Exact Solution of Terzaghi’s Consolidation Equation and Extension to Two/Three-Dimensional Cases

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Abstract
The differential equation by Terzaghi and Fröhlich, better known as Terzaghi’s one-dimensional consolidation equation, simulates the visco-elastic behaviour of soils depending on the loads applied as it happens, for example, when foundations are laid and start carrying the weight of the structure. Its application is traditionally based on Taylor’s solution that approximates experimental results by introducing non-dimensional variables that, however, contradict the actual behaviour of soils. The proposal of this research is an exact solution consisting in a non-linear equation that can be considered correct as it meets both mathematical and experimental requirements. The solution proposed is extended to include differential equations relating to two/three dimensional consolidation by adopting a transversally isotropic model more consistent with the inner structure of soils.

Keywords: Terzaghi; One-Dimensional Consolidation; Pore Overpressure; Two/Three-Dimensional Consolidation

1. Introduction
The mechanical behaviour of soils was coded only after the introduction of the “concept of effective stresses” [1] which marked the birth of Soil Mechanics starting from the general structure of Continuum Mechanics from which it derives. This concept is based on the inner structure of soils that are composed of a solid skeleton and inter-particle gaps. These pores are more or less interconnected and through them run fluids of different nature. Therefore, in view of a necessary simplification of the mathematics of associated phenomena, the concept of effective stresses requires the soils to be assimilated to bi-phasic systems composed of a solid skeleton saturated with water, i.e. two continuous means that act in parallel and share the stress status:

$$\sigma'_{ij} = \sigma_{ij} - u_0 \delta_{ij}. \quad (1)$$

In Equation (1) there is the tensor of the total stresses exerted by the solid skeleton $\sigma_{ij}$, the hydrostatic pressure exerted by the fluid $u_0$, known as interstitial pressure) and Kronecker’s delta $\delta_{ij}$; furthermore, from a merely phenomenological point of view, Equation (1) attributes the soil shear resistance only to effective stress, independent of the presence of the fluid.

It should be highlighted that the concept of effective stresses is valid only in stable conditions, when the fluid is in balance with the solid skeleton. In these conditions you can calculate the hydrostatic component in all points of the underground and in all moments through the application of the laws of balance; vice-versa, transient conditions exist necessary to introduce other elements capable of accounting for the variation of the component $u_0$ induced by stresses of various nature.

At this point, the problem focus is on the permeability coefficient (K = m/sec) that, by expressing the capacity of a soil to transmit a fluid, takes on the character of a velocity and varies approximately in the range $10^{-1}/10^{-10}$ m/sec depending on the inner structure of the solid skeleton. As a direct consequence of this extreme variability, the soils with high permeability (such as sands) behave as open hydraulic systems where compression induced, for example, by the load of a foundation, causes simultaneous drainage of the fluid from the pores. In practice, the fluid does not take part in the mechanical response and the stress induced weighs only on the solid skeleton that, in turn, subsides in association with reduced porosity. On the contrary, soils with very low permeability (such as clays) exhibit hydraulic delay in reacting to stresses, with consequent development of an initial interstitial overpressure $u_0 \neq 0$ that contradicts Equation (1), and participate in the mechanical response with the solid skeleton. A transient filtering motion fol-
laws that comes to end only when the initial value of the interstitial pressure is reset \( (u_e = 0) \).

In practice, given that a deformation of the solid skeleton occurs together with the expulsion of water, sands develop elasto-plastic settlements synchronous with load application while clays exhibit time-dependent consolidation settlements typically characterized as reverse hyperbolic functions.

Considering the above, the one-dimensional consolidation equation [1,2] describes the hydraulic behaviour of soils in transient conditions by making it possible to simulate the variation in time of interstitial overpressures \( (u_e) \), generated—for example—by the load induced by a foundation or by a road embankment (Figure 1), with consequent visco-elastic settlements to which corresponds a structural reorganisation of the solid skeleton, with reduction of porosity and, concurrently, of the degrees of freedom.

This formulation can be inferred from applying the continuity equation to supposedly saturated soils leading [3], with a few mathematical manipulations, to the following relationship that demonstrates how the transient filtering motion depends on the vertical permeability factor \( (K_v \text{ or } K_z) \), on the compressibility factor \( (m_v) \) and on the weight of the volume of water \( (\gamma_w = 10 \text{ kN/m}^3) \) that in turn identifies the fluid:

\[
2 \left( \frac{2}{e} \right) \sin \left( \frac{2}{e} \right) = \frac{\partial^2 u_e}{\partial z^2} + \left( \frac{2}{e} \right) \frac{\partial K_v \frac{\partial u_e}{\partial z}}{\partial z} = \frac{\partial u_e}{\partial t}. \tag{2}
\]

The next step consists in introducing the hypothesis (in contrast with experimental results) that the permeability factor does not change during consolidation:

\[
2 \left( \frac{2}{e} \right) \sin \left( \frac{2}{e} \right) = \frac{\partial^2 u_e}{\partial z^2} = \frac{\partial u_e}{\partial t}. \tag{3}
\]

Finally, denoting the consolidation coefficient as \( c_v \) or \( c_z \):

\[
c_v = \frac{K_v}{m_v \gamma_w}, \tag{4}
\]

you come to write the classical one-dimensional consolidation equation:

\[
c_v \frac{\partial^2 u_e}{\partial z^2} = \frac{\partial u_e}{\partial t}. \tag{5}
\]

It should be noticed that Equation (5) is analogous to Fourier’s law on heat propagation to the point that you can define the theory of consolidation as the simulation of the propagation of stress-induced interstitial pressures in the subsoil.

2. Exact Solution of Terzaghi’s Consolidation Equation

2.1. Assumptions

Let’s assume that \( u_e, k_z > 0 \) are two positive constants assigned and that:

\[
u_e : [0, +\infty) \times [0, +\infty) \rightarrow \mathbb{R} \tag{6}
\]

is the regular function given by:

\[
u_e (z,t) = u_e e^{k_z z} \cos \left( 2c_k k_z^2 t - k_z z \right). \tag{7}
\]

Now, you can easily notice that \( u_e (z,t) \) solves the differential Equation (5); indeed, given the validity of the following:

\[
\frac{\partial u_e}{\partial t} = -2c_k k_z^2 u_e e^{k_z z} \sin \left( 2c_k k_z^2 t - k_z z \right), \tag{8}
\]

\[
\frac{\partial u_e}{\partial z} = -k_z u_e e^{k_z z} \cos \left( 2c_k k_z^2 t - k_z z \right)
+ k_z u_e e^{k_z z} \sin \left( 2c_k k_z^2 t - k_z z \right), \tag{9}
\]

\[
\frac{\partial^2 u_e}{\partial z^2} = -2k_z^2 u_e e^{k_z z} \sin \left( 2c_k k_z^2 t - k_z z \right), \tag{10}
\]

you obtain:

\[
c_v \frac{\partial^2 u_e}{\partial z^2} = -2c_k k_z^2 u_e e^{k_z z} \sin \left( 2c_k k_z^2 t - k_z z \right) \frac{\partial u_e}{\partial t}, \tag{11}
\]

which is the differential equation given.

If you analyse function (7), you will find out that it simulates the time variation of interstitial overpressures in the subsoil through the consolidation constant \( k_z \)—starting from the point where they are triggered—by dampening their width as depth increases through a reverse hyperbolic function.

2.2. Connection with Experimental Data

Now, since function (7) is a solution of Equation (5), it should be necessarily extended also to experimental methods—oedometrically considered—in order to de-
termine the parameters that govern it correctly. In this sense, it may be useful to analyse an important property of the function $u_e$.

Given $H > 0$ (Figure 2), it can be useful to identify the instant $t_H > 0$ to which the following conditions apply:

$$u_e(H, t_H) = 0, u_e(H, t) \neq 0, \text{if } 0 < t < t_H.$$  \hspace{1cm} (12)

As a first passage, you should notice that equation $u_e(H, t) = 0$ reduces to the form:

$$\cos\left(2c_l k_z^2 t - k_x H\right) = 0$$  \hspace{1cm} (13)

which has infinite solutions like:

$$2c_l k_z^2 t - k_x H = \frac{\pi}{2} + h\pi \text{ upon var. of } h \in \mathbb{Z}$$  \hspace{1cm} (14)

or like:

$$t = \frac{2k_x H + \pi}{4c_l k_z^2} + \frac{\pi}{2c_l k_z^2} h \text{ upon var. of } h \in \mathbb{Z}$$  \hspace{1cm} (15)

The next passage consists in selecting the positive value $t$ closest to zero from among those determined by setting the condition $t > 0$ that provides:

$$\frac{2k_x H + \pi}{4c_l k_z^2} + \frac{\pi}{2c_l k_z^2} h > 0$$  \hspace{1cm} (16)

from which we obtain:

$$h > -\frac{2k_x H + \pi}{2\pi}.$$  \hspace{1cm} (17)

The last passage includes that, if you set also the following condition:

$$h_H = \min\left\{h \in \mathbb{Z} : h > -\frac{2k_x H + \pi}{2\pi}\right\},$$  \hspace{1cm} (18)

you obtain the formulation of the consolidation completion time:

$$t_H = \frac{2k_x H + \pi + 2h_H \pi}{4c_l k_z^2}.$$  \hspace{1cm} (19)

Finally, from Equation (19) you can extract the consolidation coefficient:

$$c_v = \frac{2k_x H + \pi + 2h_H \pi}{4t_H k_z^2}.$$  \hspace{1cm} (20)

$H$ is from Figure 2, $h_H = H_{100}$ and $t_H = t_0 t$ from Figure 3 while $k_x$ depends on the boundary conditions (uniqueness theorem).

3. Extension to the Two/Three-Dimensional Cases

Let’s put $u_e, c_x, c_z, k_x, k_z > 0$ to be positive constants assigned and be:

$$u_e : [0, +\infty) \times [0, +\infty) \rightarrow \mathbb{R}$$  \hspace{1cm} (21)

the regular function given by:

$$u_e(x, z, t) = u_e e^{-c_x x - c_z z} \cos\left[2\left(c_x k_x^2 + c_z k_z^2\right) t - k_x x - k_z z\right].$$  \hspace{1cm} (22)

Going through the same passages as in Equations (8) to (11), you can easily notice that $u(x, z, t)$ solves the differential equation:

$$c_x \frac{\partial^2 u}{\partial x^2} + c_z \frac{\partial^2 u}{\partial z^2} = \frac{\partial u}{\partial t}.$$  \hspace{1cm} (23)

Similarly, if $u_e, c_x, c_z, k_x, k_z > 0$ are positive
constants assigned and:
\[ u_e : \mathbb{R}^3 \times \mathbb{R}^+ \rightarrow \mathbb{R} \]  
(24)

the regular function given, then:
\[ u_e (x, y, z, t) = u_e e^{-k_x x - k_y y - k_z z} \]
\[ \times \cos \left[ 2 \left( c_1 k_x + c_2 k_y + c_3 k_z \right) t - k_x x - k_y y - k_z z \right] \]
(25)
solves the differential equation:
\[ c_1 \frac{\partial^2 u_e}{\partial x^2} + c_2 \frac{\partial^2 u_e}{\partial y^2} + c_3 \frac{\partial^2 u_e}{\partial z^2} = \frac{\partial u_e}{\partial t}. \]
(26)

4. Conclusions
Notwithstanding that the theory of monodimensional consolidation was first presented in 1936, it is still today taught in all geotechnical engineering courses (for example in Italy: The University of Naples Federico II [8]; The Polytechnic University of Milan [9]; The Polytechnic University of Turin [10]) and is also the only theory used in the professional procedures of engineers and geologists (to predict viscoelastic settlement of ground subjected to loads) and is likewise the only instrument applied in geotechnical laboratories to derive experimental data using oedometric tests.

The only valid alternative is provided by Biot’s theory [11] which is derived from the union of the equation of continuity (3), extended into three dimensions, with Navier’s Equations to produce a system of 4 equations with 4 unknown variables relating to the interstitial pressure and movement along three directions. However owing to its complexity it is only applicable in simple cases, for which a precise solution exists [12-21] or when the solution is arrived at using numeric methods—for example—the finite elements method implemented in modern professional and research software [22,23].

With in mind the entire Mechanics of Soils, and the study of the soil visco-elastic behaviour [1,2] in particular, the application of Terzaghi’s differential equation is historically based on Taylor’s solution [24] that approximates experimental results—limited to the one-dimensional case only—through the introduction of arbitrary and fixed non-dimensional variables, independent of the geological history of the means. This research work makes a proposal for an exact solution that can be considered correct as it solves the differential equation and, at the same time, allow correct interpretation of experimental data; then, solution has been fruitfully extended to the two- and three-dimensional cases.

To conclude, results even satisfactory have come to light from the analysis of data. At the same time, they have opened additional research channels, considering that the uniqueness theorem will be proved later on the basin of the oedometric boundary conditions.

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