

# Dimethylsulfide and Coral Bleaching: Links to Solar Radiation, Low Level Cloud and the Regulation of Seawater Temperatures and Climate in the Great Barrier Reef

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**How to cite this paper:** Jones, G., Curran, M., Swan, H. and Deschaseaux, E. (2017) Dimethylsulfide and Coral Bleaching: Links to Solar Radiation, Low Level Cloud and the Regulation of Seawater Temperatures and Climate in the Great Barrier Reef. *American Journal of Climate Change*, 6, 328-359.

<https://doi.org/10.4236/ajcc.2017.62017>

**Received:** December 14, 2016

**Accepted:** June 17, 2017

**Published:** June 20, 2017

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## Abstract

Coral reefs produce atmospheric dimethylsulfide (DMS<sub>a</sub>) which oxidises to non-sea-salt (nss) sulfate aerosols, precursors of cloud condensation nuclei (CCN) and low level cloud (LLC), reducing solar radiation and regulating sea surface temperatures (SSTs). Here we report measurements of solar radiation, SST, LLC, DMS flux, NH<sub>4</sub><sup>+</sup>, and rainfall before, during and after a major coral bleaching event at Magnetic Island in the central Great Barrier Reef (GBR). Measurements are compared with those made at the nearby fringing reef of Orpheus Island where coral bleaching did not occur. Extreme solar radiation levels occurred from November to late January and could have reflected cloud radiative effects that increased downwelling of solar radiation. High levels of LLC often coincided with high periodic fluxes of DMS from the unbleached coral reef at Orpheus Island (e.g. 14 - 20 μmol·m<sup>-2</sup>·d<sup>-1</sup>), in direct contrast to the very low fluxes of DMS that were emitted from the bleached, human-impacted Magnetic Island fringing reef (nd-0.8 μmol·m<sup>-2</sup>·d<sup>-1</sup>) when SSTs were >30°C. Continuous SSTs measurements at the Magnetic Island reef revealed various heating and cooling periods, interspersed with stable SSTs. Cooling periods (negative climate feedback) ranged from -1°C to -3°C (7 day mean -1.6°C), and often seemed to occur during low tides, periodic pulses of DMS flux and LLC, keeping SSTs < 30°C. In contrast warming periods of +1°C to +3°C (positive climate feedback, 7 day mean +1.52°C), seemed to occur during increasing tides, decreasing DMS flux and low to medium levels of LLC which increased solar radiation and caused SSTs over 30°C and corals to bleach. Alternation between these two states or types of feedback is indicated in this research and may be a function of enhanced scattering of solar

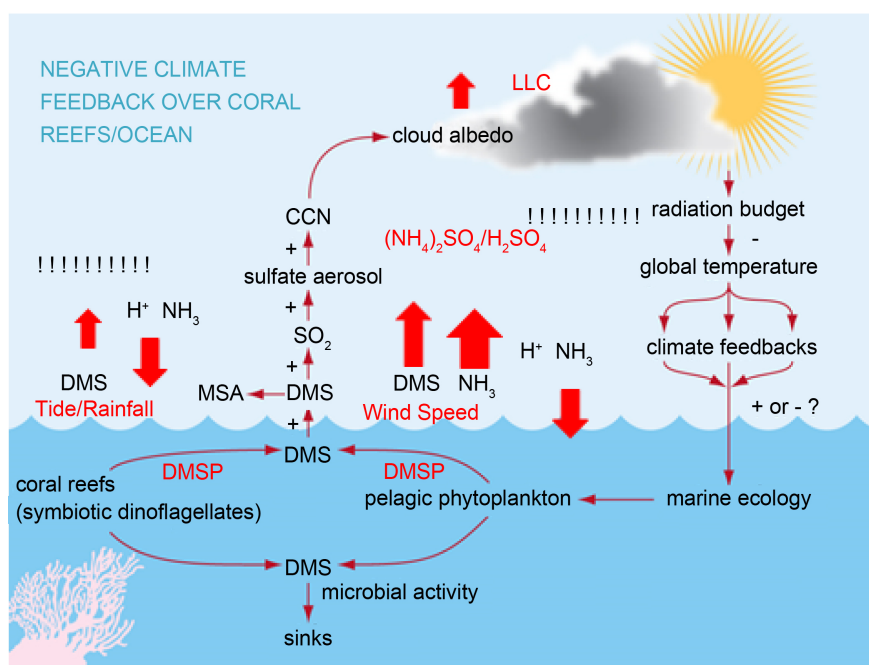
radiation from nss-sulfate aerosols that originate from oxidation of DMS<sub>a</sub> produced from the coral reefs in the GBR. The net radiative forcing from clouds can be as high as four times as large as the radiative forcing from a doubling of CO<sub>2</sub> levels in the atmosphere, which needs to be taken into account when ascribing coral bleaching events in the GBR solely to GHG warming. Further studies are needed to more critically assess the importance of this GBR coral reef-cloud feedback to the climate of northern Australia and the western Pacific, where the greatest biomass of coral reefs occurs.

## Keywords

Dimethylsulfide, Low Level Cloud, Climate Feedback, Coral Bleaching

## 1. Introduction

Reef-building corals contain high concentrations of dimethylsulfoniopropionate (DMSP) [1]-[6] that can be enzymatically converted to dimethylsulfide (DMS), a trace sulfur gas that oxidizes in the atmosphere to form nss-sulfate aerosol particles and cloud condensation nuclei (CCN) (Figure 1). The growth



**Figure 1.** The climate feedback involving production of DMS that can occur from coral reefs and adjacent ocean waters of the Great Barrier Reef. Oceanic phytoplankton and coral reefs produce seawater DMS which is transferred by air-sea transfer to the atmosphere where it oxidises to non-sea sulfate aerosol and methane sulfonate (MSA). The tiny sulfate aerosol particles grow to produce cloud condensation nuclei (CCN) forming low level cumulous and stratocumulous clouds, thus decreasing solar radiation levels, SSTs, and increasing rainfall. [Adapted from Charlson *et al.* 1987 [7] and Jones and Gabric, 2006 [8]]. Low tides and convective rainfall can increase the flux of DMS to the reef atmosphere, whilst wind speed over the ocean affects DMS transfer to the ocean atmosphere [9] [10].

of CCN can lead to the formation of low level cloud (LLC) such as cumulus and stratocumulus clouds, decreasing solar radiation and sea surface temperatures (SST), thus influencing climate as a negative climate feedback [7] [11]. Atmospheric DMS can be oxidized to nss-sulfate aerosol by hydroxyl (OH) and nitrate ( $\text{NO}_3$ ) radicals [12], although  $\text{O}_3$ ,  $\text{HO}_2$ , and  $\text{H}_2\text{O}_2$  can also be involved [13]. The products of this oxidation produce  $\text{SO}_2$ , methanesulfonate acid (MSA) and gas-phase sulfuric acid ( $\text{H}_2\text{SO}_4$ ) which can then condense onto aerosol particles [14], or nucleate to form new sulfuric acid particles [15] (Figure 1). The climate impact of atmospheric DMS ( $\text{DMS}_a$ ) depends on the ability of its oxidation products to increase CCN, cloud droplet concentration [13], LLC and rainfall [16] [17].

The CLAW hypothesis (named after Charlson, Lovelock, Andrea and Warren) [7] states that because the albedo of low level clouds (Figure 1), and thus the Earth's radiation budget is sensitive to CCN number; biological regulation of the climate is possible through the effects of sea surface temperature (SST) and sunlight on phytoplankton population and DMS production [7]. This climate regulation is also thought to occur over coral reefs [18] [19] [20]. Variations in solar radiation can increase or decrease phytoplankton activity which in turn can increase or decrease seawater and atmospheric DMS [21]. This may increase or decrease the albedo of LLC cover in the marine boundary layer (0 - 1000 m), which will decrease or increase SSTs, thus regulating climate [7]. Charlson *et al.* [7] hypothesized that as temperatures increase from GHG warming, phytoplankton DMS production will increase, thus increasing CCN and cloud albedo, decreasing solar radiation and lowering SSTs, keeping global temperatures relatively constant; a negative climate feedback (Figure 1). Such a cloud cover-SST feedback occurs in the NE Pacific [22], and over coral reefs in the GBR [23], but it is not known whether DMS is implicated in these feedbacks. An increasing amount of evidence now suggests that a LLC climate feedback involving enhanced production of  $\text{DMS}_a$  and other aerosols from coral reefs could occur in the GBR [18] [20], and indeed over reefs in the western Pacific warm pool (WPWP) [24] [25]. For example, air masses that travel over coral reefs in the southern GBR contain ammonium sulfate aerosols (Figure 1) derived mainly from the oxidation of  $\text{DMS}_a$  which are likely produced from the GBRs coral reefs [26]. Modini *et al.* [26] have characterised these freshly formed aerosol particles that consist of 60% ammonium sulfate and 40% organic material. These measurements are the first direct observation of aerosol particle formation over the GBR, and compare with earlier observations [27]. Mean aerosol particle concentrations averaged  $3200 \text{ cm}^{-3}$  and consisted of broad pulses of aerosol particles emitted in the morning and afternoon, with a small discrete aerosol burst at mid-day [26], a distribution similar to that measured for  $\text{DMS}_a$  over coral reefs in the central GBR [28]. Periodic pulses of  $\text{DMS}_a$  often occur over coral reefs during very low tides and rising tides [10] [28] [29] [30]. It is apparent that natural stress events (e.g. low tides, high SSTs, rainfall) cause the coral to utilize intracellular DMSP to cope with the resultant osmotic, solar radiation and tem-

perature stresses [31] [32] [33], and in the process produce enhanced levels of seawater and atmospheric DMS [10] [28] [30] [34].

These periodic pulses of  $\text{DMS}_a$  at low tides may be capable of increasing LLC cover over coral reefs (Figure 1), keeping SSTs  $< 30^\circ\text{C}$ , thus decreasing the incidence of coral bleaching and significantly affecting regional climate in the GBR, western and central Pacific Ocean [20] [23] [24] [28] [35]. However, the increased frequency of coral bleaching in the GBR in the last two decades could suggest that the strength of this feedback is decreasing [36] [37] [38] [39] [40].

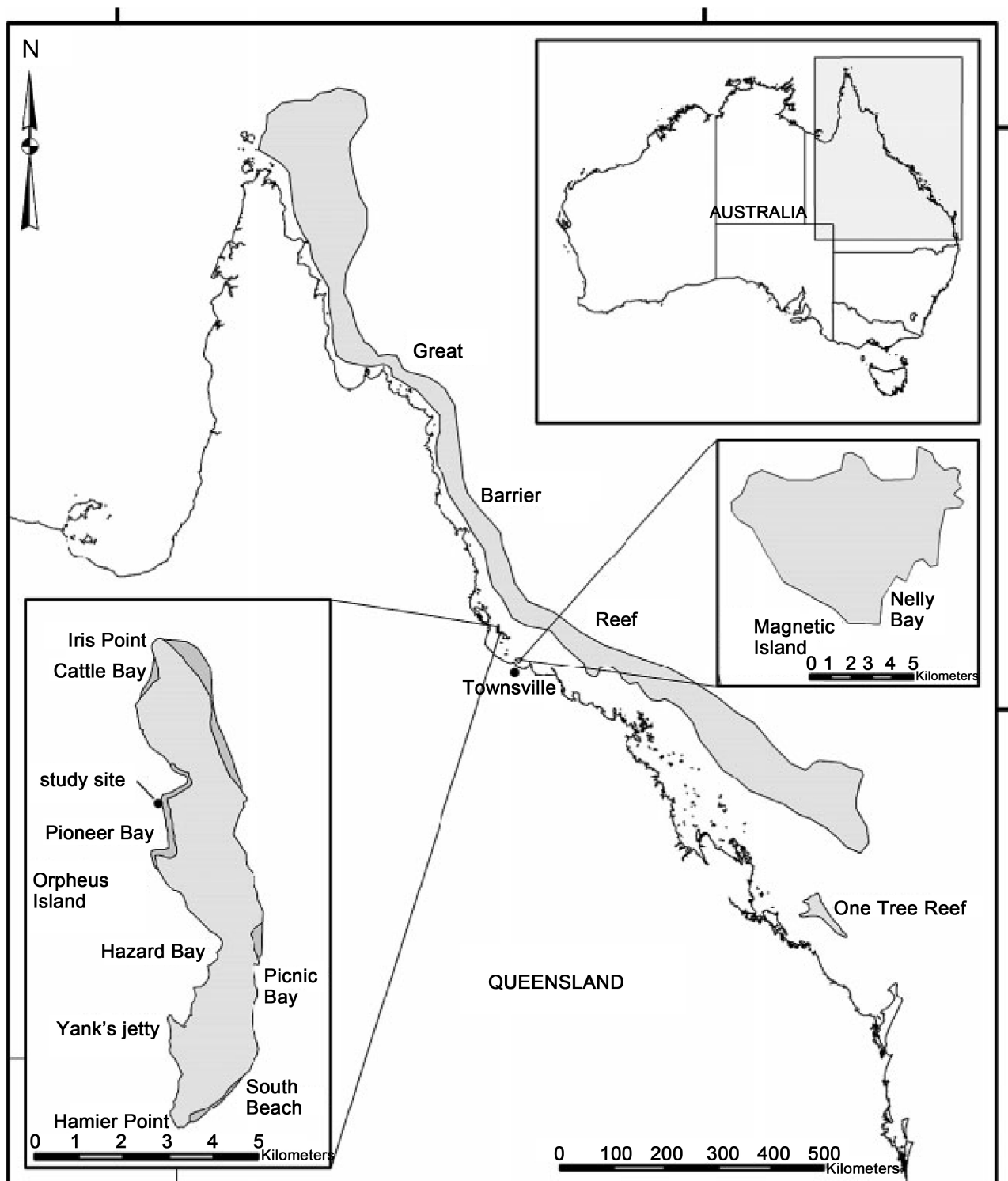
This article discusses the potential of a “coral reef-DMS-cloud” feedback involving reef production of  $\text{DMS}_a$  to constrain SSTs in the GBR, and could be the reason why coral reefs in the WPWP have not bleached so frequently over a period of 25 years [24]. In early January 1994 very high seawater temperatures ( $>33^\circ\text{C}$ ), coupled with high air temperatures ( $44.3^\circ\text{C}$ ), associated with hot westerly winds, occurred at Nelly Bay reef, Magnetic Island off Townsville in the central GBR. Many colonies of corals at this location bleached to a pale colour and colonies of some species (mainly *Acropora* and *Pocillopora*) bleached to a bone-white colour [31] [41]. In contrast, no coral bleaching was recorded at Pioneer Bay, Orpheus Island 70 km to the north of Townsville, reflecting the lower SSTs at this study site [28]. Prior, during and after the severe bleaching event at Nelly Bay seawater DMS and DMS fluxes were measured at both Pioneer Bay and Nelly Bay fringing coral reefs. The DMS database gathered during these bleaching events are unique in that it is the only DMS database we have before, during and after a coral bleaching event in the GBR. Given that DMS could have such an important role in the development of LLC cover over coral reefs, thus regulating SSTs, the aims of this paper are: 1) To highlight the changes in dissolved or seawater  $\text{DMS}_w$ , dissolved  $\text{NH}_4^+$  and DMS flux, solar radiation and SST before, during and after the bleaching event. 2) To examine the changes in LLC cover with low tides and rainfall before, during and after the bleaching event. 3) To examine any correlations or links between LLC cover, solar radiation, SST, rainfall and the sulfur substances measured in order to throw more light on this possible coral reef climate feedback in the GBR. Whilst further studies are necessary on the physics and chemistry of sulfate aerosol, CCN and LLC formation over coral reefs in the GBR, the study described here provides more momentum for such studies to take place and may provide more insights into the frequency and intensity of this coral reef-climate feedback, as well as its sustainability in the face of GHG warming and increased coral bleaching.

## 2. Study Areas and Coral Bleaching

### 2.1. Nelly Bay-Magnetic Island

Nelly Bay is one of the larger embayments on the eastern coast of Magnetic Island and approximately 7 km from Townsville and faces SE into Cleveland Bay (Figure 2). The approximate size of the reef at Nelly Bay is 43 hectares. Within Nelly Bay the reef rises sharply from the general Cleveland Bay floor to form a wide





**Figure 2.** Map of the study area showing locations of Nelly Bay fringing reef on the NE coast of Magnetic Island and the Pioneer Bay reef site at Orpheus Island north of Townsville.

reef flat that dries at lowest low tides (**Figure 2**). The reef slope area contains areas of abundant coral cover mixed with other areas dominated by brown algae. The study site ( $146^{\circ}51'N$ ,  $19^{\circ}10.5'E$ ) was approximately 700 m directly offshore. This site was chosen since it could be easily accessed from the shore, and con-

tained many species of corals that contained high concentrations of DMSP, and so produced high concentrations of  $\text{DMS}_w$  compared with phytoplankton production [3]. Human impacts such as increased suspended sediments from dredging and enhanced nutrient levels from sewage discharge into Cleveland Bay affect this site [42] [43].

## 2.2. Pioneer Bay-Orpheus Island

Orpheus Island (lat.  $18^{\circ}35'S$ , long.  $146^{\circ}29'E$ ) is ~70 km NE of Townsville on the east coast of Queensland (**Figure 2**). A major seasonal study was conducted at this site over 19 months, with 18 regular visits to the James Cook University marine research station at Pioneer Bay Reef between 7 December 1992 and 25 February 1994, and one further visit on 19 July 1994 [28]. The DMS results of this study are summarised here in order to make comparisons with Nelly Bay reef, and because the two reefs together give a more accurate idea of the flux of DMS from coral reefs in the local area during the study period. The sampling location for the present study was at the edge of the fringing coral reef in the northern part of Pioneer Bay. The fringing reef flat is ~400 m wide and the sampling site was specifically chosen because both the flooding and ebbing tide travel parallel with the coast over a significant amount of reef before reaching the sampling point [44]. At this site, there was enough water above the reef that even at the lowest of tides, the reef was just covered, enabling water sampling at all tidal heights. The reef flat and slope at Pioneer Bay contain a diverse assemblage of corals, although the coral community is likely impoverished because of human impacts from the marine station.

## 2.3. Coral Bleaching

The 1994 coral bleaching event at Nelly Bay, Magnetic Island was first observed by the Great Barrier Reef Marine Park Authority (GBRMPA) on 16<sup>th</sup> January [44]. Underwater observations of the 1994 bleaching event at Nelly Bay from 28<sup>th</sup> January to 14<sup>th</sup> February have been reported [44] [45]. Jones (1995) [44] observed that the most pronounced bleaching occurred at this site in the arborescent and tabulate acroporids, the montiporids and the pocilloporids from 8<sup>th</sup> January. In some cases heavily bleached colonies were observed immediately adjacent to normal coloured colonies of the same species. In the heavily bleached branching coral species, there were usually slight tinges of yellow/brown colour on the undersides of the branches, especially those in the interior of the colonies. Preferential loss of colour on the upward facing surfaces of the corals represented a frequent bleaching response and indicated an effect from increased levels of solar radiation [46]. The first signs of recovery from the bleaching event were noted on a dive conducted on 24<sup>th</sup> February 1994 [44]. Mottled or paler patches on partially bleached colonies appeared slightly browner and colonies readily regained their colouration in the 3 - 4 months after the bleaching event. By May 1994 few discoloured colonies could be observed at Nelly Bay, although pale yellow colonies of *Pocillopora damicornis* were observed as late in the year as July 1994 [31] [45]. Many colonies of *P. damicornis* lost in excess of 90% of

their zooxanthellae, and bleached to a bone white colour. Six months after the bleaching event many colonies of *P. damicornis* were still bleached with zooxanthellae densities up to 10 times lower than unbleached colonies ( $1.22 \times 10^5$  compared to  $2.67 \times 10^6 \text{ cm}^{-2}$ ). Changes in intracellular DMSP concentrations in this coral have been recently reported [31]. In contrast to Nelly Bay no bleaching of corals at Pioneer Bay was observed during our study period.

### 3. Methods

#### 3.1. Sea and Air Temperatures, Solar Radiation, Cloud Cover, Rainfall and Wind Speed

Continuous sea surface temperatures (SST) were measured at Geoffrey Bay (next bay to Nelly Bay), Magnetic Island by staff from AIMS which enabled heating and cooling periods to be expressed as degree heating week (DHW) or degree cooling weeks (DCW) over the November to March period. Average daily air temperatures, cloud cover and rainfall were recorded by the Townsville Bureau of Meteorology (TBM). Daily global solar radiation and low level cloud (LLC) cover was measured at Townsville airport, about 5 km from Nelly Bay by the Bureau of Meteorology (BoM) and downloaded from the BoM website. LLC was the amount of the sky covered by cloud present in the marine boundary layer (MBL = upper 500 - 1000 m) and is an estimate of marine cumulus and strato-cumulus cloud which  $\text{DMS}_a$  is reputed to form [7]. LLC measurements were made every 3 hours and averaged to give daily mean values. The daily global solar exposure is the total solar energy for a day falling on a horizontal surface. It is measured from midnight to midnight. The BoM's computer radiation model uses visible images from geostationary meteorological satellites to estimate daily global solar exposures at ground level. At each location the image brightness is used to provide an estimate of the solar irradiance at the ground or sea surface. Essentially, the irradiance at the surface can be calculated from the irradiance at the top of the earth's atmosphere, the amount absorbed in the atmosphere (dependent on the amount of water vapour present), the amount reflected from the surface (surface albedo) and the amount reflected from clouds (cloud albedo). These instantaneous irradiance values are integrated over the day to give daily solar exposure in  $\text{MJ}\cdot\text{m}^{-2}$  and are clearly a reflection of cloud cover, aerosol concentrations and water vapour in the atmosphere. Both LLC and solar radiation measurements whilst measured at Townsville airport are assumed to reflect conditions in the local geographical area.

#### 3.2. $\text{DMS}_w$ , $\text{NH}_4^+$ and DMS Flux Measurements

Water samples from Nelly Bay, Magnetic Island were collected via a small boat about 700 m offshore and over the coral reef. The site was marked by a small buoy [43]. Seawater was gently filtered using a 50 mL silanized glass syringe, with Teflon plunger, through a pre-rinsed  $0.45 \mu\text{m}$  cellulose acetate filter (Minisart, Sartorius). Water samples from Nelly Bay and Pioneer Bay were collected and analysed for  $\text{DMS}_w$  according to procedures already described [28] [46]. The

flux of DMS from Nelly Bay and Pioneer Bay fringing reefs was calculated using the method of Liss and Merlivat (1986) [9] and calculated as reported in Curran and Jones (2000) [47]. More continuous measurements of  $\text{DMS}_w$  and DMS flux at Nelly Bay reef were obtained by using the regression equations between  $\text{DMS}_w$  and SSTs ( $\text{DMS}_w = 0.5539 \text{ SST} - 12.074$ ;  $r^2 = 0.44$ ,  $p < 0.001$ ), and  $\text{DMS}_w$  and DMS flux ( $\text{DMS flux} = 2.247 \text{ DMS}_w - 1.6945$ ;  $r^2 = 0.67$ ,  $p < 0.01$ ), enabling comparison with actual measurements. Calculated  $\text{DMS}_w$  and DMS flux values compared very favorably with values obtained from seawater DMS measurements and application of the transfer velocity equation [9] (in prepn.). Samples collected from Nelly Bay were analyzed for  $\text{NH}_4^+$  by AIMS [43].

### 3.3. Zooxanthellae Densities in *Acropora formosa* Coral

Zooxanthellae densities in *A. formosa* have been reported [44]. They are replotted here as a proxy of the stresses that affected corals at Nelly Bay and provide some insights on the influence of *A. formosa* on DMS in seawater given that coral zooxanthellae produce large amounts of DMS [20]. Zooxanthellae densities in *Pocillopera damicornis*, which was still bleached at Nelly Bay in June 1994, have recently been reported, as well as intracellular levels of DMSP [3] [31].

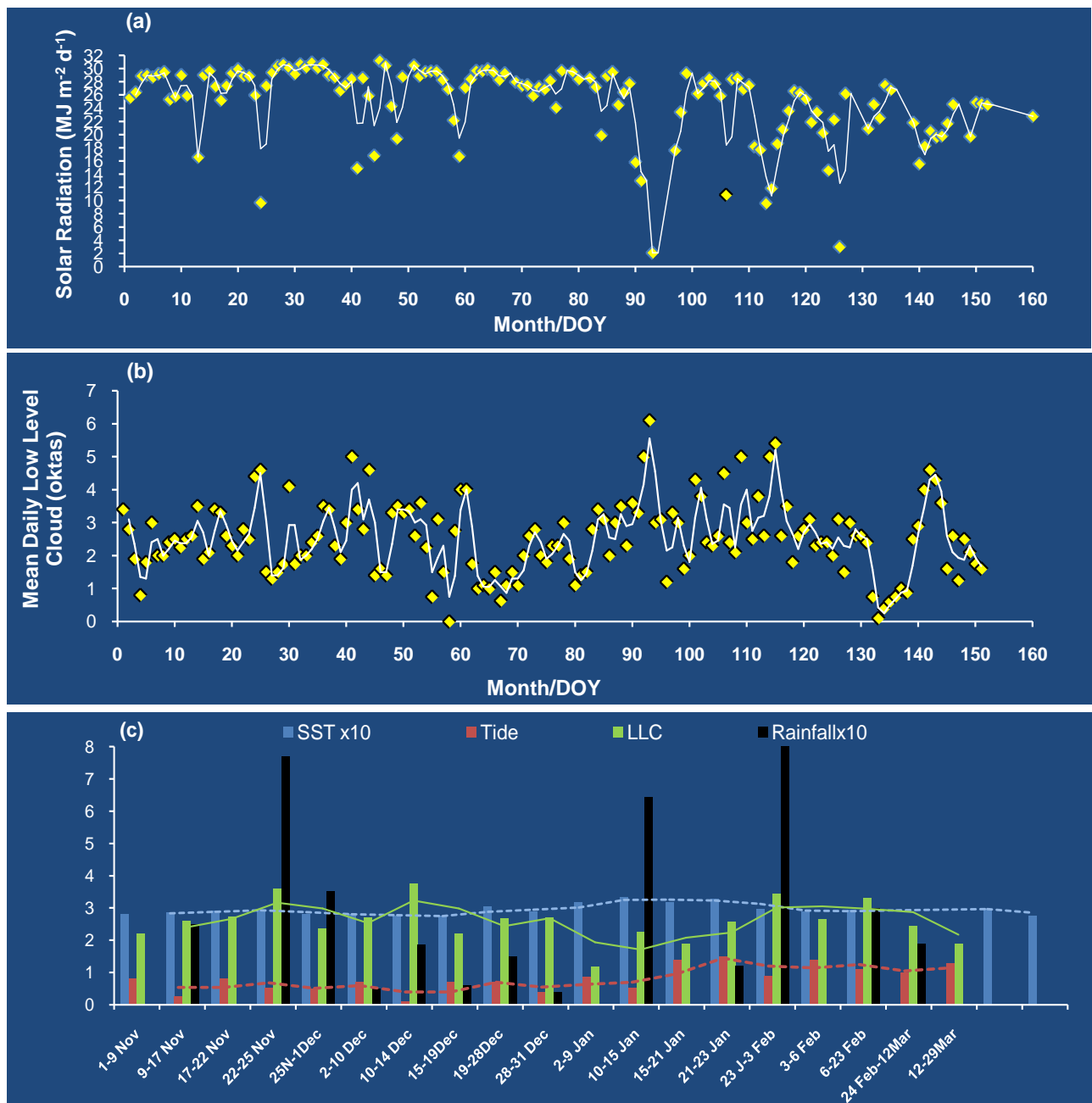
### 3.4. Statistical Analysis

SPSS 11 was used for the statistical analysis of the data sets. For the seawater data, statistically significant differences between treatments were tested using a Pearson's Product Linear Correlation analysis (parametric, normally distributed) with significance determined at the 0.01 to 0.001 level.

## 4. Results

### 4.1. Solar Radiation and Air Temperatures

Periods of very high solar radiation levels occurred over 29 days ( $>29 \text{ MJ}\cdot\text{m}^{-2}$ ) in November (DOY 1-31), December (DOY 31-61) and early January (DOY 62-72), with 9 consecutive days of extreme solar radiation levels from 27<sup>th</sup> November to 5<sup>th</sup> December (DOY 27-36) when values ranged from 29.2 to 30.9  $\text{MJ}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  (mean 30.3) (Figure 3(a)). From 1<sup>st</sup> to 6<sup>th</sup> January (DOY 62-71) prior to the bleaching event on 8<sup>th</sup> January (DOY 69), extreme solar radiation levels ranged from 28.3 to 29.9  $\text{MJ}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  (mean 29.4). The mean solar radiation levels over Nelly Bay reef in January (mean = 26.7  $\text{MJ}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ), with the early January and late November to early December values, were higher than the mean January levels for the 1990-2012 period (24.7  $\text{MJ}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ), and higher than the highest monthly mean for all 22 years (28.4  $\text{MJ}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) (BoM online data). From 4-9<sup>th</sup> January (DOY 65-70), daily air temperatures averaged over 30°C, and included a series of exceptionally high maximum daytime temperatures in excess of 36°C recorded by the Townsville Bureau of Meteorology (TBM). On 7<sup>th</sup> January 1994, the maximum air temperature was 44.3°C, the highest air temperature recorded at the time by the TBM since 1942 [44]. From 19<sup>th</sup>-23<sup>rd</sup> January



**Figure 3.** (a) Solar radiation levels in MJ·m<sup>-2</sup>·d<sup>-1</sup>, (b) Mean daily low level cloud cover (oktas), versus day of year (DOY), (c) SST (°C), lowest nighttime tides (m), LLC (oktas) and rainfall (mm) recorded from November to April by the BoM at Townsville airport and Nelly Bay.

(DOY 80-84), daily air temperatures also averaged over 30°C, and included a series of unusually high minimum daily temperatures, including the second highest minimum air temperature (29°C) recorded locally since 1942. Daily air temperatures then fell sharply from 29.3°C on 24<sup>th</sup> January, to 24.8°C on 31<sup>st</sup> January 1994 [44].

#### 4.2. Low Level Clouds, SST and Rainfall

Regular pulses of LLC cover occurred in November and December that lasted for

about a week (**Figure 3(b)**). From November-December mean daily LLC generally averaged 2 - 3 oktas with pulses often coinciding with low tides and rainfall (**Figure 3(c)**). These pulses of LLC coincided with relatively stable SSTs of 27.5°C - 29.5°C (mean 28.37°C), except from 19-28<sup>th</sup> December when SSTs reached 30.3°C. From the 2<sup>nd</sup> to 9<sup>th</sup> January (DOY 63-70) mean daily LLC cover markedly decreased (~1 okta) and coincided with the extreme solar radiation levels that caused corals to bleach white on their upper surfaces at Nelly Bay reef [44] with SSTs increasing to 32°C - 34°C which continued to 23<sup>rd</sup> January (**Figure 3(c)**). From 10<sup>th</sup> January to 26<sup>th</sup> February (DOY 71-118) increasing pulses of LLC cover occurred (2 - 6 oktas) as the corals slowly recovered, and then remained low and relatively constant (daily mean ~2 - 3 oktas) up to 9<sup>th</sup> March (DOY 118-129). As LLC increased in February, and low tides increased in height from February-March, SSTs gradually decreased (**Figure 3(c)**). Stable SSTs often coincided with significant rainfall that occurred in November (monthly mean 102 mm) (**Figure 3(c)**), with lower, but more regular, and smaller amounts falling in December (monthly mean 44 mm), which often fell during low tides. On the 10<sup>th</sup> and 12<sup>th</sup> January 54 and 12 mm of rain occurred, with 70 mm falling on 31<sup>st</sup> January, yielding a monthly mean rainfall of 149 mm. In February very high rainfall occurred on 1<sup>st</sup> and 2<sup>nd</sup> February (66, 25 mm), with smaller amounts occurring throughout the month, yielding a monthly mean of 212 mm. From 9-13<sup>th</sup> March (DOY 129-133) mean daily LLC cover rapidly decreased to zero and then rapidly increased to ~ 4.5 oktas over 9 days to 22<sup>nd</sup> March (DOY 142) when the mean daily LLC cover decreased gradually to about 1.5 oktas on 31<sup>st</sup> March (DOY 151) (**Figure 3(b)**). No rainfall occurred in March, and in April only 0.6 mm of rainfall was recorded on 28<sup>th</sup> April.

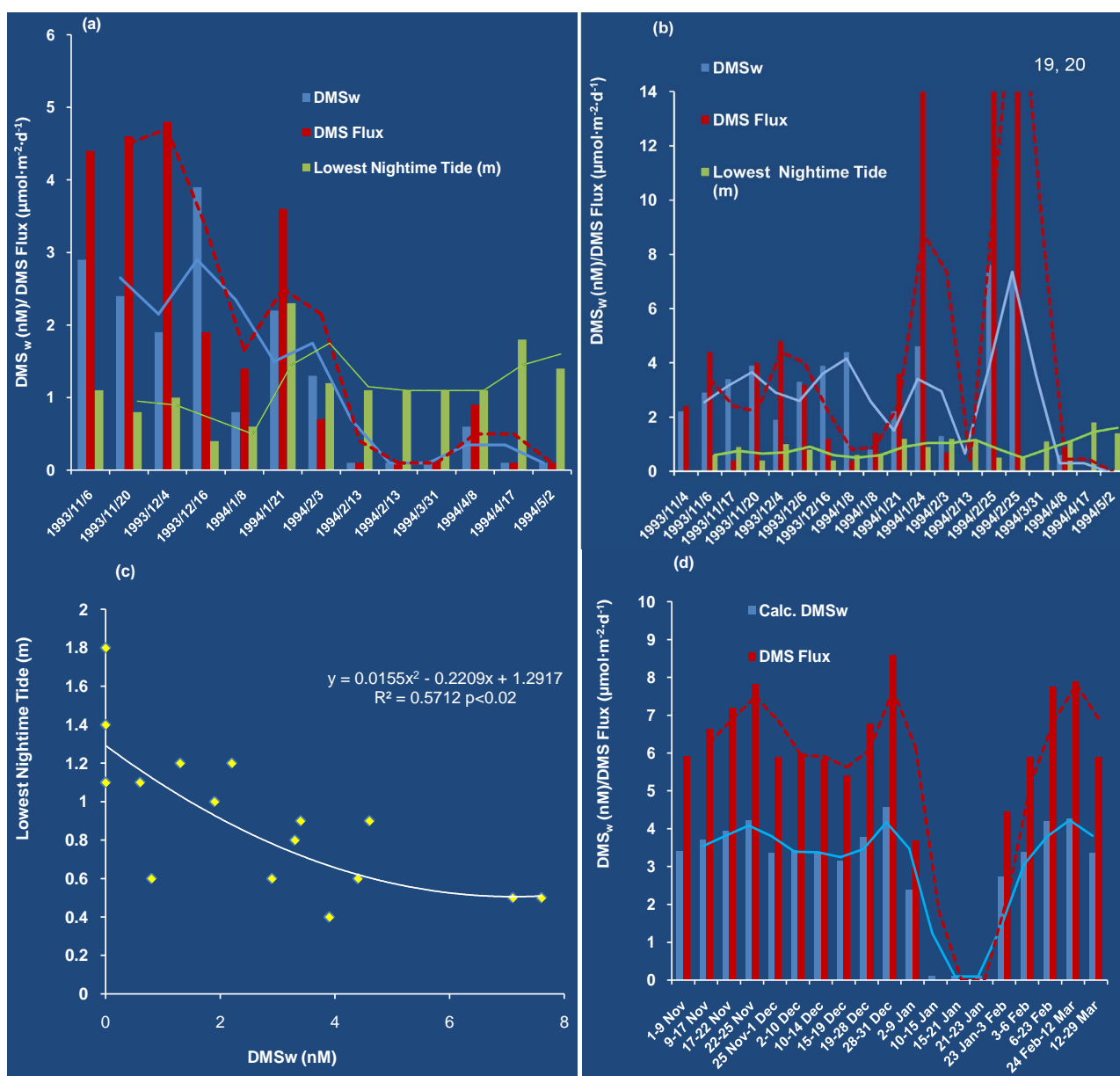
### 4.3. Variation in DMS<sub>w</sub>, DMS Flux, NH<sub>4</sub><sup>+</sup> and Coral Zooxanthellae

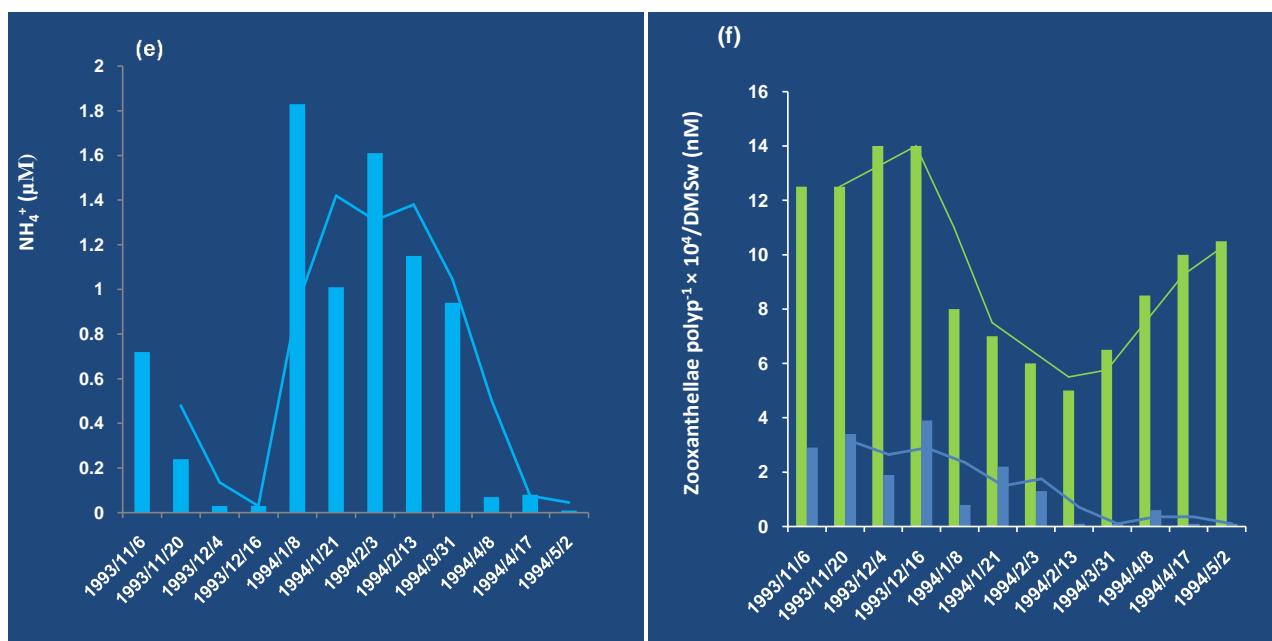
DMS<sub>w</sub> concentrations at Nelly Bay Reef ranged from nd-3.9 nM (mean 2.2 nM) with generally elevated concentrations (>1 µM) occurring from 6<sup>th</sup> November to the 3<sup>rd</sup> February, decreasing to very low concentrations (<1 µM) from 13<sup>th</sup> February to 2<sup>nd</sup> May (**Figure 4(a)**). During mid-February to early May DMS<sub>w</sub> concentrations were particularly low at Nelly Bay (nd-0.6 nM, mean = 0.2 nM) (**Figure 4(a)**). DMS flux concentrations ranged from nd-4.8 µmol·m<sup>-2</sup>·d<sup>-1</sup> (mean 1.64) with high values before the bleaching event (**Figure 4(a)**), and elevated flux levels after the bleaching event on 21<sup>st</sup> January, decreasing to nd-0.8 µmol·m<sup>-2</sup>·d<sup>-1</sup> from 3<sup>rd</sup> February to 2<sup>nd</sup> May. A measure of DMS emitted from coral reefs in the wider local region was obtained by combining the DMS<sub>w</sub> and DMS flux measurements from the fringing coral reef at Nelly Bay fringing reef (Magnetic Island) with those of Pioneer Bay (Orpheus Island) (**Figure 4(b)**). These measurements again highlight increasing pulses of DMS<sub>w</sub> and DMS flux that occurred at these two reefs over the period November to the end of February. For example from 4<sup>th</sup> November-17<sup>th</sup> November DMS flux emitted from these two coral reefs ranged from 0.4 - 4.4 µmol·m<sup>-2</sup>·d<sup>-1</sup> (mean = 2.4). From 8<sup>th</sup> January to 3<sup>rd</sup> February much more DMS was emitted from Pioneer Bay reef than Nelly Bay



reef with fluxes ranging from 0.4 - 14  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  (mean 4.1). On the 24<sup>th</sup> January the huge pulse of DMS<sub>a</sub> emitted from Pioneer Bay reef (14  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) occurred when significant rainfall occurred in the region (10 mm at Townsville airport on 23<sup>rd</sup> January) and low tides occurred at night (0.9 m) (**Figure 3(c)**).

Similarly on 25<sup>th</sup> February an even higher flux of DMS<sub>a</sub> was emitted from the fringing reef at Pioneer Bay ranging from 19 - 20  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  (mean 19.5), when medium low tides occurred (0.8 m) (**Figure 4(b)**). These huge fluxes of DMS emitted from Pioneer Bay reef were in contrast to the very low fluxes of DMS that were emitted from 3<sup>rd</sup> February to 2<sup>nd</sup> May from the Nelly Bay fringing reef (nd-0.8  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) where coral bleaching occurred (**Figure 4(a)**), and calculated fluxes from 10th January to 23rd January (**Figure 4(d)**). As the lowest nighttime tides decreased at both Nelly Bay and Pioneer Bay reefs seawater



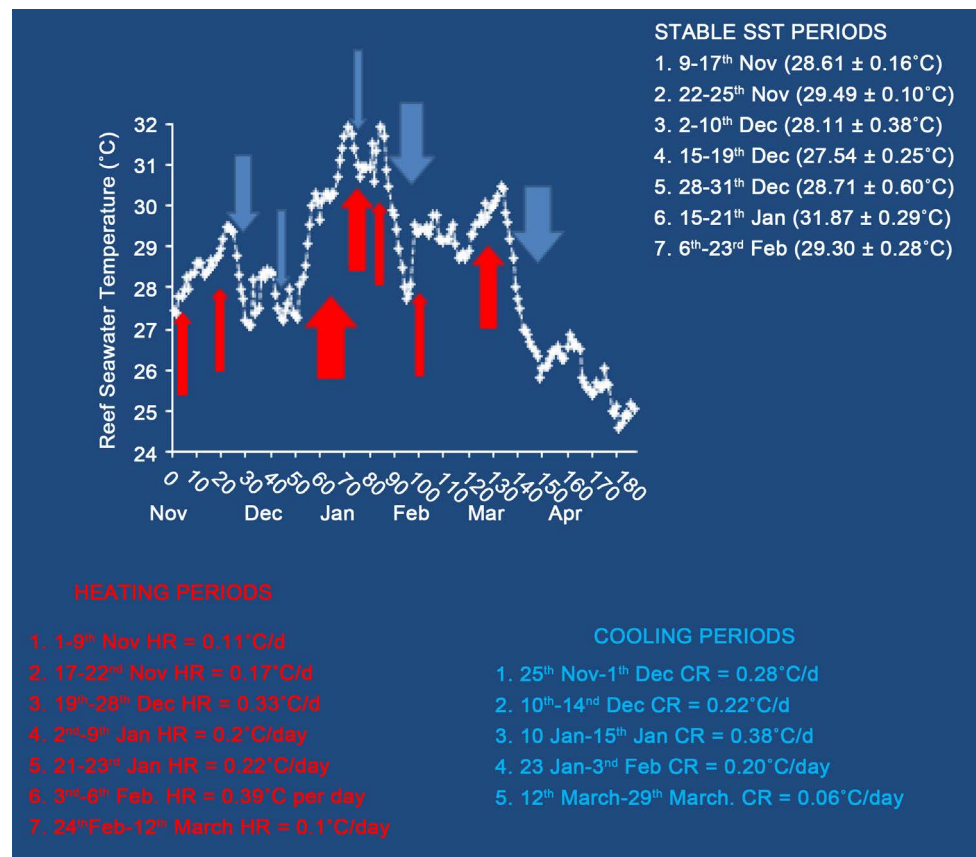


**Figure 4.** (a)  $\text{DMS}_w$  (nM), DMS flux ( $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) and lowest nighttime tide (m) at Nelly Bay reef. (b)  $\text{DMS}_w$  (nM), DMS flux ( $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) and lowest nighttime tide (m) at Nelly Bay and Pioneer Bay fringing reefs. (c) Seawater  $\text{DMS}_w$  (nM) plotted against lowest nighttime tide (m) for the two reef sites. (d) Calculated  $\text{DMS}_w$  (nM) and DMS flux ( $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) for Nelly Bay fringing reef. (e) Dissolved  $\text{NH}_4^+$  ( $\mu\text{M}$ ) at Nelly Bay fringing reef. (f) Zooxanthellae concentrations (zooxs  $\text{polyp}^{-1} \times 10^4$ ) in green and seawater  $\text{DMS}_w$  (nM) in blue for Nelly Bay reef.

$\text{DMS}_w$  increased (Figure 4(c)), suggesting that DMS fluxes increase from reefs as more of the reefs are exposed at low tides. Ammonium concentrations in Nelly Bay from 6<sup>th</sup> November 1993 to 2<sup>nd</sup> May 1994 ranged from 0.03 - 1.83  $\mu\text{M}$  (Figure 4(e)), with elevated levels in early November, decreasing to very low levels in early to mid-December, and then a huge increase from the 8<sup>th</sup> January (1.83  $\mu\text{M}$ ) during the coral bleaching period. Very high concentrations ( $>1 \mu\text{M}$ ) occurred up to 31<sup>st</sup> March, with levels decreasing again to very low levels in April and May (Figure 4(e)). Zooxanthellae densities in *Acropora formosa* started to rapidly decrease on 8<sup>th</sup> January as this coral started to bleach, reaching lowest levels on 13<sup>th</sup> February [44] (Figure 4(f)). On 31<sup>st</sup> March as SSTs decreased to 27°C at Nelly Bay, zooxanthellae densities reached  $6.5 \times 10^4$  zooxanthellae per polyp, and then slowly increased their numbers, stabilizing at  $10.5 \times 10^4$  zooxanthellae per polyp on 2<sup>nd</sup> May (Figure 4(f)). These concentrations are still very low for this coral suggesting this coral had not completely recovered from the bleaching, and may even have been bleaching in November to December. At Pioneer Bay, Orpheus Island no bleaching of corals was reported.

#### 4.4. Continuous SST Record

A continuous record of SSTs was recorded by AIMS at Geoffrey Bay (Magnetic Island) at 2.1 m depth on the reef flat. Several warming, cooling and stable periods of SSTs were evident over the November to March period (Figure 5). For example, from 1<sup>st</sup>-9<sup>th</sup> November SSTs increased from 27.55°C - 28.81°C at a rate of 0.11°C per day. SSTs then remained relatively stable from 9<sup>th</sup>-17<sup>th</sup> November



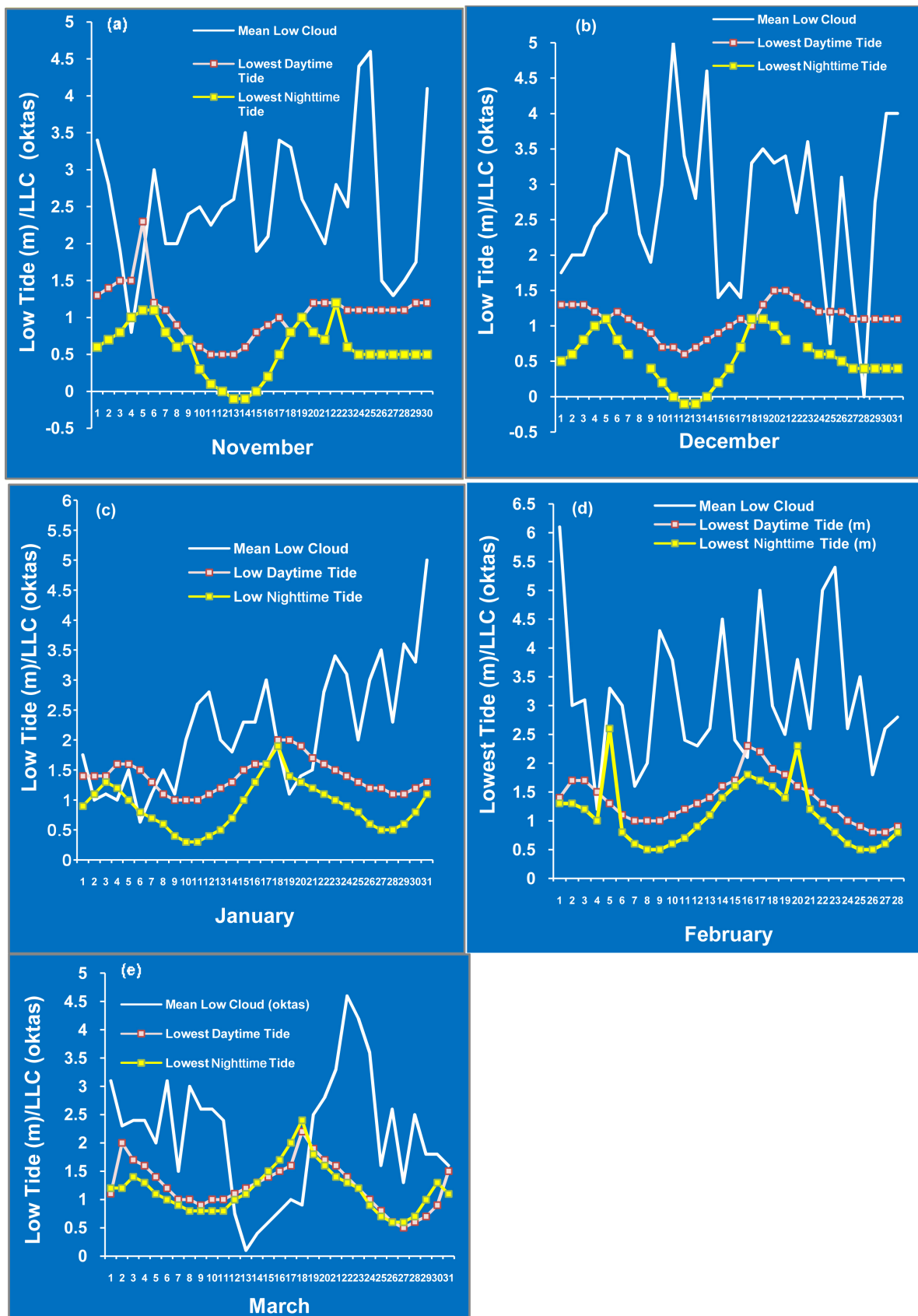
**Figure 5.** SST ( $^{\circ}\text{C}$ ) for Magnetic Island reef flat seawater from November-April. Heating and cooling periods, together with stable SST periods are highlighted, with the thickness of the arrows indicating the magnitude of the cooling and warming.

(mean  $28.61^{\circ}\text{C} \pm 0.16^{\circ}\text{C}$ ) until another warming period occurred from 17<sup>th</sup>-22<sup>nd</sup> November with a heating rate of  $0.17^{\circ}\text{C}$  per day. From 22<sup>nd</sup>-25<sup>th</sup> November another stable SST period occurred (mean  $29.49^{\circ}\text{C} \pm 0.11^{\circ}\text{C}$ ), followed by a cooling period from 25<sup>th</sup> November to 1<sup>st</sup> December when SSTs cooled from  $29.52^{\circ}\text{C}$  to  $27.31^{\circ}\text{C}$  at a rate of  $0.28^{\circ}\text{C}$  per day. This was followed by yet another period of stable SSTs from 2<sup>nd</sup>-10<sup>th</sup> December (Mean  $28.11^{\circ}\text{C} \pm 0.38^{\circ}\text{C}$ ). From 10<sup>th</sup>-14<sup>th</sup> December another cooling period occurred with SSTs decreasing from  $28.32^{\circ}\text{C}$  to  $27.2^{\circ}\text{C}$  at a cooling rate of  $0.22^{\circ}\text{C}$  per day. From 15<sup>th</sup>-19<sup>th</sup> December SSTs were relatively stable at  $27.54^{\circ}\text{C} \pm 0.25^{\circ}\text{C}$ . From the 18<sup>th</sup> December to 27<sup>th</sup> December, a period of 10 days, SSTs increased from  $27.37^{\circ}\text{C}$  to  $30.32^{\circ}\text{C}$  (Figure 5) increasing at a rate of  $\sim 0.30^{\circ}\text{C}$  per day. From 28<sup>th</sup> to 31<sup>st</sup> December a stable period of SSTs again occurred ranging from  $28.02^{\circ}\text{C}$  -  $29.40^{\circ}\text{C}$  (mean =  $28.71^{\circ}\text{C} \pm 0.6^{\circ}\text{C}$ ). SSTs from 2<sup>nd</sup> to 9<sup>th</sup> January rapidly increased from  $30.3^{\circ}\text{C}$  to  $34^{\circ}\text{C}$  (mean  $31.7^{\circ}\text{C}$ ) at a rate of  $\sim 0.46^{\circ}\text{C}$  per day. From 10<sup>th</sup>-15<sup>th</sup> January SSTs cooled significantly ( $0.38^{\circ}\text{C}$  per day) to  $30.71^{\circ}\text{C}$  on 15<sup>th</sup> January. Another stable but high SST period then occurred from 15<sup>th</sup> to 21<sup>st</sup> January, with SSTs stabilizing at  $31.87^{\circ}\text{C} \pm 0.25^{\circ}\text{C}$  (above the coral bleaching threshold), followed by a short, sharp warming period from 21<sup>st</sup>-23<sup>rd</sup> January when SSTs reached  $31.95^{\circ}\text{C}$  (increasing at  $0.22^{\circ}\text{C}$  per day). A sustained and rapid cooling period then occurred

from 23<sup>rd</sup> January to 3<sup>rd</sup> February, decreasing to 27.75°C at a rate of 0.20°C per day (**Figure 5**). From 3<sup>rd</sup>-6<sup>th</sup> February SSTs rapidly increased at the highest warming rate of 0.39°C per day, peaking at 29.56°C on 6<sup>th</sup> February. This was followed by the longest stable period of SSTs from 6<sup>th</sup> February to 23<sup>rd</sup> February (mean 29.30°C  $\pm$  0.28°C) when more regular pulses of LLC occurred (DOY 98-125; **Figure 3(b)**), and when significant amounts of rain fell over this period (**Figure 3(d)**). From 24<sup>th</sup> February to 12<sup>th</sup> March a second warming period occurred at a much lower warming rate of 0.1°C per day (28.72°C - 30.46°C), followed by decreasing SSTs from 12<sup>th</sup> March (30.26°C) to 29<sup>th</sup> March (25.81°C). SSTs then cooled to 24.9°C on 4<sup>th</sup> May.

#### 4.5. Low Level Clouds and Links to Low Tides and Rainfall

Mean daily LLC cover over Cleveland Bay reefs was compared with changes in the lowest nighttime and daytime tides (**Figures 6(a)-(e)**), and presence of rainfall (**Figure 3(d)**), as these natural factors can release large amounts of DMS<sub>a</sub> from coral reefs [10] [29]. These large pulses of DMS<sub>a</sub> can potentially assist LLC formation (see **Figure 1**). Increasing pulses of LLC occurred from 1-25<sup>th</sup> November and often coincided with very low nighttime tides (-0.1 to 0.8 m) (**Figure 6(a)**), with 25 mm rain falling between 13<sup>th</sup>-17<sup>th</sup> November during some of the lowest nighttime tides of the month (-0.1 - 0.5 m) (**Figure 3(c)**). Pulses of LLC cover also occurred from 16<sup>th</sup>-21<sup>st</sup> November on rising tides of 0.2 m to 1 m (no rainfall), whilst from 23<sup>rd</sup>-26<sup>th</sup> November a large pulse of LLC occurred during low nighttime tides of 0.5 - 0.6 m and high rainfall (77 mm) (**Figure 6(a)**, **Figure 3(c)**). On 30<sup>th</sup> November high levels of LLC coincided with a low nighttime tide of 0.5 m and 4 mm of rainfall. The lowest daytime tides from November-February were always higher than the lowest nighttime tides with both low tides often coinciding with elevated LLC (**Figures 6(a)-(e)**). In December, tides were again very low, mostly at night. Pulses of LLC cover often coincided with decreasing lowest nighttime tides (**Figure 6(b)**) and very high DMS fluxes of 6 - 8.6  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  (**Figure 4(d)**), in the more continuous flux profile, and more regular rainfall (**Figure 3(c)**). Rainfall was not as high as in November but occurred intermittently throughout December and often occurred during periods of the lowest nighttime tides (-0.1 to 0.6 m; **Figure 3(c)**). In the first week of January when corals bleached at Nelly Bay there was almost no LLC cover in the Townsville region (**Figure 6(c)**), whilst DMS fluxes were low ranging from 0.4 - 3.4  $\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  (**Figure 4(a)**) during medium low nighttime tides of ~1 m. From 9<sup>th</sup> January increasing pulses of LLC cover occurred over coral reefs in the Townsville region and again often occurred during periods of the lowest nighttime tides (0.3 and 0.4 m) and intense bursts of tropical rainfall on 10<sup>th</sup> January (54.4 mm) and 12<sup>th</sup> January (9.8 mm) (**Figure 3(c)**). From 22<sup>nd</sup> to 30<sup>th</sup> January rainfall ranged from 0.2 mm to 10.4 mm with the highest rainfall occurring on a 1m low tide on 23<sup>rd</sup> January (**Figure 3(c)**). On 31<sup>st</sup> January to 2<sup>nd</sup> February the highest levels of LLC cover occurred (3 - 5 oktas) during intense bursts of tropical rainfall and coincided with relatively high nighttime low tides (1 - 1.3 m) (**Figure 6(c)**, **Figure 6(d)**, **Figure 3(c)**). More regular pulses of LLC cover



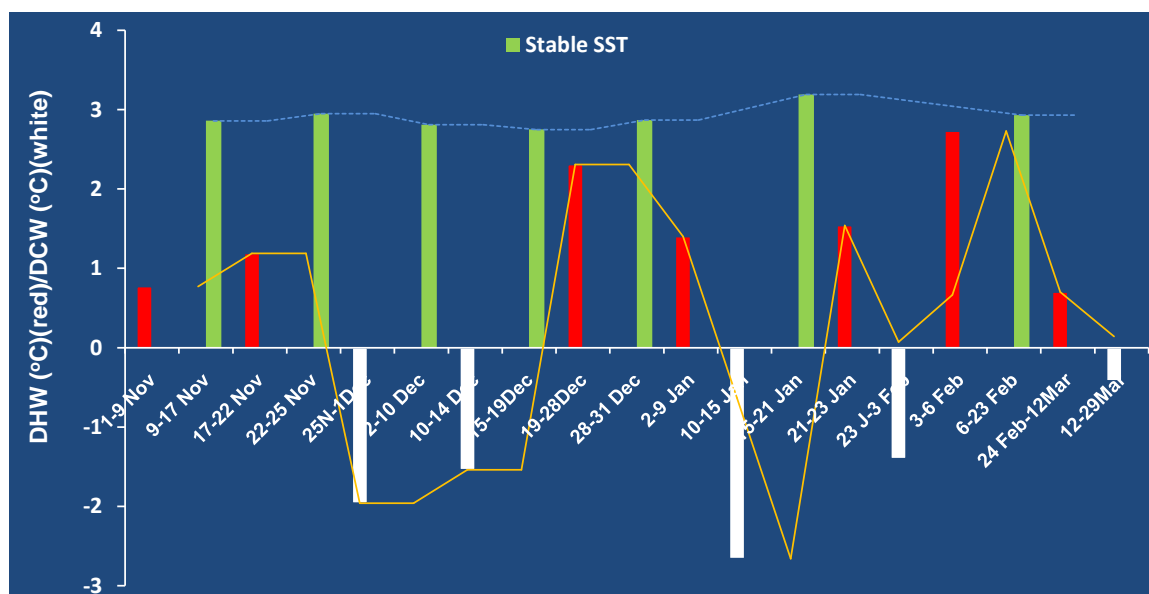
**Figure 6.** Mean low level cloud (LLC) in oktas and lowest daytime and nighttime tides (m) for the Townsville-Cleveland Bay region from November to March (a)-(e).

often occurred during low tides after 2<sup>nd</sup> February, and often coincided with rainfall bursts of <10 mm (**Figure 6(d)**). Smaller differences in low tides occurred between the lowest nighttime tides and lowest daytime tides in February (**Figure 6(d)**) and these prolonged periods of low tides and heavy rainfall (>70 mm) may have caused the extremely large DMS fluxes ( $20 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) from Pioneer Bay on 25<sup>th</sup> February (**Figure 6(d)**, **Figure 4(b)**, **Figure 3(c)**). In March the lowest nighttime tides ranged from 0.6 - 0.8 m with pulses of LLC cover often coinciding with lowest tides from 1<sup>st</sup>-6<sup>th</sup> March (**Figure 6(e)**), whilst high LLC cover occurred in the third week of March during decreasing daytime and nighttime low tides (1.5 m - 0.5 m) (e.g. 18<sup>th</sup>-24<sup>th</sup> March). No rainfall occurred in March.

#### 4.6. Alternating Warming and Cooling Periods

The heating and cooling periods of the Nelly Bay reef flat over the 5 month period (**Figure 5**) were first calculated as changes in SST per day and then were scaled up to “Degree Heating Week” (DHW) or “Degree Cooling Week” (DCW) as coral ecologists often use a similar index to compare coral bleaching events around the world. We have done this so that we can compare each warming and cooling period over a standard amount of time (*i.e.* 7 days). For example, the first heating period from 1-9<sup>th</sup> November (DHW +0.77°C) was followed by relatively stable SSTs from 9-17<sup>th</sup> November ( $28.61^\circ\text{C} \pm 0.16^\circ\text{C}$ ) (**Figure 7**). This was followed by a heating period from 17-22<sup>nd</sup> November with a DHW of +1.19°C followed by a SST stabilization period from 22-25<sup>th</sup> November of  $29.49^\circ\text{C} \pm 0.10^\circ\text{C}$ , followed by a significant cooling period from 25<sup>th</sup> November to 1<sup>st</sup> December with a DCW of -1.96°C which surprisingly was equal to the sum of the two previous DHWs (*i.e.* +0.77 +1.19 = +1.96°C). These SST heating and cooling periods in November (**Figure 7**) were generally accompanied by elevated levels of DMS<sub>w</sub>, DMS flux at Nelly Bay and Pioneer Bay (**Figure 4(a)**, **Figure 4(b)**), elevated LLC (**Figure 6(a)**) and intermittent rainfall (**Figure 3(c)**), during periods of low daytime and nighttime tides (0.26 - 0.83 m, mean 0.59 m) (**Figure 6(a)**). Over 9 days from 2-10<sup>th</sup> December (**Figure 7**) another SST stabilization period occurred ( $28.11^\circ\text{C} \pm 0.38^\circ\text{C}$ ). From 10-14<sup>th</sup> December another cooling period occurred with a DCW of -1.54°C, which was again followed by a SST stabilization period from 15-19<sup>th</sup> December ( $27.54^\circ\text{C} \pm 0.25^\circ\text{C}$ ) (**Figure 7**). From 19-28<sup>th</sup> December the second highest warming event occurred (**Figure 5**) giving a DHW of +2.31°C, with SSTs at 30.3°C, just over the thermal bleaching threshold of *Acropora*. This was followed by a SST stabilization period from 28-31<sup>st</sup> December when SSTs were ~28.7°C. From 15<sup>th</sup>-31<sup>st</sup> December DMS<sub>w</sub> and DMS flux gradually increased at Nelly Bay reef (**Figure 4(d)**) in the more continuous record, whilst LLC increased from 17-23<sup>rd</sup> and 28-31<sup>st</sup> December (**Figure 6(b)**). These periods of enhanced DMS<sub>w</sub> and DMS flux in December often occurred during the lowest nighttime tides (0.1 - 0.7 m; mean 0.52 m) and periods of high LLC (**Figure 6(b)**). From 2-9<sup>th</sup> January mean SSTs reached 31.7°C with a DHW of +1.4°C, followed by a short rapid cooling period from 10-15<sup>th</sup> January of





**Figure 7.** DCW (white) and DHW (red) in  $^{\circ}\text{C week}^{-1}$  and stable SSTs (green) ( $^{\circ}\text{C}$ ) for Magnetic Island reef flat waters at 2.1 m depth for the summer period November to March.

$-2.66^{\circ}\text{C}$  (Figure 7). This was followed by high mean SSTs from 15-21<sup>st</sup> January of  $31.9^{\circ}\text{C}$  and relatively high  $\text{DMS}_w$  and DMS fluxes from the reefs (Figure 4(a), Figure 4(b), Figure 4(d)).

From 21-23<sup>rd</sup> January a short warming period occurred with a DHW of  $+1.54^{\circ}\text{C}$  coinciding with enhanced DMS fluxes from Pioneer Bay reef ( $14 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) (Figure 4(b)), and increasing levels of LLC from 21<sup>st</sup> January to 3<sup>rd</sup> February (Figure 6(c), Figure 6(d)). This was followed by a similar magnitude cooling period from 23<sup>rd</sup> January to 3<sup>rd</sup> February of  $-1.4^{\circ}\text{C}$  (Figure 7). A short intense warming period from 3-6<sup>th</sup> February recorded the highest DHW of  $+2.73^{\circ}\text{C}$  when  $\text{DMS}_w$  and DMS fluxes were very low on 3<sup>rd</sup> February ( $0.7 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) (Figure 4(a)) and corals were still stressed as evidenced by low zooxanthellae counts in *Acropora* (Figure 4(f)). From 6<sup>th</sup>-23<sup>rd</sup> February an exceptionally long period of stable SSTs of  $29.3^{\circ}\text{C} \pm 0.28^{\circ}\text{C}$  occurred (Figure 7), coinciding with regular pulses of LLC (Figure 6(d)) and high DMS fluxes at Pioneer Bay reef ( $14 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) in late January, and on 25<sup>th</sup> February) ( $20 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) (Figure 4(b)), and in the more continuous flux record (Figure 4(d)). The Pioneer Bay reef did not experience bleaching since SSTs were  $< 30^{\circ}\text{C}$  over the 19 month study, whilst at Nelly Bay reef in early to mid-February DMS flux was very low at  $\sim 0.4 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  and SSTs were exceptionally high  $30^{\circ}\text{C}$  -  $34^{\circ}\text{C}$ . A small warming period of  $+0.7^{\circ}\text{C}$  DHW occurred from 24 February-12<sup>th</sup> March, followed by a cooling period of  $-0.42^{\circ}\text{C}$  as the region exited the warm summer months.

#### 4.7. Regression Analysis

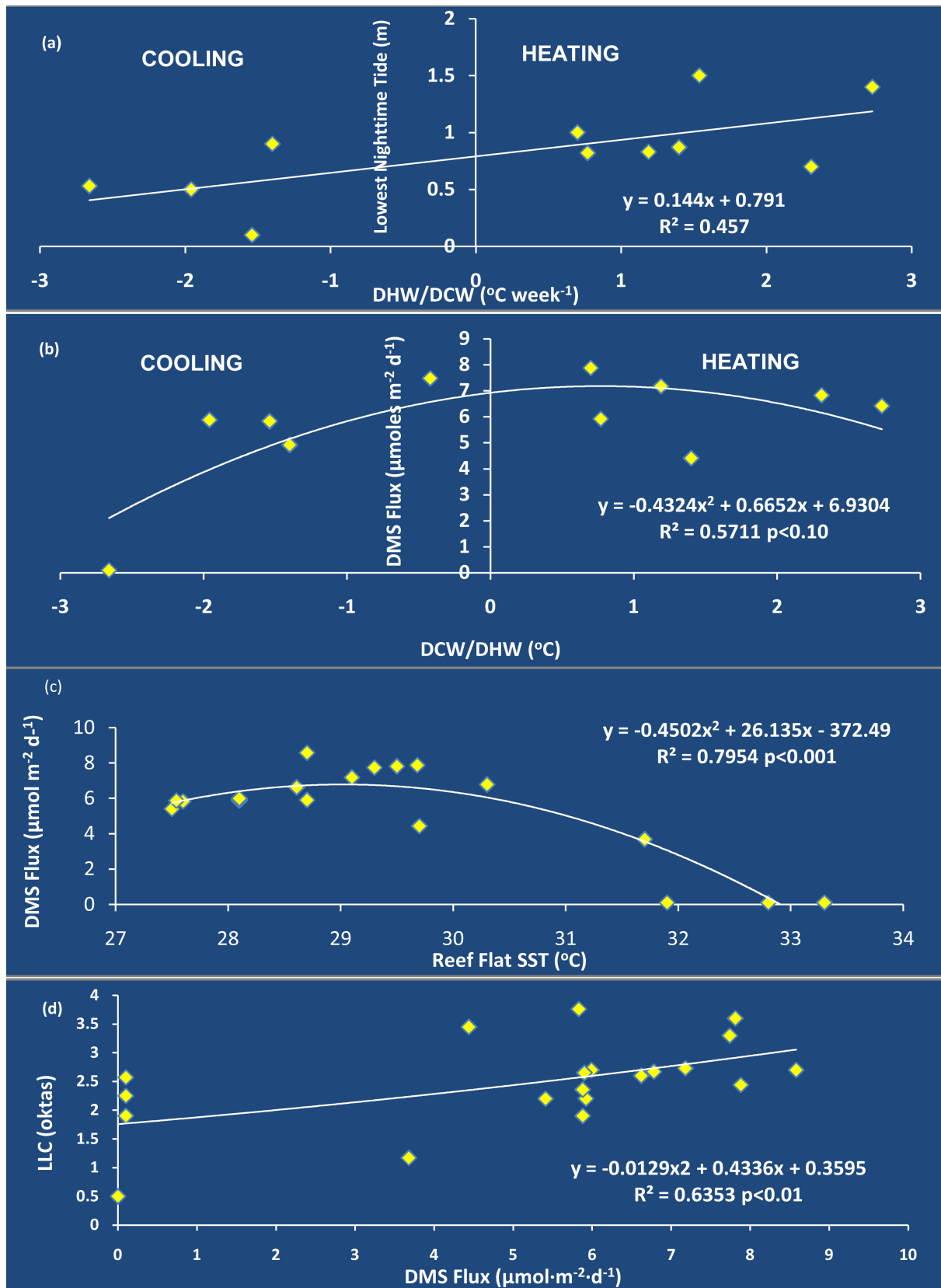
Solar radiation was significantly positively correlated with  $\text{DMS}_w$  and DMS flux ( $p < 0.05$ ) indicating that as solar radiation increases both  $\text{DMS}_w$  and DMS flux also increase (see Table 1). One of the most interesting findings in this work was

**Table 1.** Regression analysis and level of significance for Nelly Bay and Pioneer Bay.

Correlation	Equation	Regression	Significance
DMS <sub>w</sub> -SR (NB)	DMS <sub>w</sub> = 0.56 SR - 13.9	r <sup>2</sup> = 0.73 (9)	p < 0.05
DMS Flux-SR (NB)	DMS flux = 0.054 SR <sup>2</sup> - 2.29 SR + 24	r <sup>2</sup> = 0.72 (9)	p < 0.05
DMS Flux-SR (NB + PB) <sup>#</sup>	DMS flux = -0.15 SR <sup>2</sup> + 1.9 SR + 24	r <sup>2</sup> = 0.48 (13)	p < 0.10
DMS <sub>w</sub> -SST (PB)	DMS <sub>w</sub> = 0.55 SST - 12	r <sup>2</sup> = 0.44 (76)	p < 0.001*
DMS <sub>w</sub> -SST (NB)	DMS <sub>w</sub> = 0.9 SST - 23.6	r <sup>2</sup> = 0.83 (8)	p < 0.01
DMS Flux-SST (NB)	DMS flux = -0.46 SST <sup>2</sup> + 26 SST - 373	r <sup>2</sup> = 0.80 (19)	p < 0.001*
DMS Flux-SST (NB) <sup>##</sup>	DMS flux = 1.14 SST - 25.8	r <sup>2</sup> = 0.67 (13)	p < 0.01
DMS <sub>w</sub> -NH <sub>4</sub> <sup>+</sup> (NB)	DMS <sub>w</sub> = -1.67 NH <sub>4</sub> <sup>+</sup> + 3.9	r <sup>2</sup> = 0.99 (6)	p < 0.001*
NH <sub>4</sub> <sup>+</sup> -SST (NB)	NH <sub>4</sub> <sup>+</sup> = 0.19 SST - 4.9	r <sup>2</sup> = 0.56 (12)	p < 0.05
DMS <sub>w</sub> -Zoox (NB)	DMS <sub>w</sub> = 0.29 Zoox - 1.4	r <sup>2</sup> = 0.43 (12)	p > 0.05
Zoox-DMS Flux (NB)	Zoox. = 0.81 DMS Flux + 7	r <sup>2</sup> = 0.42 (12)	p > 0.02
NH <sub>4</sub> <sup>+</sup> -Zoox (NB)	NH <sub>4</sub> <sup>+</sup> = -0.14 Zoox + 2	r <sup>2</sup> = 0.50 (10)	p < 0.10
DMS <sub>w</sub> -LNT (NB + PB)	LNT = 0.016 DMS <sub>w</sub> <sup>2</sup> - 0.22 DMS <sub>w</sub> + 1.29	r <sup>2</sup> = 0.57 (16)	p < 0.02
DMS Flux-DMS <sub>w</sub> (NB + PB)	DMS flux = 2.25 DMS <sub>w</sub> - 1.7	r <sup>2</sup> = 0.67 (16)	p < 0.001*
DMS <sub>w</sub> -LLC (NB + PB)	DMS <sub>w</sub> = 0.34 LLC <sup>2</sup> + 2.28 LLC - 0.06	r <sup>2</sup> = 0.70 (16)	p < 0.001*
SR-LLC (NB)	SR = -0.88 LLC <sup>2</sup> + 2.4 LLC + 26.7	r <sup>2</sup> = 0.38 (100)	p < 0.001*

Units: DMS<sub>w</sub> = seawater DMS (nM). SR = Solar radiation (MJ·m<sup>-2</sup>). DMS flux = μmol·m<sup>-2</sup>·d<sup>-1</sup>. NH<sub>4</sub><sup>+</sup> = μmol. Zooxanthellae concentration = zooxanthellae polyp<sup>-1</sup> × 10<sup>4</sup>. LLC = Low Level Cloud (oktas). LNT = Lowest night-time tide (m). NB = Nelly Bay. PB = Pioneer Bay. \* = highly significant. # = Highest DMS flux values for PB omitted in correlation. ## = SST < 30°C.

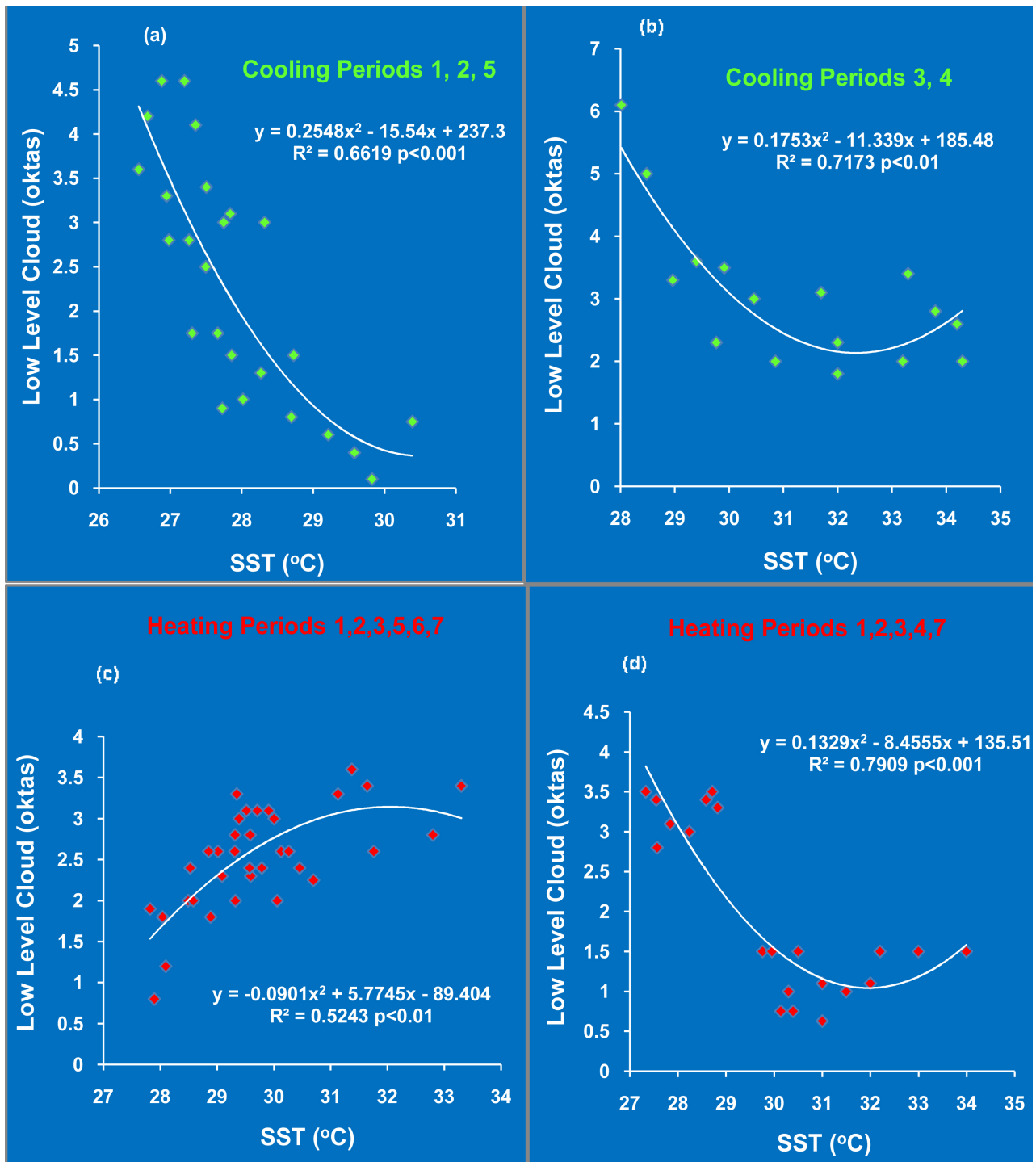
the very significant positive correlation of DMS<sub>w</sub> with SST (r<sup>2</sup> = 0.44, p < 0.001) for Pioneer Bay fringing reef (see **Table 1**). However, there was no significant correlation of DMS<sub>w</sub> and SSTs for Nelly Bay reef if we included SSTs > 30°C. However, when SSTs > 30°C are excluded there was a significant positive correlation between DMS<sub>w</sub> and SSTs for Nelly Bay reef (r<sup>2</sup> = 0.83, p < 0.01). Interestingly almost the same linear regression equation was obtained for the DMS<sub>w</sub>-solar radiation correlation for Nelly Bay reef as the DMS<sub>w</sub>-SST correlation for Pioneer Bay reef (see **Table 1**). However, the DMS<sub>w</sub>-SST linear regression equation for Nelly Bay had a much greater slope by as much as 162%, compared with Pioneer Bay. It is interesting that the regression equation that correlates the lowest nighttime tide with DCW and DHW (**Figure 8(a)**), although just short of significance at p > 0.1, indicates that cooling of reef flat waters generally occurs during low nighttime tides that ranged from 0.5 - 0.9 m (mean 0.5 m). In contrast warming of reef flat waters generally occurred during tides that ranged from 0.7 - 1.4 m (mean 0.97 m) which seems the reverse of what you might expect. Also the regression equation that describes the significant (p < 0.10) relationship between DMS flux and DCW and DHW (**Figure 8(b)**) indicates that cooling of reef flat waters generally occurred at lower DMS fluxes that ranged from 0.1 - 7.48 μmol·m<sup>-2</sup>·d<sup>-1</sup> (mean 4.8), whilst warming of reef flat waters generally occurred at slightly higher DMS fluxes that ranged from 4.4 - 7.8 μmol·m<sup>-2</sup>·d<sup>-1</sup> (mean 6.4). DMS flux decreased linearly with DCW, although the correlation was not significant at p > 0.10 (**Figure 8(b)**). DMS flux was very significantly (p < 0.001) correlated with reef flat SSTs (r<sup>2</sup> = 0.80) for SSTs ranging



**Figure 8.** Results for Nelly Bay fringing reef flat. (a) DHW and DCW plotted against the lowest nighttime tides (m). (b) DMS flux ( $\mu\text{mol m}^{-2} \text{d}^{-1}$ ) versus DCW/DHW ( $^{\circ}\text{C week}^{-1}$ ). (c) DMS flux ( $\mu\text{mol m}^{-2} \text{d}^{-1}$ ) versus reef flat SSTs ( $^{\circ}\text{C}$ ). (d) LLC versus DMS flux ( $\mu\text{mol m}^{-2} \text{d}^{-1}$ ).

from 27.5°C to 33.3°C, with the correlation described by a polynomial equation ( $n = 2$ ) (**Figure 8(c)**). As SSTs breached the coral bleaching threshold and reached maximum SSTs of 33.3°C, DMS flux decreased to nd levels (**Figure 8(c)**). When SSTs > 30°C were omitted from the correlation DMS flux was significantly ( $p < 0.01$ ) linearly correlated with increases in SSTs from 27.5°C to 29.68°C (**Figure 8(c)**). Seawater  $\text{NH}_4^+$  was also significantly positively correlated with SSTs at Nelly Bay reef ( $r^2 = 0.56$ ,  $p < 0.05$ ) (**Table 1**).  $\text{DMS}_w$  measured at Nelly Bay and Pioneer Bay reefs was significantly positively correlated with the lowest nighttime tides ( $r^2 = 0.57$ ,  $p < 0.02$ ) with the correlation predicting that as the lowest nighttime tide decreases,  $\text{DMS}_w$  often increases. Also of interest was the highly significant positive correlation of  $\text{DMS}_w$  and DMS flux for Nelly Bay and Pioneer Bay ( $r^2 = 0.67$ ,  $p < 0.001$ ) (see **Table 1**). Whilst the correlation between zooxanthellae counts in *Acropora formosa* and DMS flux was not quite significant ( $p > 0.10$ ), more data points may establish a statistically significant correlation. However, the correlation suggests that as zooxanthellae counts decrease in *Acropora* coral the flux of DMS from reefs decreases. There was also a significant correlation of  $\text{NH}_4^+$  with zooxanthellae counts ( $r^2 = 0.50$ ,  $p < 0.10$ ), and an almost significant correlation between  $\text{DMS}_w$  and zooxanthellae counts ( $r^2 = 0.43$ ,  $p > 0.10$ ). Whilst the overall correlation between  $\text{NH}_4^+$  and  $\text{DMS}_w$  was not significant, when data points from the 13<sup>th</sup> February are not included (*i.e.* when *Acropora* reached its lowest zooxanthellae concentrations) the correlation was highly statistically significant ( $r^2 = 0.99$ ,  $p < 0.001$ ). Whilst more data points are needed the  $\text{DMS}_w$ -LLC correlation is highly significant ( $p < 0.001$ ) and suggests that as  $\text{DMS}_w$  in reef environments increases LLC cover also increases (see **Table 1**). The correlation between DMS flux and LLC is also highly significant (**Figure 8(d)**) (see **Table 1**).

Different LLC-SST correlations occurred for different cooling and heating periods experienced at the Nelly Bay fringing reef. For example, for cooling periods not affected by coral bleaching (*i.e.* cooling periods 1, 2, 5), increasing levels of LLC from ~0.5 - 4.5 oktas were positively correlated with SSTs ( $p < 0.001$ ) when SSTs were < 30°C (**Figure 9(a)**). However, for cooling periods after coral bleaching (*i.e.* cooling periods 3, 4), SSTs were negatively correlated with LLC ( $p < 0.01$ ) and indicated that SSTs could be extremely high (>30°C) at very low levels of LLC cover (*i.e.* 2 - 3 oktas) (**Figure 9(b)**). The regression equation also indicated that this cooling took some time to reduce SSTs to < 30°C. For the heating periods (*i.e.* periods 1, 2, 3, 5, 6, 7) that did not include the coral bleaching period 4 (*i.e.* 2-9<sup>th</sup> January), increasing LLC was significantly positively correlated with increased SSTs up to 33.5°C ( $p < 0.01$ ) (**Figure 9(c)**). For heating periods that included the coral bleaching event (*i.e.* periods 1, 2, 3, 4, 7), but not period 5 (21-23<sup>rd</sup> January) and period 6 (the highest heating rate from the 3<sup>rd</sup>-6<sup>th</sup> February), decreasing levels of LLC were accompanied by increases in SSTs as you might expect (**Figure 9(d)**). Solar radiation levels were significantly correlated with LLC ( $p < 0.001$ ) (**Figure 10**). However, although increasing LLC generally attenuated solar radiation there were clearly occasions when extreme

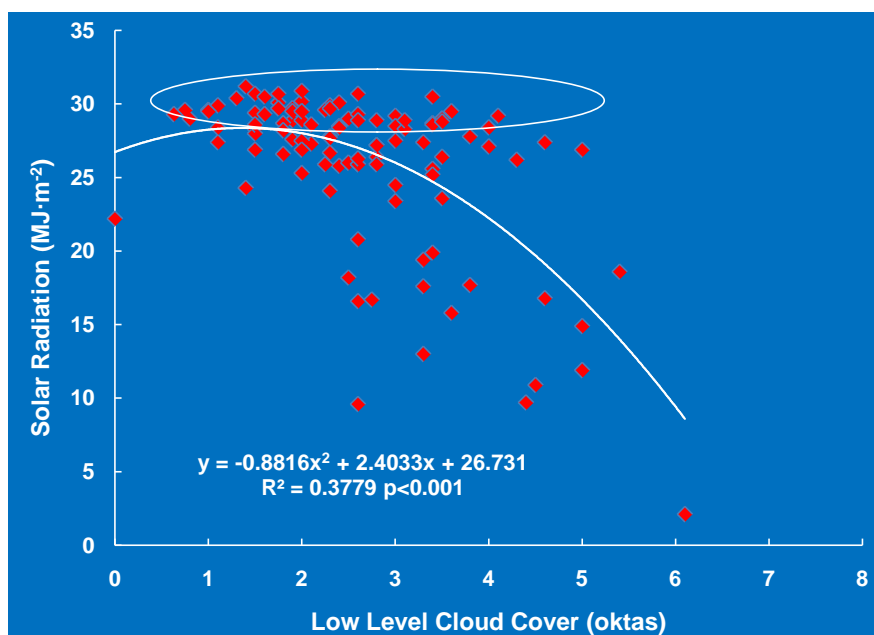


**Figure 9.** LLC (oktas) plotted against SSTs for different cooling (a, b) and heating periods (c, d) for Nelly Bay reef flat waters.

solar radiation ( $\sim 29 - 30 \text{ MJ}\cdot\text{m}^{-2}$ ) occurred at both low levels of LLC ( $< 2$  oktas) and medium levels of LLC (3 - 4 oktas) (circled area, **Figure 10**).

## 5. Discussion

A detailed comparison of mean daily LLC with the lowest daytime and nighttime tides indicated that on many occasions high levels of LLC often occurred on days



**Figure 10.** Mean daily solar radiation levels in  $\text{MJ}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  are plotted against LLC cover at Townsville airport for the November-April period. [Circled area = values of extreme solar radiation at low and medium LLC].

with the lowest daytime and nighttime tides from November-February (**Figures 6(a)-(d)**), and often coincided with high sporadic pulses of  $\text{DMS}_w$  and  $\text{DMS}$  fluxes (**Figures 4(a)-(d)**). For example, from November-December high levels of LLC often occurred during low tides and high  $\text{DMS}$  flux (**Figures 4(a)-(d)**, **Figures 6(a)-(d)**). However, in the first week of January when corals were bleaching at Nelly Bay low levels of LLC occurred ( $<2$  oktas) when  $\text{DMS}$  fluxes from Pioneer Bay and Nelly Bay reefs were very low ( $0.4 - 1.4 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  (**Figure 4(b)** and **Figure 4(c)**), coinciding with high solar radiation levels and high SSTs (**Figure 3(b)**, **Figure 5**). From 9<sup>th</sup> January to 3<sup>rd</sup> February the flux of  $\text{DMS}$  emitted from the two reefs ranged from  $0.2 - 14 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  (mean 5), with the highest flux emitted from the more pristine Pioneer Bay reef on 24<sup>th</sup> January ( $14 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) after significant rainfall in the region (10 mm at Townsville airport on 23<sup>rd</sup> January), and when a nighttime low tide of 0.9 m occurred (**Figure 6(c)**).

From late January to early December mean daily LLC increased dramatically reaching levels as high as 5 - 6 oktas (**Figure 6(d)**). Regular pulses of high LLC then occurred for the rest of February, often occurring during low tides and sporadic pulses of very high  $\text{DMS}$  flux at the Pioneer Bay reef (**Figure 4(b)**). For example, on 25<sup>th</sup> February extremely high fluxes of  $\text{DMS}_a$  were emitted from the fringing reef at Pioneer Bay ranging from  $19 - 20 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$  (mean 19.5), when medium low tides again occurred (0.8 m) (**Figure 4(b)**). These huge fluxes of  $\text{DMS}$  emitted sporadically from Pioneer Bay reef were in contrast to the very low fluxes of  $\text{DMS}$  that were emitted from 13<sup>th</sup> February to 2<sup>nd</sup> May from the Nelly Bay fringing reef ( $\text{nd}-0.8 \mu\text{mol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) where coral bleaching occurred and zooxanthellae counts in *Acropora formosa* reached their lowest concentra-



tions on 13<sup>th</sup> February (**Figure 4(f)**). Whilst high levels of LLC in the summer can lower solar radiation levels and SSTs thus protecting coral reefs from the full effects of high solar radiation levels [23] [24] optically thin LLC can also increase solar radiation levels in the GBR [47] [48].

For example, elevated LLC (>4 oktas) can clearly attenuate solar radiation (**Figure 10**), but extreme solar radiation levels ( $\sim 30 \text{ MJ}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) occurred in the local region at both low (1 - 2 oktas) and medium (3 - 4 oktas) levels of LLC (see circled area in **Figure 10**), suggesting a positive radiative effect from low and medium levels of LLC as reported for this region [47]. From November to early January over a period of 29 days solar radiation levels were often extreme ( $>29 \text{ MJ}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ). From 19<sup>th</sup>-28<sup>th</sup> December solar radiation levels often ranged from 29 - 30  $\text{MJ}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ , contributing to the second highest DHW of  $+2.31^\circ\text{C}$  in Magnetic Island reef flat waters (**Figure 7**), with SSTs ranging from  $27.33^\circ\text{C}$  -  $31.76^\circ\text{C}$ , strongly suggesting that coral bleaching could have occurred earlier than the first week of January. This suggestion is supported by the extremely low levels of zooxanthellae in *Acropora formosa* at Nelly Bay from 6<sup>th</sup> November to 16<sup>th</sup> December ( $1.2 - 1.4 \times 10^5$  zooxanthellae polyp<sup>-1</sup>), compared with zooxanthellae counts in recovering *Pocillopera damicornis* colonies of  $2.67 \times 10^6$  at the same location in July [31]. Bleaching of *A. formosa* in early January further reduced zooxanthellae concentrations to  $0.5 \times 10^5$  zooxanthellae polyp<sup>-1</sup> on 13<sup>th</sup> February (**Figure 4(f)**). The solar radiation levels over Nelly Bay reef in January were exceptional and bleached the upper surfaces of *Acropora* coral at Nelly Bay reef [46]. These high levels of solar radiation undoubtedly contributed to the very high SSTs of  $30.5^\circ\text{C}$  -  $33.3^\circ\text{C}$  at Nelly Bay in the first week of January [41] [44] and most probably led to the corals bleaching more severely, with corals continuing to lose zooxanthellae right up to mid-February (**Figure 4(f)**). However, directly after the bleaching from the 10<sup>th</sup>-15<sup>th</sup> January Magnetic Island surface waters cooled significantly with a DCW of  $-2.66^\circ\text{C}$  (**Figure 7**), most probably aided by the high levels of LLC (**Figure 6(c)**), which periodically reached 4 - 6 oktas (mean of 3 hourly observations) and produced 64 mm of rainfall (**Figure 3(c)**). From 19<sup>th</sup>-23<sup>rd</sup> January, daily air temperatures included a series of unusually high minimum daily temperatures, including the second highest minimum air temperature ( $29^\circ\text{C}$ ) recorded locally since 1942 at the time and which coincided with increasing levels of LLC during the nighttime (**Figure 6(c)**) [44]. Daily air temperatures then fell sharply from  $29.3^\circ\text{C}$  on 24<sup>th</sup> January, to  $24.8^\circ\text{C}$  on 31<sup>st</sup> January 1994, aided by very high levels of LLC from 21<sup>st</sup> January to 3<sup>rd</sup> February (**Figure 6(c)**, **Figure 6(d)**).

As solar radiation levels increased so did  $\text{DMS}_w$  and DMS flux and were significantly and directly correlated with each other (see **Table 1**), possibly reflecting increased photosynthesis of reef organisms that produce DMS [28] [49], and/or effects of SST and tidal stress that releases  $\text{DMS}_w$  and  $\text{DMS}_a$  from coral reefs [10] [20] [33]. Production of coral mucus enriches reefs with extremely high DMS concentrations, particularly when stressed [34] [49], and together with other volatile organic substances such as isoprene from corals [50] can ra-

pidly enhance the oxidation of  $\text{DMS}_a$ . As solar radiation levels increased over coral reefs so did  $\text{DMS}_w$  and DMS flux and explains why there is such a good correlation of  $\text{DMS}_w$  with increasing SST for Pioneer Bay reef (see **Table 1**). However, no such correlation occurred for Nelly Bay reef when SSTs  $> 30^\circ\text{C}$  were included. When SSTs  $> 30^\circ\text{C}$  were excluded a significant DMS-SST correlation was also observed for Nelly Bay reef (**Table 1**).  $\text{NH}_4^+$  concentrations were also significantly positively correlated with increasing SSTs and significantly negatively correlated with  $\text{DMS}_w$  (see **Table 1**) as previously reported [1]. The huge increase of  $\text{NH}_4^+$  concentrations from January to February often coincided with high sporadic fluxes of DMS from Nelly Bay and Pioneer Bay reefs (**Figure 4(b)**, **Figure 4(e)**) and could explain the high  $(\text{NH}_4)_2\text{SO}_4$  concentrations (sulfate aerosol) that occur periodically in marine air masses over healthy coral reefs in the GBR [26]. This evidence strongly suggests there may be a link between  $\text{DMS}_a$  emitted from coral reefs in the GBR, sulfate aerosol particles from oxidation of  $\text{DMS}_a$  [26], and the elevated levels of condensation nuclei (CN) measured over the GBR [51] which are precursors of CCN and LLC (**Figure 1**).

The chemical composition of CCN is predominantly sulfate neutralized to varying degrees by ammonium [52]. Once a cloud droplet is formed it will scavenge ammonia (Liss and Galloway, 1993 [53]). This production of  $\text{NH}_4^+$  is vital to the formation of sulfate aerosol since it charge balances the sulfate in the atmosphere (**Figure 1**). Much of the light scattering properties of sulfate aerosol has a range of compositions from  $\text{H}_2\text{SO}_4$  to  $(\text{NH}_4)_2\text{SO}_4$ , which can enhance solar radiation levels [7]. The interaction of the sulfur and reduced nitrogen cycles in the atmosphere over the GBR can influence the productivity of the GBR. The more sulfate aerosols present in the atmosphere, the greater the rate of nitrogen deposition due to scavenging of ammonia by sulfuric acid. In this way sulfate aerosols can essentially act as a carrier of nutrient nitrogen into GBR waters which could be greater than land-inputs of nitrogen [53]. Further studies are urgently needed to quantify this atmospheric nitrogen and sulfur input to the GBR [J. Ivey; PhD studies].

Continuous SSTs measurements for Magnetic Island reef flat waters revealed various heating and cooling periods, interspersed with stable SSTs (**Figure 7**). The most extreme heating occurred in late December and early February when solar radiation levels were extreme and LLC was high (**Figures 3(a)-(c)**). Most cooling periods (negative climate feedback) occurred during low tides ( $<0.8$  m) and decreasing DMS fluxes (**Figure 8(b)**, **Figure 8(c)**), whilst most heating periods were often associated with higher low tides ( $>0.8$  m) and elevated DMS fluxes, solar radiation and LLC (**Figure 8(b)**, **Figure 8(c)**). Alternation between these two states or types of feedback (cooling and warming) is indicated in this research (**Figure 8(a)**, **Figure 8(b)**, **Figure 9**). A cooling period (period 4 in **Figure 5**) of DCW  $-1.4^\circ\text{C}$ , associated with a low tide of 0.8 m coincided with high levels of LLC and rainfall of 163 mm during the 23<sup>rd</sup> January to 3<sup>rd</sup> February period (**Figure 3(d)**). The combination of low spring tides and rainfall on the

aerially exposed reef flat could result in extremely high DMS fluxes from the two coral reefs and could explain the high levels of LLC over reefs in the region from 23<sup>rd</sup> January to 3<sup>rd</sup> February (**Figures 6(c)-(d)**), causing lowered SSTs. In our study the cooling and heating periods identified were significantly correlated with DMS fluxes from the two reefs ( $p < 0.10$ ) and suggest that as the flux of DMS from coral reefs increases cooling of reef flat waters occurs (**Figure 7, Figure 8(c)**). This could reflect increases in reef-generated  $\text{DMS}_a$  and oxidation to sulfate aerosol and CCN formation (Modini *et al.* 2009), leading to enhanced LLC and periodic rainfall (**Figure 1**). Whilst detailed atmospheric measurements still need to be carried out over the GBR to more fully understand this climate feedback there is now more evidence that SST cooling by LLC does in fact occur in the wider GBR [18] [20] [23], thus implicating reef emissions of  $\text{DMS}_a$ .

This cooling trend in the central GBR has been examined across all study summers and shelf positions [23]. Leahy *et al.* found SSTs exhibited distinct drops during periods of high cloud cover, and conversely SST increases during periods of low cloud cover, with a 3-day temporal lag between a change in cloud cover and a subsequent change in SST. The relationship was strongest in both El Nino (2005) and La Nina (2008) study summers, at the inner-shelf position in those study summers, where the LLC cover tends to accumulate [23]. Interestingly these authors found occasions when SSTs decreased by  $\sim 0.5^\circ\text{C} - 1.5^\circ\text{C}$  during cloud cover increases (January and March), and situations where SSTs increased (February), a situation we have found in our study.

In our study different LLC-SST correlations occurred for different cooling and heating periods. For cooling periods in late November, mid-December, and almost all of March (periods not affected by coral bleaching), increasing levels of LLC were significantly correlated ( $p < 0.001$ ) with decreases in SSTs (**Figure 9(a)**). However, in mid-January very high SSTs ranging from  $31^\circ\text{C} - 33^\circ\text{C}$  occurred in Magnetic Island reef flat seawater and coincided with very low levels of LLC cover ( $< 2.5$  oktas) (**Figure 4(c)**). For a cooling period from 23<sup>rd</sup> January to 3<sup>rd</sup> February, high levels of LLC (**Figure 4(c), Figure 4(d)**) reduced SSTs to below  $30^\circ\text{C}$  (**Figure 5**), coinciding with enhanced DMS fluxes from Pioneer Bay reef within this period (**Figure 4(b), Figure 4(d)**). During almost all heating periods two correlations again occurred between LLC and SSTs. For example, for heating periods that did not include the coral bleaching event in the first week of January (*i.e.* periods 1, 2, 3, 5, 6, 7) (**Figure 9(c)**), increasing levels of LLC were accompanied by increases in SSTs. For heating periods 1, 2, 3, 4, 7 (which included the coral bleaching period 4 from 2-9<sup>th</sup> January) decreases in LLC were accompanied by increases in SSTs as one might expect (**Figure 9(d)**). It is thought that these contrasting effects on the LLC-SST correlation reflect the high solar radiation levels that can occur when LLC contains high concentrations of small diameter (nm) sulfate aerosol particles [26], generated by the oxidation of  $\text{DMS}_a$  produced from the reefs during low tides [28] [29]. These increases in  $\text{DMS}_a$  have been observed as spikes or pulses, lasting from  $\sim 1$  to 8.5 h, at Heron Island coral cay in the southern GBR [10]. These researchers have also observed

that when convective rain falls onto exposed coral reefs in the GBR at low tides even higher spikes of  $\text{DMS}_a$  can occur over reefs, possibly caused by intracellular changes in DMSP in corals [31]. It is thought that these natural stress events (low tides, high SSTs, rainfall) cause the coral to utilize intracellular DMSP to cope with the resultant osmotic, solar radiation and temperature stresses [10] [31] [32] [33], and in the process produce enhanced levels of seawater and atmospheric DMS. Swan *et al.*, [10] [50] have proposed that CCN produced from coral reef-derived  $\text{DMS}_a$  oxidation products could contribute to cloud formation in the southern GBR as part of the convective internal boundary layer (CIBL) produced over Heron Reef [54]. These sporadic spikes of enhanced fluxes of  $\text{DMS}_a$  occurred intermittently over the Heron Reef and were 2 to 30 times the DMS flux from the adjacent ocean surrounding the coral reef [10] [55]. According to Swan *et al.*, [10] the island cay could assist the convective upward movement of DMS and its oxidation products via the CIBL to form localized convection clouds which may precipitate over the island and surrounding reef. Processes such as this could generate the proposed coral reef climate feedback described in this paper. Such a process would lower reef flat SSTs, which has been observed in this study at Magnetic Island.

In contrast, the high SSTs that accompany low or medium levels of LLC, could reflect increased scattering of solar radiation from a build-up of these small diameter sulfate aerosols in the marine boundary layer [7] [18] and could explain the presence of optically thin clouds that are often observed during coral bleaching in the GBR (see **Figure 6** in Elvidge *et al.* [56]). Changes in nss-sulfate aerosol number and LLC cloud amount could increase the long-wave and short-wave cloud radiative forcing [57] over coral reefs in the GBR, as suggested by the good correlations between  $\text{DMS}_w$  and DMS flux and solar radiation (see **Table 1**), and could exacerbate coral bleaching events from this increased scattering of solar radiation. The emission of aerosol particles from the 1,000 km length of reefs in the GBR has been calculated to be  $10^{19}$  particles  $\text{s}^{-1}$ , comparable with emissions from land surfaces and forest fires [51]. These aerosol particles over the GBR consist of 60% ammonium sulfate and 40% non-DMS volatile organics, oxidation products of  $\text{DMS}_a$  [26]. The results of Modini *et al.* [26] are the first direct observation of aerosol particle formation over the GBR with exceptionally high growth rates and very high aerosol particle concentrations. In light of our recent research these high aerosol concentrations could clearly reflect enhanced levels of  $\text{DMS}_a$  emitted from coral reefs in the GBR during low tides [10] [28] [29]. Oxidation of  $\text{DMS}_a$  to sulfate aerosols then occurs (**Figure 1**) and these aerosols grow to form CN [51], precursors of CCN, potentially leading to increases in LLC cover which could increase net downwelling solar radiation and SSTs in the GBR [23] [54] [57], causing cooling and warming of SSTs during the summer season.

An important parameter that governs the relative scattering and absorption of solar radiation and thus the net downwelling solar radiation is the single scattering albedo (SSA), which refers to the ratio of scattering to the sum of scatter-

ing and absorption [58] [59]. Aerosols have a net negative surface-cooling effect when  $SSA < 0.95$ , and a net positive surface warming effect when  $SSA > 0.95$  [59] [60]. The net effect can change from a negative to larger positive forcing depending on the cloud fraction, surface albedo, and vertical aerosol concentration and cloud distribution [59]. Aerosols, in particular sulfates (from the oxidation of DMS, **Figure 1**), larger than  $\sim 0.05 \mu\text{m}$  provide the nuclei for most cloud droplets. Cloud droplets are the most important factor controlling the albedo of clouds and hence the temperature of our planet. Newly formed CN scatter radiation, but when they grow larger, acting as CCN ( $5 - 10 \mu\text{m}$ ), their scattering cross-section increases  $10^7$ . Under these conditions channeling effects occur where direct radiation is reflected by the clouds and solar UV radiation is enhanced. Such a situation has been observed in the Townsville region of the central GBR [48], in the wider GBR [61] and during coral bleaching at Keppel Islands and Heron Island in the central and southern GBR [54] [56], and could have accentuated the mass coral bleaching events in the GBR in 1998, 2002, and 2016 [40]. Low tides have a major and variable impact on emissions of  $\text{DMS}_a$  from coral reefs in the GBR [10] [28] [55] with oxidation forming variable sulfate aerosol concentrations that would affect the net downwelling solar radiation [44] [54]. Under light winds and low tides in summer at the Heron Island reef flat, daily net downwelling shortwave solar radiation can exceed  $+800 \text{ W}\cdot\text{m}^{-2}$ ; with up to 95% of the net radiation during the morning heating the water column and benthic coral cover [54]. These extreme solar radiation levels caused heating of the reef which was exacerbated by a mid-afternoon low tide when the shallow reef water reached  $34^\circ\text{C}$  and near-bottom temperatures  $33^\circ\text{C}$ , exceeding the thermal tolerance of corals at Heron Island, causing coral bleaching [54]. Enhanced scattering of solar radiation from nss-sulfate aerosols derived from oxidation of  $\text{DMS}_a$  produced from coral reefs, and associated cloud radiative effects, may have a bearing on the severity of coral bleaching events in the GBR [20] [63] and may reflect variation in the liquid water depth of the aerosols [62]. This net radiative forcing from clouds can be as high as four times as large as the radiative forcing from a doubling of  $\text{CO}_2$  levels in the atmosphere [57] and needs to be taken into account when ascribing coral bleaching events in the GBR solely to GHG warming.

## 6. Conclusion

Our research has provided preliminary evidence suggesting that variations in the flux of  $\text{DMS}_a$  from coral reefs in the GBR may enhance the scattering of solar radiation from nss-sulfate aerosols derived from oxidation of  $\text{DMS}_a$ , which could be responsible for the extreme solar radiation and SSTs observed during mass coral bleaching events. Changes in the LLC-SST relationship for the Magnetic Island fringing coral reef could indicate that the cooling and warming effects observed at this reef in summer may be due to changes in cloud radiative effects in the solar UV and thermal IR affecting downwelling radiation to the coral reef. These cloud radiative effects, as suggested by the different LLC-SST correlations,

may have a bearing on the increased frequency of mass coral bleaching events we are now seeing in the GBR as SSTs increase from global warming [37] [40] [63]. Global warming in combination with cloud radiative effects from aerosols produced from the coral reef marine environment may partly explain the mass coral bleaching in February 2016 when one third of coral reefs in the northern GBR experienced severe bleaching decimating hundreds of coral reefs. Clearly more in-depth studies are needed in the GBR on this coral reef-climate feedback in order to critically assess the factors that affect cooling and warming of inshore GBR waters from aerosols produced from the coral reefs and how this may affect the climate of northern Australia and the western Pacific.

## Acknowledgements

Thanks are extended to the staff of the Orpheus Island research station for their assistance and to Advan Muslim who collected samples from Nelly Bay reef. Dr Ray Berkelmans (AIMS) is thanked for providing the SST data from Magnetic Island. The completion of this research was assisted by a grant to Associate Professor Graham Jones by ARC Discovery Project grant (ARCDP) DP150101649 with Professor Zoran Ristovski of the Queensland Institute of Technology.

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