caused by periodicity of approach of Venus with the Earth due to their relative motion. Period of these approaches is $T_{BE} = T_E T_V / (T_E - T_V) = 1.598$ yr, where $T_E$ and $T_V$ are sidereal periods of the Earth and Venus. The graph $\Delta \theta$ on Figure 7(a) shows 1 cyr interval where oscillations with the period $T_2 = T_{BE}$ are made against oscillations with the bigger period $T_3 = 8.12$ yr and the amplitudes $\psi_{a2} = 1.65\cdot10^{-9}$ rad = $0.341$ mas and $\theta_{a2} = 6.16\cdot10^{-10}$ rad = $0.127$ mas. These oscillations with the period $T_3$ are caused by periodicity of approach of Venus with the Earth at the longest distance from the equatorial plane ($B_1$ and $B_3$ on Figure 1(c)). Really, after the first approach, for example in the point $B_1$, following approach will be in $T_2 = 1.6$ yr, i.e. will be aside of the point $B_2$ for 0.1 yr. Therefore 0.5 yr/0.1 yr = 5 approaches, i.e. $T_3 = 5 T_2 \approx 8.1$ yr so that the last approach will occur again in the point $B_1$ or $B_2$ is required.

Periodicity of approach in points $B_1$ and $B_2$ can be given by expression

$$T_{BC} = 0.5 T_{BE} / \left( T_{BE} - 0.5 \text{Int} (2 T_{BE}) \right)$$

where $T_{BE}$ is the period of approach of the body B with the Earth, and function Int defines an integer part of a number.

Figure 7(b) shows 10 kyr $\psi$ and $\theta$ dynamics. The angles change linearly with the mean (for 10 kyr) rates $\dot{\psi}_{mn} = -9.53\cdot10^{-8}$ rad/cyr and $\dot{\theta}_{mn} = 8.05\cdot10^{-9}$ rad/cyr, and the nutation rate is in 12 times less than the precession rate. Difference of the mean nutation rate from zero on 10 kyr interval is caused by the influence of precession of Venus orbital plane relative to the Earth’s equatorial plane during this time. Apparently, further it will
lead to oscillations with a new period.

Graphs of the other planets action on the Earth’s rotational motion look similarly to graphs for the Sun, the Moon and Venus. The most difference proves for the nutation angle. Figure 8 represents the comparison of influence of the Sun and the planets on the nutation angle. Under the action of the Sun, Mercury (Figure 8) and Venus (Figure 7) the nutation angle increases. The action of the other planets leads to its reduction. These changes of the angle \( \theta \) occur, most likely, via the orbital precession of the acting bodies. In the beginning they change linearly for all the bodies (under the sine law), except for the Sun and the Moon. The semi-annual nutation oscillations for the action of the Sun (Figure 4) are big enough and shade the trend \( \theta \). Therefore, Figure 8 presents the change of sliding average \( \Delta \theta_s \) on 100 next points. From the diagram follows that \( \Delta \theta_s \) changes, apparently, under the cosine law with average rate for 1000 yr \( \dot{\theta}_{s,0} = 3.3 \times 10^{-8} \) rad/cyr.

There are traced different kinds of the nutation oscillations. All the bodies have the daily oscillations with the period \( T_1 \). Oscillations with the semi-orbital period \( T_2 \) are brightly expressed for the Sun and external planets, except for Pluto, which has the oscillations \( \theta \) with the period \( 2T_2 \). For near planets from Mercury to Mars there are oscillations having horizontal maxima with orbital periods \( T_{BE} \) of the bodies (B) relative to the Earth (E), for example, 1.6 yr for Venus and 2.14 yr for Mars. Horizontal parts of graphs of the angle \( \theta \) are caused by long contact of the planet and the Earth during approach. There are brightly expressed oscillations with the period \( T_3 = T_{BC} = 8.1 \) yr for Venus, which are caused by approach with the Earth in the points \( B_1 \) and \( B_3 \). For Mars the doubled periods with \( T_3 = 2T_{BC} = 2 \times 7.9 = 1.58 \) yr are received. The cause of the periods doubling is the result of that the action of the bodies in diametric opposition points relative to the Earth differs due to eccentricity of the orbits and their inclination to the equatorial plane. Therefore the action in one of the points becomes prevailing, thereby the period of oscillations is fixed on it.

The periods and the amplitudes of the Earth’s axis oscillations and the rates of precession \( \psi_m \) caused by the action of the Sun, the Moon and the planets are given in Table 3. The Moon acts mostly on the amplitudes and the precession rate. Venus renders the most action among the planets.

Having comparing precession rate \( \dot{\psi}_m \) with the mean value \( \dot{\psi}_m \) calculated under the Equation (49), it is visible (Table 1) that they are close for the Sun, the Moon and the far planets from Jupiter to Neptune. The approximate Equation (49) gives big errors for the close planets via strong influence on averaging of the planets approach with the Earth. Pluto’s error at averaging arises via the big obliquity of its orbit and big eccentricity.

Table 3 gives oscillations of several types. The first, with daily period \( T_1 = 0.99348 \) d, is caused by the Earth’s rotational motion. Here the greatest amplitudes have the rates \( \dot{\psi} \) and \( \dot{\theta} \). Other kinds of oscillations are caused by periodicity of intersection of the Earth’s equatorial plane by the bodies, periodicity of approaches of the planets with the Earth and periodicity of approaches in points of the greatest removal from the equatorial plane (points \( B_1 \) and \( B_3 \) on Figure 1). In these cases the angles \( \psi \) and \( \theta \) reach the greatest amplitudes. At the Moon action the greatest amplitudes of oscillations are caused by precession of its orbital plane.

7. Comparison with Works of Other Researchers

Comparison of the obtained differential Equations (41) and (42) with the differential equations of other authors...
becomes complicated as the majority of authors do not submit these equations in a final kind. With the purpose of their approximate solution the differential equations of the 2nd order are simplified to equations of the first order (the Poisson equations). Additionally, different reference systems, different variables and different designations are used, and no one work known to us has the equations in a final form similar to (41), (42). P. Bretagnon et al. [9] gives the fullest kind of the differential equations from the latest works. At substitution in the equations of these authors the moment \( J_2 = J_3 \), and angular velocity according to (40), terms, caused by the angular momentum, coincide with corresponding terms of our Equations (41) and (42). These authors express terms, caused by the torques, by Legendre polynomials, therefore they cannot be compared with our terms.

Our solution (40) for the Earth’s angular velocity \( \dot{\phi} \) and kinematic terms of Equations (41) and (42), taking into account the signs of \( \psi \) and \( \theta \), also coincide with the Equations (9.1.04) and kinematic terms of Equations (9.1.05) given by G. Duboshin [17], if to consider two typing errors in Equations (9.1.05). These typing errors in the subsequent Poisson equations (9.1.06) are eliminated. By comparison of the equations, typing errors in other works are found out. This circumstance is an additional argument in necessity of the work on the derivation of the Earth’s rotational motion equations.

Our derivation of the torques coincides with a traditional derivation [15] up to expression (31) inclusively before the force function \( U \). After that the values \( r^2_y \) and \( z/r \) are usually represent in series on elements of elliptical motion on the basis of the theory of secular perturbations. So these results also are non-comparable to ours.

As a result it is possible to make the following conclusion. The terms of Equations (41) and (42), caused by the angular momentum, coincide with the equations of other researchers. The terms of the equation, caused by the torques, also coincide at identical stages of the derivation.

The accepted way of derivation of the equations, the approach to their solution, the initial condition setting and the final kind of (41), (42) do not repeat the equations, known to us in the literature. Despite originality of these methods for the Earth’s rotational motion equations, they are typical in the mechanic and are traditionally used in its different areas.

In our solutions daily oscillations of the angles of precession \( \psi \) and nutation \( \theta \) are received. P. Bretagnon et al. [9] carried out numerical integration for check of the analytical solutions. Differences between the numerical and analytical results, submitted in their work, have daily oscillations, i.e. numerical solutions of these authors also give daily oscillations. Daily frequencies of oscillations \( \psi \) and \( \theta \) also exist in the observation data presented in the form of series [8].

Table 4 gives the comparison of the amplitudes and the periods of the Earth’s axis oscillations of the second and third types received by us (Sm) with the results of Bretagnon et al. [9] analytical solutions (Br). The mean precession and nutation rates, caused by the action of the Sun, the Moon and the planets, are compared here also. The amplitudes of oscillations \( \psi_a \) and \( \theta_a \), their period \( T \) for the Sun and the mean precession speed \( \psi_{20} \) coincide well. It is necessary to notice that the rate \( \theta_m \), according to Figure 8, changes nonlinearly, and its almost zero mean value during the initial moment against considerable semi-annual oscillations cannot be defined precisely. Due to this reason it is impossible to obtain the Moon’s rate \( \theta_m \).

The periods of oscillations \( T \) of the angles, as we see from Table 4, well coincide for all the planets. Amplitudes of oscillations \( \psi \) and \( \theta \) are also well harmonized for the given planets from Jupiter to Neptune. Smaller conformity for the near planets is explained by difficult structure of oscillations \( \psi \) and \( \theta \). In work [9] amplitudes are defined by representation of solutions as

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**Table 3. Periods of the Earth’s axis oscillations, their amplitudes and mean nutation rates at separate action of the Sun, the Moon and the planets:**

<table>
<thead>
<tr>
<th>Bodies</th>
<th>( T_1, \text{d} )</th>
<th>( \psi_{a1} ), rad</th>
<th>( \theta_{a1} ), rad</th>
<th>( T_2 (T_1) ), yr</th>
<th>( \psi_{a2, a3} ), rad</th>
<th>( \theta_{a2, a3} ), rad</th>
<th>( \psi_m ), rad/cyr</th>
</tr>
</thead>
<tbody>
<tr>
<td>Sun</td>
<td>0.99348</td>
<td>6.52987·10^{-8}</td>
<td>2.60203·10^{-8}</td>
<td>0.5/0.5</td>
<td>5.97569·10^{-6}</td>
<td>2.69668·10^{-8}</td>
<td>-7.73671·10^{-3}</td>
</tr>
<tr>
<td>Mercury</td>
<td>0.99348</td>
<td>1.03196·10^{-14}</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
</tr>
<tr>
<td>Venus</td>
<td>0.99348</td>
<td>2.15159·10^{-12}</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
</tr>
<tr>
<td>Mars</td>
<td>0.99348</td>
<td>2.78597·10^{-16}</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
</tr>
<tr>
<td>Jupiter</td>
<td>0.99348</td>
<td>2.32699·10^{-11}</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
</tr>
<tr>
<td>Saturn</td>
<td>0.99348</td>
<td>1.01170·10^{-13}</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
</tr>
<tr>
<td>Uranium</td>
<td>0.99348</td>
<td>4.65987·10^{-16}</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
</tr>
<tr>
<td>Neptune</td>
<td>0.99348</td>
<td>5.34857·10^{-17}</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
</tr>
<tr>
<td>Pluto</td>
<td>0.99348</td>
<td>5.18906·10^{-19}</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
<td>disp.</td>
</tr>
<tr>
<td>Moon</td>
<td>0.99348</td>
<td>7.28690·10^{-1}</td>
<td>2.28844·10^{-8}</td>
<td>disp.</td>
<td>18.6/3.7·10^{-2}</td>
<td>0.98351·10^{-4}</td>
<td>5.23492·10^{-5}</td>
</tr>
</tbody>
</table>

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harmonious series; their main harmonics are given in Table 4. In our work actual amplitudes are defined by a sliding average method. These methods yield different results at difficult structure of oscillations.

The mean rates of change Ψm and θm at action of the planets, as we see from Table 4, are well correlated. Note, that exact coincidence of our and P. Bretagnon et al. [9] results is impossible, since there are difference in ways of parameters definition and in the initial data also. For example, at identical dynamical ellipticity Ed = EdS our mean rate of precession for the action of the Moon is -3445.90, which practically coincides with P. Bretagnon et al. [9] result (-3445.50), see Table 4.

As the solutions of P. Bretagnon et al. [9] are normalized by the observation data, coincidence of our results with results of these authors testifies their confirmation by observation data. The coincidences of our results (given to Table 4) testify reliability of the obtained differential equations and their solutions. This coincidence testifies also the correctness of the initial data. The mathematical model of motion of the planets and the Moon, which gives their positions in process of numerical integration of Equations (41) and (42), proves to be true also. It is necessary to notice also that in [9] the Earth is considered as non-axisymmetrical, power function is estimated by potential of the surface distribution of gravity and geodetic precession is taken into account. All these additional influences give additives which are in limits of differences of our solutions from Bretagnon et al. [9].

8. Conclusions

1) The law of angular momentum change is analyzed and on its basis the differential equations of the Earth’s rotational motion for the gravitational action of the Solar system bodies is derived.

2) The algorithm and the program of numerical inte-
Presidium of the Russian Academy of Sciences.

10. References


